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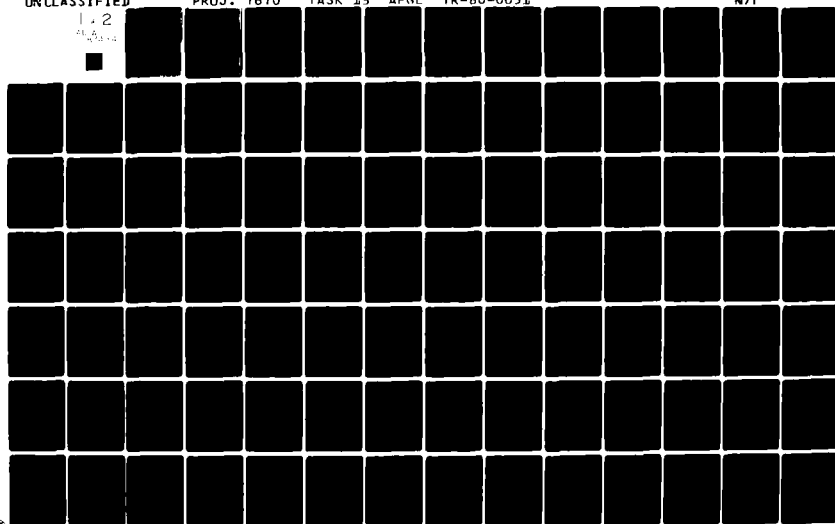
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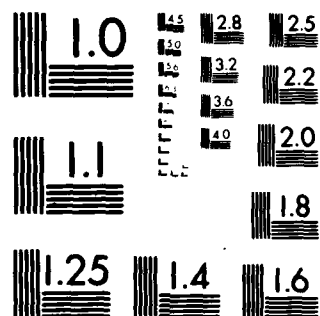
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**DEMONSTRATION OF THE RESEARCH NARRATIVE TRANSFER PROGRAM
FOR ARMY AIRBORNE: APPLICATIONS TO
AND SOUTHERN**

by

**W. H. Lee, Paul T. Ryan, Grant C. Aufderhaar
and Sun-Young Yeh**

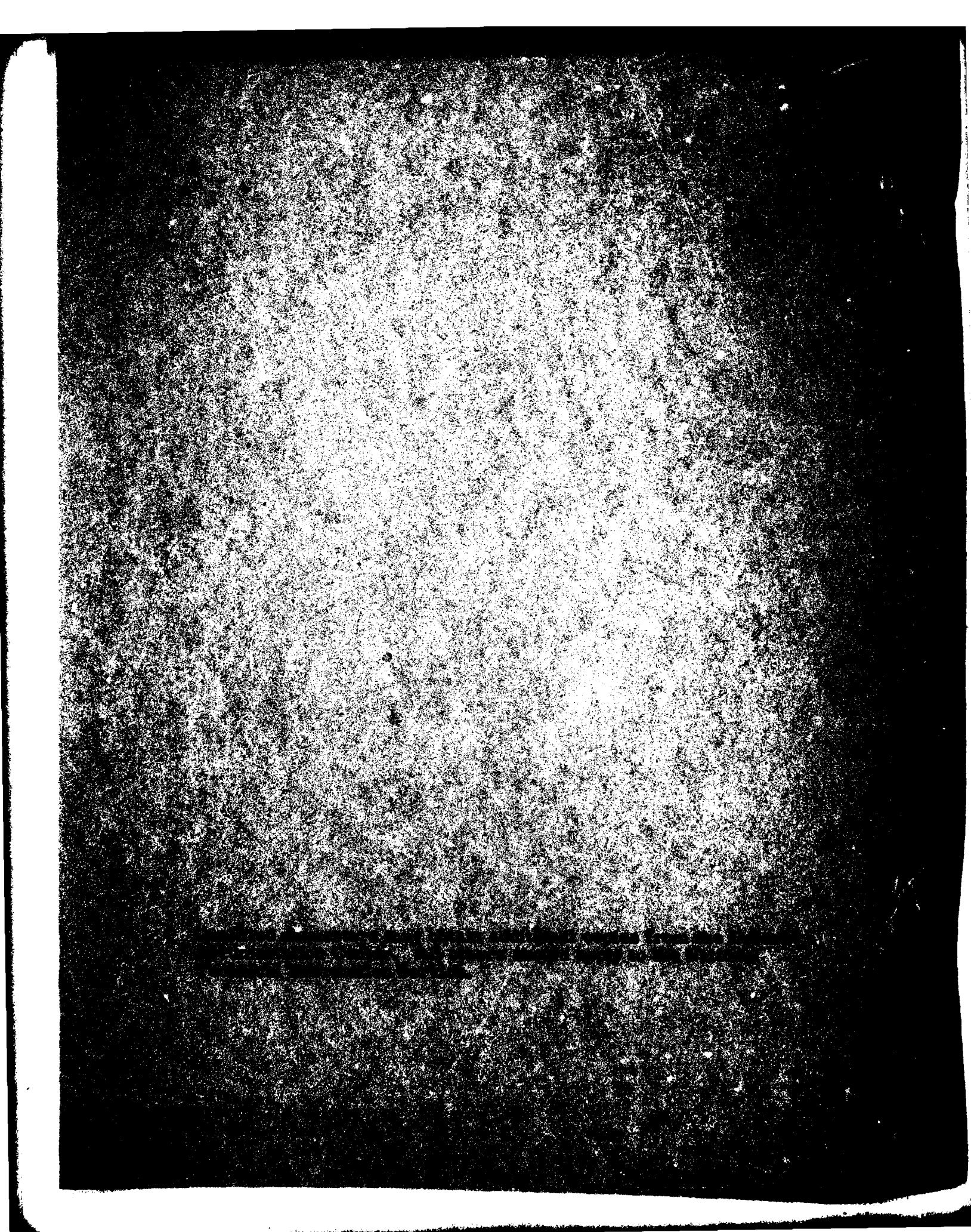
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ocean surfaces for two different atmospheric profiles. Over the land surface, it is shown that increasing the cloud liquid water content and cloud thickness significantly reduces the brightness temperatures for SSM/T channels 1 - 3. Over the ocean surface, however, the brightness temperatures for channels 1 - 3 are increased as the cloud liquid water content increases and reach to maximum values depending on the cloud thickness involved. Moreover, we show the importance of the cloud position and the atmospheric temperature profile on the brightness temperature values. In addition, investigation of the effects of precipitation on the temperature retrieval using both the theoretically simulated values and real SSM/T data has also been carried out on the basis of a statistical method. The hypothetical retrieval experiments reveal that the temperatures close to the surface suffer increased degradation as the rainfall rate increases. This finding is further supported by the analysis employing the real SSM/T data for a number of case studies in which temperature profiles from radiosondes are available for comparisons.

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SECTION 1

INTRODUCTORY REMARKS

The first microwave sounder intended for operational use was flown aboard the Air Force Defense Meteorological Satellite Program (DMSP) Block 5D satellite system launched in June, 1979. The microwave sensor (SSM/T) is a 7-channel scanning radiometer in the 50-60 GHz oxygen band region. The prime objective of SSM/T is to provide data for deriving global basis temperature profiles in the troposphere and lower stratosphere. Channel 1 (50.5 GHz) is a window channel responding strongly to the earth's surface characteristics, dense clouds, and precipitation. Moreover, the peaks of weighting functions for channels 2, 3, and 4 (53.2, 54.35, and 54.9 GHz) are below about 10 km.

It is known that the prime advantage of microwave over infrared temperature sounders is that the longer wavelength microwaves are much less affected by clouds and precipitation. However, to what degree and quantitative extent will clouds and precipitation influence the brightness temperature values which would have been observed under clear conditions? The objectives of the present project under the sponsorship of the Air Force Geophysics Laboratory have been (1) to develop a program for the transfer of microwave radiation in cloudy and precipitating atmospheres, (2) to compare the theoretical simulated microwave and infrared radiances with

observed values utilizing the required ground truth data when the DMSP microwave and infrared data become available, and (3) to investigate the inference of temperature profiles under all weather conditions.

Within the scope of these objectives and subject to the availability of the DMSP data, this final report first presents the theoretical foundation for the transfer of microwave radiation in clear and cloudy atmospheres. Applications of the microwave program to DMSP SSM/T channels are then described. Sensitivity analyses on the effects of clouds and precipitation on the upwelling brightness temperatures for channels 1 - 4 are subsequently carried out. In addition, we report the results of temperature profile retrieval exercises using the simulated brightness temperatures computed from the microwave radiative transfer program described previously. Lastly, a number of selected case studies utilizing the real SSM/T data recently made available to us are also carried out for the temperature retrieval exercises by means of a statistical method. These case studies cover the conditions for clear, cloudy and precipitating atmospheres.

SECTION 2

THEORY OF MICROWAVE RADIATIVE TRANSFER IN PLANE-PARALLEL ATMOSPHERES

2.1 Microwave Radiative Transfer in Clear Atmospheres

Consider a non-scattering, plane-parallel atmosphere which is in local thermodynamic equilibrium and assume that thermal radiation from the earth atmosphere is independent of the azimuthal angle ϕ , the general equation of transfer may be expressed in the frequency domain as

$$\mu \frac{d I_{\nu}(z, \mu)}{k_{\nu} \rho dz} = - I_{\nu}(z, \mu) + B_{\nu}[T(z)] , \quad (2.1)$$

where I_{ν} is the radiance at frequency ν , k_{ν} is the absorption coefficient, ρ is the density, $\mu = \cos \theta$, θ is the emergent angle, z is the height, and B_{ν} denotes the Planck function. For satellite sounding applications, the observations are normally taken close to the upwelling direction ($\mu \approx 1$). Thus, for the simplicity of discussions on the transfer of thermal radiation in a non-scattering atmosphere, the dependence on μ contained in the radiance expression is omitted.

The radiance solution of Eq. (2.1) at the top of the atmosphere in the upwelling direction is given by

$$I_{\nu}(\infty) = I_{\nu}(0) T_{\nu}(0, \infty) + \int_0^{\infty} B_{\nu}[T(z)] \frac{\partial T_{\nu}(z, \infty)}{\partial z} dz, \quad (2.2)$$

where $I_{\nu}(0)$ represents the radiance contribution from the surface,

and the transmittance $T_v(z, \infty)$ is expressed with respect to the top of the atmosphere in the form

$$T_v(z, \infty) = \exp\left[-\int_z^{\infty} k_v(z') \rho(z') dz'\right] . \quad (2.3)$$

In the microwave region, since the emissivity ϵ_v of the surface is normally less than unity, there will be a reflection contribution from the surface. The radiance emitted from the surface having a temperature T_s would therefore be given by

$$I_v(0) = \epsilon_v B_v(T_s) + (1-\epsilon_v) \int_{\infty}^0 B_v[T(z)] \frac{\partial T_v(0, z)}{\partial z} dz . \quad (2.4)$$

The first terms in the right-hand side of Eq. (2.4) denotes the surface emission contribution, whereas the second term represents the emission contribution from the entire atmosphere to the surface, which is reflected back to the atmosphere at the same frequency. The transmittance $T_v(0, z)$ is now expressed with respect to the surface.

Inserting the lower boundary condition into Eq. (2.2), we find

$$\begin{aligned} I_v(\infty) = & \epsilon_v B_v(T_s) T_v(0, \infty) + (1-\epsilon_v) T_v(0, \infty) \int_{\infty}^0 B_v[T(z)] \frac{\partial T_v(0, z)}{\partial z} dz \\ & + \int_0^{\infty} B_v[T(z)] \frac{\partial T_v(z, \infty)}{\partial z} dz . \end{aligned} \quad (2.5)$$

In the frequency domain, the Planck function is given by

$$B_v(T) = 2h \nu^3 / [c^2 (e^{h\nu/KT} - 1)] , \quad (2.6)$$

where h is the Planck constant, K is the Boltzmann constant, and c is the velocity of light. Moreover, in the microwave region $h\nu/KT \ll 1$,

the Planck function may be approximated by

$$B_{\nu}(T) \approx (2K\nu^2/c^2)T. \quad (2.7)$$

This is the Rayleigh-Jeans law, which states that the Planck radiance is linearly proportional to the temperature. Furthermore, radiometers which measure the thermal emission are usually calibrated with sources at certain reference temperatures. Thus, we may define an equivalent brightness temperature T_B such that

$$I_{\nu} = (2K\nu^2/c^2)T_B(\nu). \quad (2.8)$$

Substituting Eqs. (2.7) and (2.8) into Eq. (2.5), the solution of microwave radiative transfer may now be written in terms of temperature as follows:

$$\begin{aligned} T_B(\nu) = & \epsilon_{\nu} T_s T_{\nu}(0, \infty) + (1 - \epsilon_{\nu}) T_{\nu}(0, \infty) \int_0^{\infty} T(z) \frac{\partial T_{\nu}(0, z)}{\partial z} dz \\ & + \int_0^{\infty} T(z) \frac{\partial T_{\nu}(z, \infty)}{\partial z} dz. \end{aligned} \quad (2.9)$$

The transmittance is generally available with respect to the top of the atmosphere; i.e., $T_{\nu}(z) = T_{\nu}(z, \infty)$. Thus, for computational purposes, it is desirable to express $T_{\nu}(0, z)$ in terms of $T_{\nu}(z, \infty)$. For monochromatic frequencies, the transmittance is an exponential function of the optical depth as shown in Eq. (2.3).

Hence, we may write

$$\begin{aligned} T_{\nu}(0, z) &= \exp\left[-\int_0^z k_{\nu}(z') \rho(z') dz'\right] \\ &= \exp\left[-\int_0^{\infty} k_{\nu}(z') \rho(z') dz' + \int_z^{\infty} k_{\nu}(z') \rho(z') dz'\right] \\ &= T_{\nu}(0, \infty) / T_{\nu}(z, \infty). \end{aligned} \quad (2.10)$$

Moreover, since $T_v(0, \infty)$, the transmittance of the entire atmosphere, is a constant value, we also find

$$\frac{\partial T_v(0, z)}{\partial z} = - \frac{T_v(0, \infty)}{[T_v(z, \infty)]^2} \frac{\partial T_v(z, \infty)}{\partial z} . \quad (2.11)$$

Substituting Eq. (2.11) into Eq. (2.9), rearranging terms, and letting $T_v(z, \infty) = T_v(z)$, Eq. (2.9) may be rewritten to yield

$$T_B(v) = \epsilon_v T_s T_v(0) + \int_0^\infty J_v(z) \frac{\partial T_v(z)}{\partial z} dz , \quad (2.12)$$

where the atmospheric source term is given by

$$J_v(z) = \{1 + (1 - \epsilon_v) [T_v(0)/T_v(z)]^2\} T(z) . \quad (2.13)$$

2.2 Microwave Radiative Transfer in Scattering Cloudy Atmospheres

The basic equation of transfer for a plane-parallel cloud layer consisting of absorbing gases in local thermodynamic equilibrium may be written in the form

$$\begin{aligned} \mu \frac{d I_v(\tau, \mu)}{d\tau} = I_v(\tau, \mu) - \frac{\tilde{\omega}_v}{2} \int_{-1}^1 P_v(\mu, \mu') I_v(\tau, \mu') d\mu' \\ - (1 - \tilde{\omega}_v) B_v[T(\tau)] . \end{aligned} \quad (2.14)$$

In Eq. (2.14), I_v represents the monochromatic radiance of frequency ν , μ the cosine of the emergent angle with respect to the zenith, τ the optical depth for cloud particles and gases within the cloud, P_v the normalized axially symmetrical scattering phase function, T the cloud temperature, and the single scattering albedo

$$\tilde{\omega}_v = \frac{\beta_{s,v}}{\beta_{s,v} + \beta_{a,v} + k_v \rho} , \quad (2.15)$$

where $\beta_{s,v}$ and $\beta_{a,v}$ denote the volume scattering and absorption cross sections, respectively, for cloud particles at frequency v . The normalized phase function can be expressed as a Legendre polynomial of finite terms. By approximating the integration in Eq. (2.14) utilizing Gauss' quadrature formula, a set of first order inhomogeneous differential equations can be derived. Upon searching the homogeneous and particular solutions of the differential equations as outlined by Chandrasekhar (1950), the complete solution of the scattered intensity from an isothermal cloud having a temperature T_c at a given discrete stream can be expressed as

$$I_v(\tau, \mu_i) = \sum_m L_m \phi_m(\mu_i) \exp(-k_m \tau) + B_v(T_c) , \quad (2.16)$$

where m is the number of discrete streams employed, ϕ_m and k_m represent eigenfunctions and eigenvalues for the differential equations, and are associated with the scattering phase function and single-scattering albedo, and L_m are a set of constants of proportionality which can be determined from the radiation boundary conditions above and below the cloud layer.

In the previous subsection, we pointed out that in the microwave region, the Planck function is normally expressed in terms of the temperature and the measured radiance is given by an equivalent brightness temperature. Thus, using Eqs. (2.7) and (2.8), Eq. (2.14) may be written in terms of the brightness temperature as follows:

$$\mu \frac{dT_B(\nu, \tau; \mu)}{d\tau} = T_B(\nu, \tau; \mu) - \frac{\tilde{\omega}_\nu}{2} \int_{-1}^{+1} P_\nu(\mu, \mu') T_B(\nu, \tau; \mu') d\mu' - (1 - \tilde{\omega}_\nu) T(\tau) . \quad (2.17)$$

The solution of Eq. (2.17) can be derived in a manner similar to that of Eq. (2.14) and is given by

$$T_B(\nu, \tau; \mu_i) = \sum_m L'_m \phi_m(\mu_i) \exp(-k_m \tau) + T_c . \quad (2.18)$$

Here the constants of proportionality L'_m differ from those given in Eq. (2.16) and are to be evaluated from the radiation boundary conditions described below. Note that T_c , the cloud temperature, is independent of frequency ν .

The inhomogeneous and nonisothermal structure of a cloud layer may be taken into consideration by dividing the cloud layer into a number of sublayers, each of which has a mean isothermal temperature and is homogeneous with respect to the cloud composition. At the cloud top, the downward brightness temperature is equal to the brightness contributions from every point in the atmosphere above the cloud top. This can be expressed by

$$T_B(\nu, z_t; -\mu_i) = \int_{z=z_t}^{\infty} T(z) d\tau_\nu(z, z_t; -\mu_i) , \quad (2.19)$$

where z_t is the height at the cloud top and the negative sign on μ_i simply indicates downward transfer. Note here that we change the τ -coordinate to the height coordinate for the convenience of discussion. The transmittance in this equation is

$$T_v(z_t, z; \mu_i) = T_v(z, z_t; -\mu_i) = \exp\left[-\frac{1}{\mu_i} \int_{z_t}^z k_v(z') \rho(z') dz'\right], \quad (2.20)$$

and in terms of transmittance with respect to the top of the atmosphere, we find

$$\begin{aligned} T_v(z_t, z; \mu_i) &= T_v(z_t, \infty; \mu_i) / T_v(z, \infty; \mu_i) \\ &= T_v(z_t; \mu_i) / T_v(z; \mu_i), \quad z_t > z. \end{aligned} \quad (2.21)$$

Within the cloud layer, where scattering occurs, continuity of brightness values from all directions is required. Thus,

$$T_B(v, z_\ell; \mu_i) = T_B(v, z_{\ell+1}; \mu_i), \quad \ell=1, 2, \dots, N-1, \quad (2.22)$$

where N is the total number of sublayers within the cloud.

At the lower boundary of the cloud, three brightness contributions are immediately apparent. These include (a) the surface contribution, (b) the direct atmospheric contribution from below the cloud, and (c) the reflected atmospheric contribution from below the cloud. A fourth, and perhaps less obvious, brightness contribution at the lower boundary must be considered. Since the reflection of downward brightness by the earth's surface is significant for microwave radiation, the emergent brightness at the cloud bottom will contribute to the lower boundary condition. That is to say, the solution to the radiative transfer through the cloud affects the boundary conditions used to obtain the solution. This suggests that an iterative approach to the correct solution is required. Practically, this is accomplished by assuming initially that the top-down throughput

of the cloud equals one. The emergent intensity at the cloud bottom obtained by this assumption can then be attenuated by two trips through the atmosphere between the earth's surface and the cloud base and by the reflection at the earth's surface. This value is added to the lower boundary condition and the solution to the radiative transfer equation then produces a new value for the emergent brightness temperature at the cloud bottom. From this solution a new top/down throughput can be calculated and the process is repeated until the new throughput varies from the old by less than one tenth. In all cases examined by this study the iteration halted after only two steps. The lower boundary conditions can then be expressed as

$$\begin{aligned}
 T_B(\nu, z_b; +\mu_i) = & \epsilon_\nu T_s T_\nu(0, z_b; +\mu_i) + \int_{z=0}^{z_b} T(z) dT_\nu(z, z_b; +\mu_i) \\
 & + (1-\epsilon_\nu) T_\nu(0, z_b; +\mu_i) \left[\int_{z=0}^{z_b} T(z) dT_\nu(0, z; -\mu_i) \right. \\
 & \left. + T_B(\nu, z_t; -\mu_i) T_\nu^C(-\mu_i) T_\nu(z_b, 0; -\mu_i) \right], \quad (2.23)
 \end{aligned}$$

where $T_\nu^C(\mu_i)$ is defined to be the top/down throughput of the cloud for the stream defined by μ_i and z_b is the height of the cloud base. It is given by

$$T_\nu^C(\mu_i) = T_B(\nu, z_b; -\mu_i) / T_B(\nu, z_t; -\mu_i). \quad (2.24)$$

In Eq. (2.23), all the transmittances can be expressed in reference to the top of the atmosphere. The numerical technique for solving a set of simultaneous linear equations denoted in Eqs. (2.19), (2.22) and (2.23) has been well documented in our previous infrared studies

(Liou et al., 1977; Feddes and Liou, 1977).

The microwave transfer program for cloud layers will give the upward brightness temperature at the cloud top, i.e., $T_B(\nu, z_t; \mu_i)$. Thus, the upward brightness temperature at the satellite point of view in overcast cloudy conditions may then be written as

$$T_B(\nu, \infty; \mu_i) = T_B(\nu, z_t; \mu_i) T_V(z_t, \infty; \mu_i) + \int_{z_t}^{\infty} T(z) dT_V(z, \infty; \mu_i) , \quad (2.25)$$

where the first term in the right-hand side of Eq. (2.25) represents the contribution from the cloud top brightness temperature which is being attenuated to the top of the atmosphere, and the remaining term is the atmospheric contribution above the cloud top.

SECTION 3

APPLICATIONS OF MICROWAVE RADIATIVE TRANSFER PROGRAM TO DMSP BLOCK 5D SSM/T CHANNELS

3.1 Characteristics of DMSP Block 5D SSM/T Channels

The Passive Microwave Temperature Sounder (SSM/T) was launched in June, 1979, in a sun-synchronous polar orbit, by the United States Air Force as part of the Defense Meteorological Satellite Program (DMSP) Block 5D package. Other equipment of meteorological interest on board the space craft include visual and infrared imagery channels and a scanning infrared spectroradiometer. The SSM/T sensor is a cross track scanning radiometer, which acquires data at 32 second intervals and at seven angular positions separated by 12 degrees. The "footprints" of a scan for the SSM/T along the satellite subtrack are shown in Figure 1. The horizontal resolution is a near circle of 174 km diameter in the nadir direction, while it is an ellipse with a major axis of 304 km and minor axis of approximately 213 km at the maximum scan angle of 36° from nadir. Seven operational frequencies were chosen in the vicinity of a strong oxygen absorption band in the 50-60 GHz region. Figure 2 depicts the weighting functions ($\partial T/\partial z$) of the SSM/T channels in the nadir direction for a surface emissivity of 0.97. Channel 1, 50.50 GHz, is a "window" channel and senses at or near the earth's surface. Channels 2, 3, and 4 (53.20, 54.35 and 54.90 GHz) have weighting functions that

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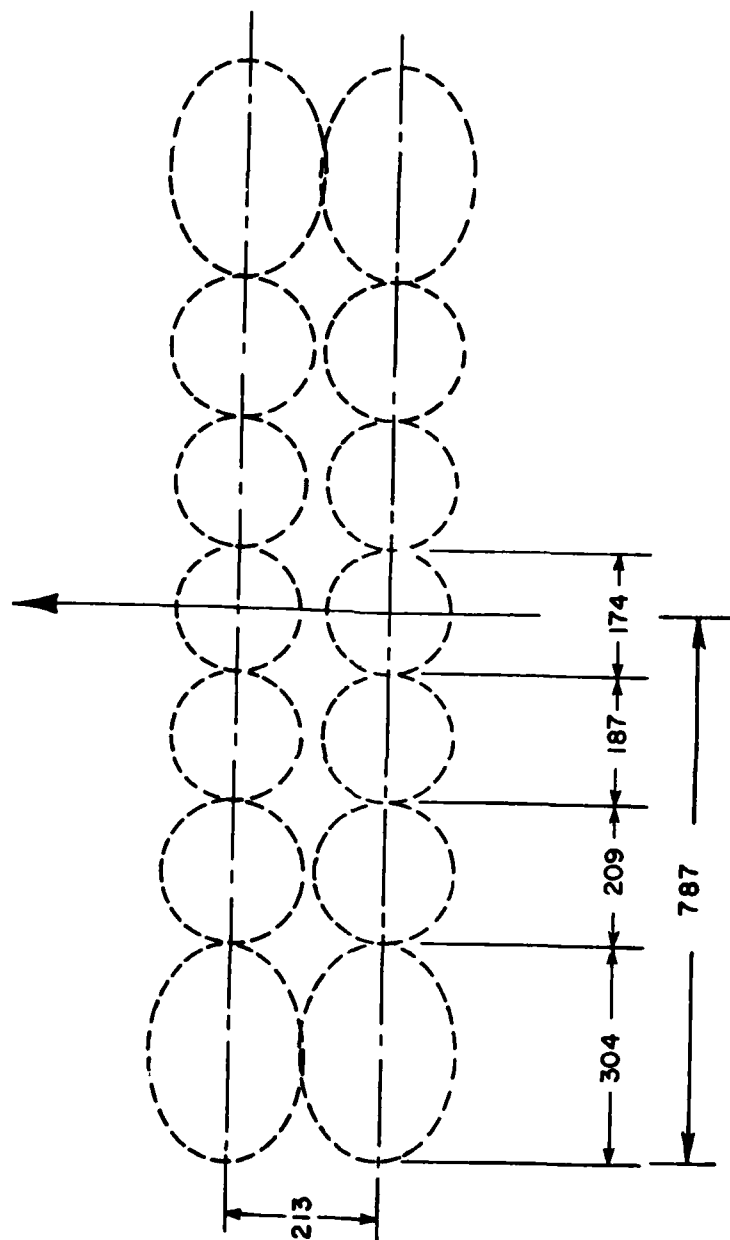


Figure 1. The scan pattern of SSM/T.

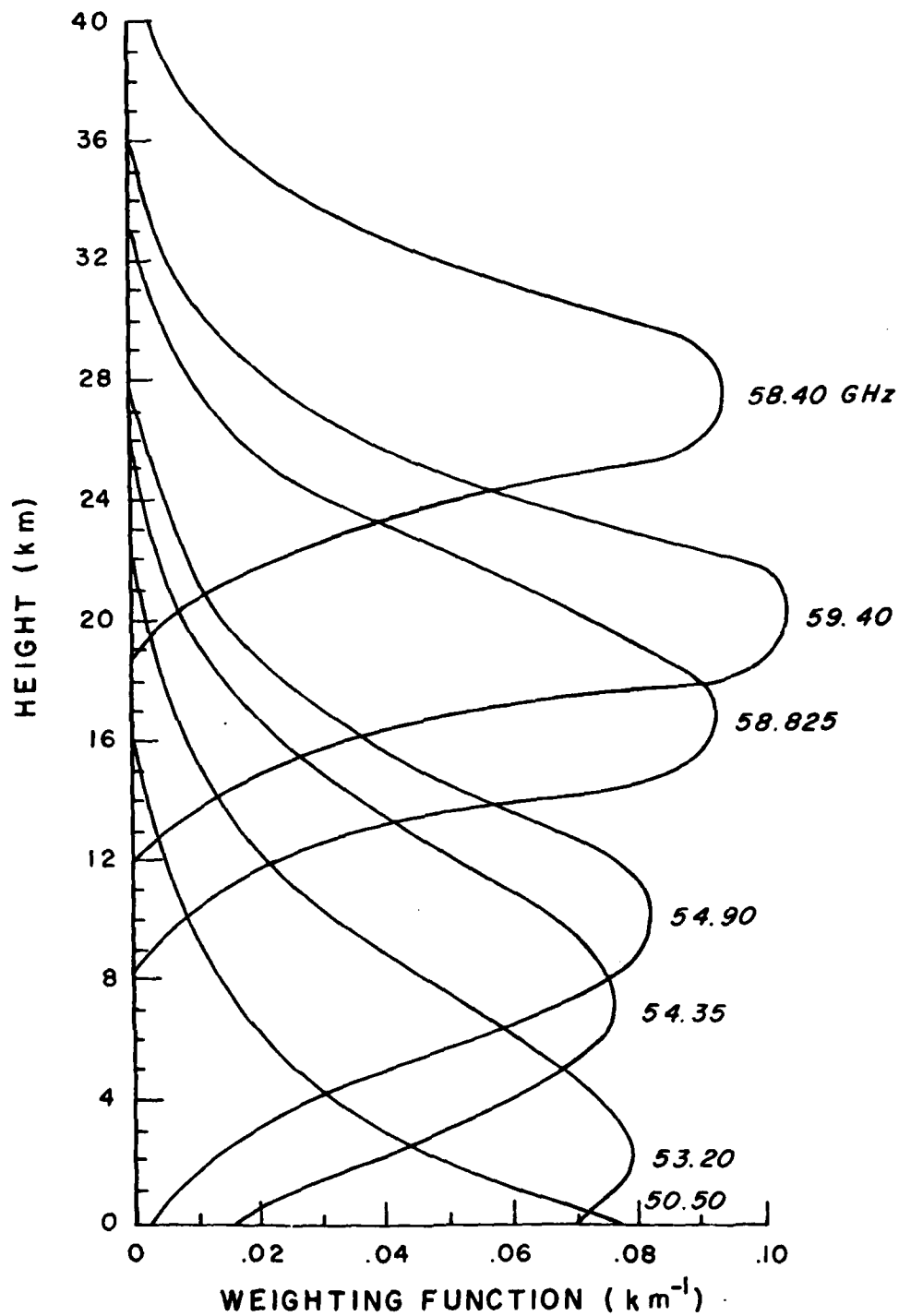


Figure 2. SSM/T weighting functions (nadir) with antenna gain characteristics included for a surface emissivity of 0.97.

peak at increasing heights in the troposphere. The weighting functions of Channels 5, 6 and 7 peak in the stratosphere with Channel 5 peaking highest in the atmosphere near 28 km. The weighting functions result from consideration of atmospheric absorption due to water vapor and molecular oxygen. The peaks of these weighting functions indicate the approximate location in the atmosphere from which most of the energy that reaches the top of the atmosphere originates. The weighting functions presented in Figure 2 also include the transmittance corrections for antenna gain characteristics. The DMSP Block 5D satellite also carries channels in the infrared. The Special Sensor H (SSH) is the infrared temperature-humidity-ozone sounder consisting of six channels in the $15\text{ }\mu\text{m}$ CO_2 band, eight channels in the $18\text{-}30\text{ }\mu\text{m}$ H_2O rotational band, and one channel in the $9.6\text{ }\mu\text{m}$ O_3 band.

3.2 Atmospheric Absorption and Scattering Parameters for SSM/T Frequencies

Microwave frequencies between 50.0 and 60.0 GHz are affected by several different atmospheric constituents. Gases in the atmosphere absorb and emit microwave radiation. Cloud droplets and rain drops not only absorb and emit but also scatter the incoming radiation.

3.2.1 Absorption due to gaseous constituents. The primary gaseous absorbers in the microwave region are H_2O and O_2 . Atmospheric nitrogen gradually becomes a more significant absorber at frequencies greater than 120 GHz, but has little influence in the vicinity of 60 GHz. Other gaseous constituents occur in very low concentrations and, in general, their effects are very small.

The O_2 absorption spectrum in the microwave region is due to the molecule's magnetic dipole moment. Figure 3 presents the absorption coefficient of dry air versus frequency. The absorption coefficient (decibels/km) shows a dramatic peak near 60 GHz and a secondary maximum near 120 GHz. Both peaks are attributable to molecular oxygen. The characteristics of this absorption spectrum were first investigated by Van Vleck (1947). Other investigators have built on his studies, culminating in a paper by Meeks and Lilley (1963). For this study, the absorption coefficients for O_2 were calculated using Meeks and Lilley's parameters.

Water vapor absorption in the microwave is chiefly due to an electric dipole moment. Initial studies were published by Van Vleck (1947), Becker and Autler (1946), and King, et al. (1947). Rosenblum (1961) gave an excellent summary of work done up to 1961. Figure 4 presents water vapor absorption (decibels/km) versus frequency in the microwave region. It is seen that the values of absorption increase gradually through the range of frequencies under investigation. It is apparent that the absorption due to water vapor is less than that of molecular oxygen for all the channels chosen for the SSM/T sounder. Note, however, that even though the absorption coefficients for O_2 are generally higher than those for H_2O , the two values are roughly equivalent for the window channel at 50.50 GHz. H_2O absorption coefficients were calculated using the method of Barrett and Chung (1962). Computer programs for determination of both of these coefficients were kindly provided by Falcone (personal communication).

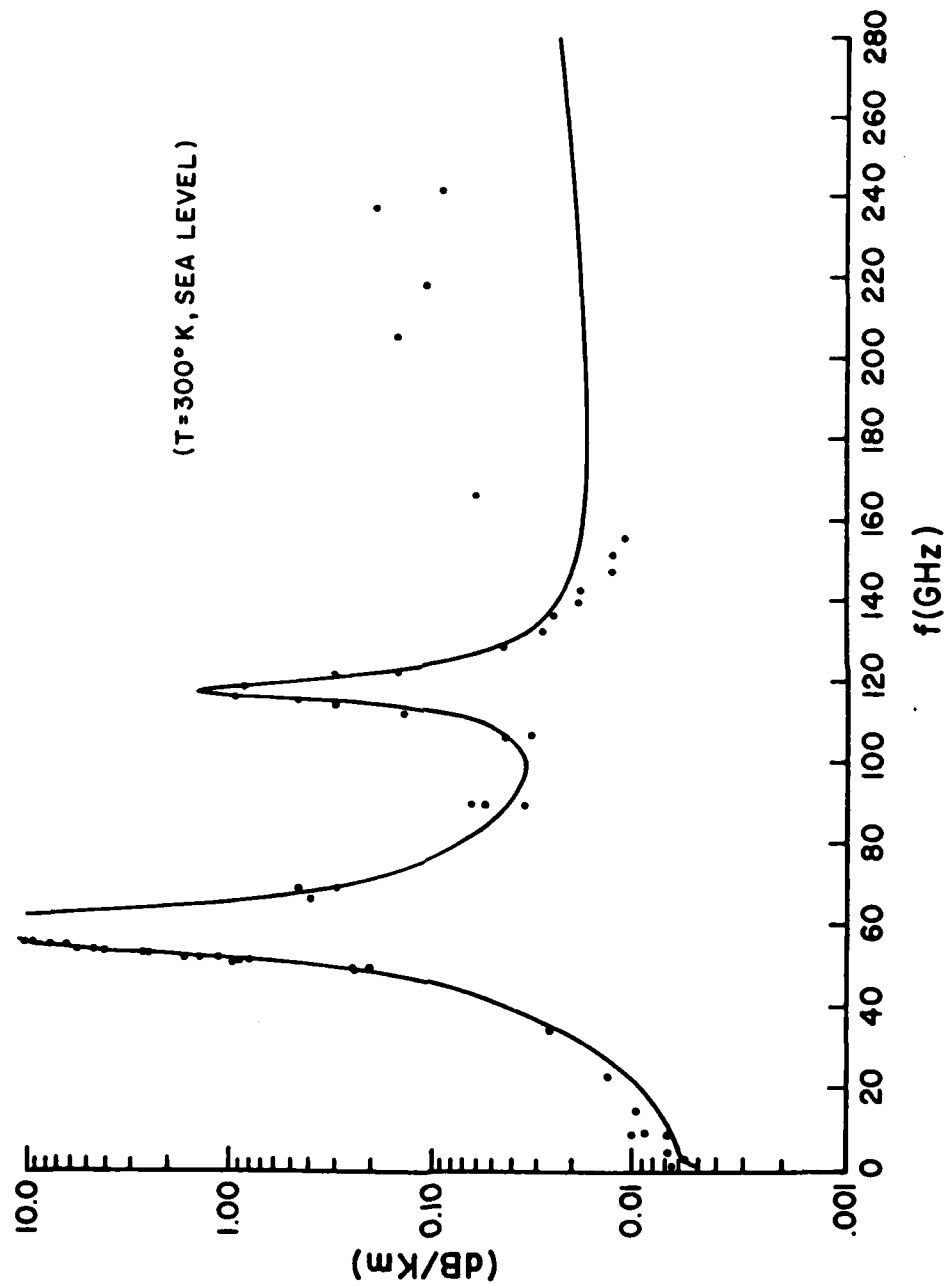


Figure 3. Absorption coefficient (decibels/km) of dry air as a function of frequency (after Meeks and Lilley, 1963). The dots are measured points by Hans Leibe during the period from 1974 to 1976 (courtesy of Rigone).

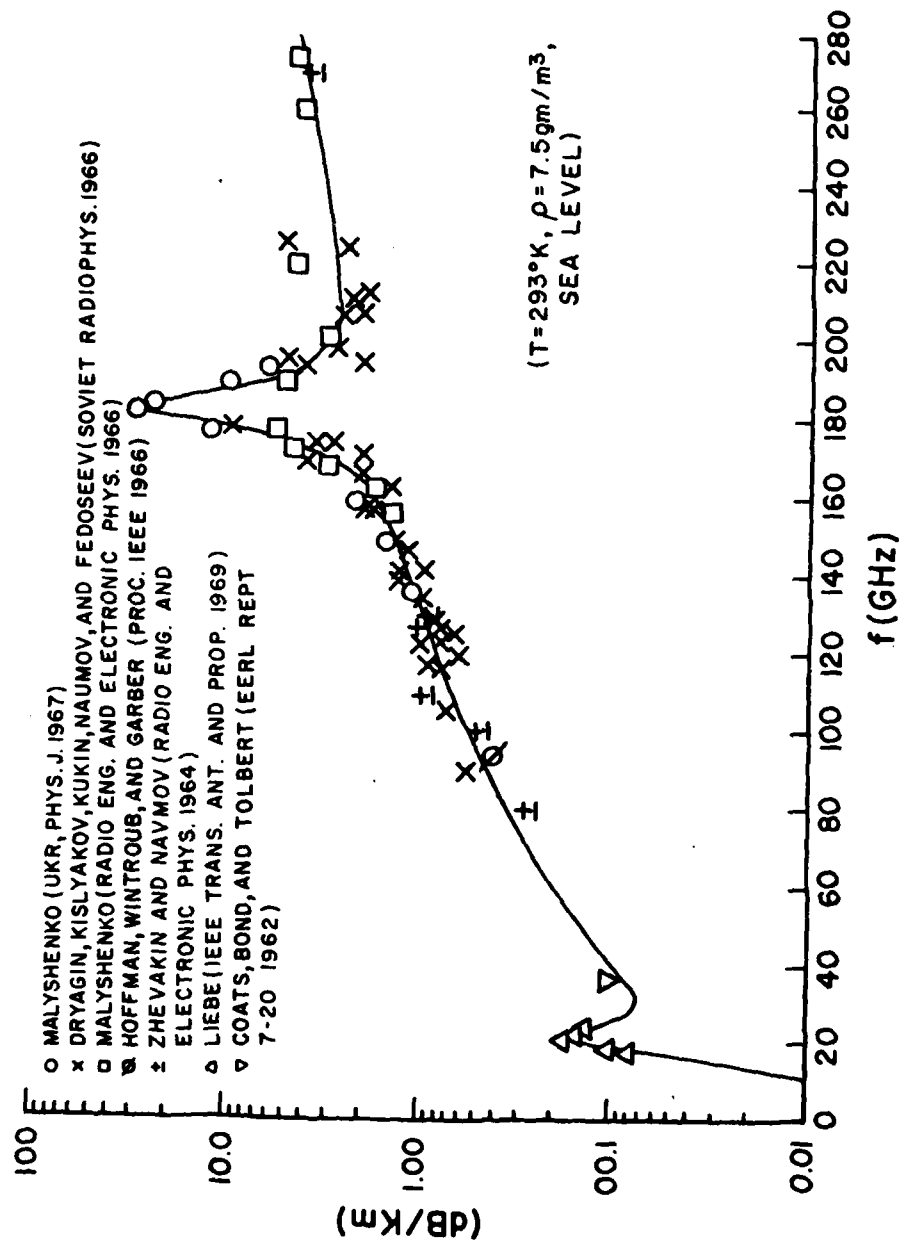


Figure 4. Water vapor absorption (decibels/km) as a function of frequency in the microwave region.

3.2.2 Scattering and absorption due to cloud droplets. The application of Mie scattering theory enables the calculation of optical parameters for assumed spherical particles given the complex index of refraction, wavelength of incident radiation, and particle size distribution. Although larger rain drops are known to deviate somewhat from the spherical shape, Mie theory still remains probably the best currently available theoretical tool for analyzing the effects of cloud droplets and rain in the atmosphere.

The Mie program used for the calculations in this study is a modification for Univac computers of the program by Liou and Hansen (1971). The output of the Mie computations include the coefficients of the Legendre polynomial approximation to the phase function, the extinction coefficient (km^{-1}) and the single scattering albedo. Note that the extinction coefficient and single scattering albedo determined here involve scattering and absorption by the particles as the only mechanisms for extinction. These values are modified to include absorption by the gaseous atmosphere within the cloud prior to their use in the solution of the radiative transfer equation.

The complex indices of refraction for pure liquid water at temperatures of $+10^\circ$ and -10° C have been determined by Savage (1976) for several different frequencies. These values (Table 1) are based on a least squares fit to the data of Saxton and Lane (1962) given by Hollinger (1973). The complex indices of refraction for the SSM/T channels as shown in Table 2 were obtained from Table 1 by linear interpolation in frequency. For the purpose of this study the values for 0° C are used throughout.

Table 1. Refractive index $m=m_r-im_i$ of pure liquid water.

Frequency (GHz)	T=10°C		T=0°C		T=-10°C	
	m_r	m_i	m_r	m_i	m_r	m_i
19.35	5.87	2.96	5.20	2.94	4.64	2.81
37.00	4.32	2.60	3.83	2.37	3.46	2.12
50.30	3.76	2.29	3.37	2.03	3.08	1.78
89.50	3.00	1.67	2.76	1.42	2.60	1.20
100.00	2.89	1.56	2.68	1.31	2.54	1.10
118.00	2.75	1.39	2.58	1.15	2.46	0.96
130.00	2.68	1.30	2.53	1.07	2.43	0.89
183.00	2.49	1.00	2.39	0.81	2.33	0.66
231.00	2.40	0.82	2.33	0.66	2.29	0.53

Table 2. Refractive index of pure liquid water (SSM/T).

	Channel Frequencies (GHz)						
	50.50	53.20	54.35	54.90	58.40	58.82	59.40
m_r	3.37	3.32	3.31	3.30	3.24	3.24	3.23
m_i	2.03	1.98	1.97	1.96	1.90	1.90	1.89

Utilizing the wavelength and the associated index of refraction, Mie scattering calculations for single particles may be carried out. Once the absorption and scattering efficiencies for single particles have been determined in terms of effective cross sectional area, the efficiencies for a size distribution of particles can be determined by integrating over the distribution. This area per volume relationship then has dimensions of inverse length and is described as the absorption or scattering coefficient. The extinction coefficient is then the sum of the absorption and scattering coefficients. The drop size distributions used for this study are necessarily numerical models. Two models have been chosen to represent nonprecipitating and precipitating clouds, respectively. Deirmendjian's L-Model cloud (Deirmendjian, 1969) drop size distribution is used for the nonprecipitating case. It is given by

$$N(r) = ar^{\alpha} \exp(-br^{\delta}) , \quad (3.1)$$

where r is the drop radius in μm and note that $N(r)$ is in units of cm^{-4} . By varying the constants this model can be made to represent a variety of cloud types. For this study, $a=4.975 \times 10^7$, $\alpha=2$, $b=15.1186$, and $\delta=.5$. The selection of the nonprecipitating cloud model here is hypothetical and we fully realize that it is not a representative cloud model for typical clouds occurring in the atmosphere which have normally insignificant effect in the microwave frequencies. For the precipitating case, a theoretical drop size distribution based on rainfall rate is used (Marshall and Palmer, 1948). The distribution is expressed by

$$N(r) = 0.16 \exp(-2r\delta) , \quad (3.2)$$

where $\delta = 41 R^{-.21}$ and R denotes rainfall rate (mm/hr.). This exponential behavior of drop size distribution has been experimentally verified by Gunn and Marshall (1958) and Sekhon and Srivastava (1970).

The phase function in terms of Legendre polynomial coefficients, extinction coefficients, and single scattering albedos calculated for each SSM/T channel from Mie theory are tabulated in Tables 3-13. Since the extinction coefficient input to the radiative transfer program must be in units of (km^{-1}), the final values output by the Mie program have been multiplied by the proper constant to obtain these units.

3.3 Atmospheric Profile and Cloud Models Used

Two climatological profiles were used for the sensitivity study. The first profile was derived from the northern hemispheric mid-latitude Spring/Fall climatology contained in the U.S. Standard Atmospheric Supplements, 1966. The second profile was derived from the climatology for 30° north latitude, July, contained in the same source. For both profiles temperature and height values were interpolated to 40 standard pressure levels used for this study. Mixing ratio values for levels below 250 mb were obtained from a Skew-T diagram and hence are saturation values. Above 250 mb a constant mixing ratio of 3 parts per million was used. The profiles for the mid-latitude Spring/Fall and the 30° N latitude, July, are depicted in Figure 5.

To account for the total water within the cloud, the vertical liquid water content is customarily used. For spherical particles the droplet volume per unit volume of atmosphere is given by

Table 3. Optical parameters (Deirmendjian L-Model)

Freq(GHz)	$\bar{\omega}_0$	$\bar{\omega}_1$	$\bar{\omega}_2$	$\bar{\omega}_3$	$\bar{\omega}_4$	$\bar{\omega}_5$
	$\bar{\omega}_6$	$\bar{\omega}_7$	$\bar{\omega}_8$	$\bar{\omega}_9$	$\bar{\omega}_{10}$	$\bar{\omega}_{11}$
	$\bar{\omega}_{12}$	$\bar{\omega}_{13}$	$\bar{\omega}_{14}$	$\bar{\omega}_{15}$	β_{ext}	$\bar{\omega}_v$
50.50 GHz						
.10000000E+1	.33071515E+0	.49360468E+0	.11251515E+0	-.75388403E-2	-.64620544E-2	
-.61193558E-1	-.19706414E-1	-.72342400E-1	-.21862288E-1	-.73393241E-1	-.21055428E-1	
-.56718646E-1	-.18075541E-1	-.53308599E-1	-.13297797E-1	.19639240E+0	.371442	
53.20 GHz						
.10000000E+1	.36502548E+0	.50176195E+0	.12244376E+0	-.16683441E-2	-.59117595E-2	
-.60940647E-1	-.21526171E-1	-.72936066E-1	-.24082668E-1	-.74087993E-1	-.23228360E-1	
-.67363482E-1	-.19934813E-1	-.53830418E-1	-.14667078E-1	.21121971E+0	.378722	
54.35 GHz						
.10000000E+1	.37639293E+0	.50474406E+0	.12667598E+0	.92690317E-3	-.54816648E-2	
-.60773288E-1	-.22136718E-1	-.73163086E-1	-.24870594E-1	-.74365129E-1	-.24001107E-1	
-.67622238E-1	-.20600170E-1	-.54040190E-1	-.15157736E-1	.21741939E+0	.382306	
54.90 GHz						
.10000000E+1	.38338237E+0	.50662057E+0	.12879932E+0	.22010716E-2	-.53178290E-2	
-.60704117E-1	-.22500922E-1	-.73293393E-1	-.25328362E-1	-.74521266E-1	-.24448971E-1	
-.67767713E-1	-.20985495E-1	-.54157977E-1	-.15441639E-1	.22034972E+0	.383564	
58.40 GHz						
.10000000E+1	.42759183E+0	.51955854E+0	.14307608E+0	.10864534E-1	-.39435839E-2	
-.60144436E-1	-.24772093E-1	-.74167010E-1	-.28268136E-1	-.75591876E-1	-.27337488E-1	
-.68768864E-1	-.23473136E-1	-.54969336E-1	-.17275603E-1	.23852110E+0	.391119	
58.82 GHz						
.10000000E+1	.43064151E+0	.52059478E+0	.14467422E+0	.11900079E-1	-.36744970E-2	
-.60041932E-1	-.24929679E-1	-.74243871E-1	-.28504663E-1	-.75695095E-1	-.27573876E-1	
-.68966560E-1	-.23677469E-1	-.55048812E-1	-.17426715E-1	.24066970E+0	.392445	
59.40 GHz						
.10000000E+1	.43805500E+0	.52296755E+0	.14716554E+0	.13424878E-1	-.33915394E-2	
-.59527715E-1	-.25301872E-1	-.74397610E-1	-.29002240E-1	-.75888122E-1	-.28065275E-1	
-.69047813E-1	-.24101133E-1	-.55195863E-1	-.17739194E-1	.24356932E+0	.393526	

Table 4. Optical parameters (Marshall-Palmer 1mm/hr)

Freq(GHz)	$\bar{\omega}_0$	$\bar{\omega}_1$	$\bar{\omega}_2$	$\bar{\omega}_3$	$\bar{\omega}_4$	$\bar{\omega}_5$
	$\bar{\omega}_6$	$\bar{\omega}_7$	$\bar{\omega}_8$	$\bar{\omega}_9$	$\bar{\omega}_{10}$	$\bar{\omega}_{11}$
	$\bar{\omega}_{12}$	$\bar{\omega}_{13}$	$\bar{\omega}_{14}$	$\bar{\omega}_{15}$	ρ_{ext}	$\bar{\omega}_v$
50.50GHz						
.1000000E+1	.19833052E+0	.47688411E+0	.55118323E-1	-.39233125E-1	-.87157051E-2	
-.62096694E-1	-.11531642E-1	-.70222380E-1	-.12313806E-1	-.71009004E-1	-.11824920E-1	
-.64517421E-1	-.10137258E-1	-.51524593E-1	-.74482281E-2	.84232398E-1	.292315	
53.20GHz						
.10000000E+1	.22402792E+0	.47839620E+0	.59819801E-1	-.37492504E-1	-.95918234E-2	
-.62194212E-1	-.12922219E-1	-.70374822E-1	-.13802233E-1	-.71164137E-1	-.13254064E-1	
-.64659236E-1	-.11362062E-1	-.51639036E-1	-.83475256E-1	.92745402E-1	.304235	
54.35GHz						
.10000000E+1	.23205810E+0	.47878575E+0	.61910084E-1	-.36678740E-1	-.98396532E-2	
-.62242228E-1	-.13385122E-1	-.70430895E-1	-.14298761E-1	-.71221342E-1	-.13730905E-1	
-.64711583E-1	-.11770879E-1	-.51681400E-1	-.86479728E-1	.96396855E-1	.309626	
54.90GHz						
.10000000E+1	.23750060E+0	.47918551E+0	.62929877E-1	-.36273642E-1	-.10013095E-1	
-.62247228E-1	-.13680903E+0	-.70468713E-1	-.14615722E-1	-.71259849E-1	-.14035253E-1	
-.64746761E-1	-.12031728E-1	-.51709807E-1	-.88395184E-2	.98139872E-1	.311786	
50.40GHz						
.10000000E+1	.27220688E+0	.48214605E+0	.69720240E-1	-.33448139E-1	-.11069380E-1	
-.62405476E-1	-.15579415E-1	-.70740726E-1	-.16653212E-1	-.71537185E-1	-.15991775E-1	
-.65000162E-1	-.13708748E-1	-.51914319E-1	-.10071212E-1	.10920593E+0	.324571	
58.92GHz						
.10000000E+1	.27427582E+0	.48225550E+0	.70554734E-1	-.33082321E-1	-.11109789E-1	
-.62412982E-1	-.15711972E-1	-.70762828E-1	-.16796611E-1	-.71559851E-1	-.16129550E-1	
-.65020926E-1	-.13826951E-1	-.51931131E-1	-.10158213E-1	.11054591E+0	.326468	
59.40GHz						
.10000000E+1	.28023853E+0	.48286078E+0	.71740599E-1	-.32564901E-1	-.11279653E-1	
-.62442334E-1	-.16038949E-1	-.70815697E-1	-.17148129E-1	-.71613804E-1	-.16467157E-1	
-.65070226E-1	-.14116323E-1	-.51970903E-1	-.10370743E-1	.11236057E+0	.328350	

Table 5. Optical parameters (Marshall-Palmer 2mm/hr)

Freq(GHz)	$\bar{\omega}_1$	$\bar{\omega}_2$	$\bar{\omega}_3$	$\bar{\omega}_4$	$\bar{\omega}_5$
$\bar{\omega}_0$					
$\bar{\omega}_6$					
$\bar{\omega}_{12}$					
50.50GHz					
.10000000E+1	.23305293E+0	.47669795E+0	.69520500E-1	-.33475318E-1	-.95957994E-2
-.62175845E-1	-.13776946E-1	-.70478158E-1	-.14730295E-1	-.71271311E-1	-.14146668E-1
-.64758072E-1	-.12129186E-1	-.51720168E-1	-.89139807E-2	.17904705E+0	.333189
53.20GHz					
.10000000E+1	.26325986E+0	.47933453E+0	.75807102E-1	-.30618404E-1	-.10409493E-1
-.62303234E-1	-.15445792E-1	-.70736236E-1	-.16526853E-1	-.71535191E-1	-.15872090E-1
-.64999268E-1	-.13608205E-1	-.51914898E-1	-.10000476E-1	.19526609E+0	.343557
54.35GHz					
.10000000E+1	.27304477E+0	.48016100E+0	.78578498E-1	-.29300388E-1	-.10608633E-1
-.62339426E-1	-.16019093E-1	-.70836625E-1	-.17147537E-1	-.71638278E-1	-.16468370E-1
-.65093569E-1	-.14119549E-1	-.51991134E-1	-.10376438E-1	.20213311E+0	.348307
54.90GHz					
.10000000E+1	.27938366E+0	.48085325E+0	.79948634E-1	.28644508E-1	-.10758894E-1
-.62367953E-1	-.16371191E-1	-.70899995E-1	-.17527982E-1	-.71703261E-1	-.16833752E-1
-.65152969E-1	-.14432824E-1	-.52039092E-1	-.10606620E-1	.20538624E+0	.350154
58.40GHz					
.10000000E+1	.31965204E+0	.48595885E+0	.89119851E-1	-.24116337E-1	-.11591706E-1
-.62553437E-1	-.18625856E-1	-.71353265E-1	-.19972884E-1	-.72168922E-1	-.19182567E-1
-.65578578E-1	-.16446644E-1	-.52382578E-1	-.12086596E-1	.22574856E+0	.360992
58.82GHz					
.10000000E+1	.32229704E+0	.48625726E+0	.90221748E-1	-.23540332E-1	-.11595119E-1
-.62559037E-1	-.18795815E-1	-.71394044E-1	-.20160143E-1	-.72211288E-1	-.19362601E-1
-.65617340E-1	-.16601176E-1	-.52413899E-1	-.12200346E-1	.22818576E+0	.362670
59.40GHz					
.10000000E+1	.32915882E+0	.48728307E+0	.91830136E-1	-.22720526E-1	-.11714991E-1
-.62591200E-1	-.19181183E-1	-.71480898E-1	-.20579983E-1	-.72300772E-1	-.19766032E-1
-.65669129E-1	-.16947077E-1	-.52479871E-1	-.12454616E-1	.23146830E+0	.364242

Table 6. Optical parameters (Marshall-Palmer 3mm/hr)

Freq(GHz)	$\bar{\omega}_0$	$\bar{\omega}_1$	$\bar{\omega}_2$	$\bar{\omega}_3$	$\bar{\omega}_4$	$\bar{\omega}_5$
	$\bar{\omega}_6$	$\bar{\omega}_7$	$\bar{\omega}_8$	$\bar{\omega}_9$	$\bar{\omega}_{10}$	$\bar{\omega}_{11}$
	$\bar{\omega}_{12}$	$\bar{\omega}_{13}$	$\bar{\omega}_{14}$	$\bar{\omega}_{15}$	β_{ext}	$\bar{\omega}_v$
50.50GHz						
.10000000E+1	.25862318E+0	.47744717E+0	.79897807E-1	-.28583795E-1	-.99653228E-1	-.15858597E-1
-.62224346E-1	-.15418045E-1	-.70739437E-1	-.16511622E-1	-.71540767E-1	-.355154	
-.65005033E-1	-.13597997E-1	-.51920506E-1	-.99950262E-1	.27303706E+0		
53.20GHz						
.10000000E+1	.29162371E+0	.48117814E+0	.8782297E-1	-.24840494E-1	-.10646886E-1	-.17782016E-1
-.62359262E-1	-.17264665E-1	-.71090662E-1	-.18513792E-1	-.71901866E-1	-.364537	
-.65335110E-1	-.15247116E-1	-.52186991E-1	-.11206979E-1	.29677251E+0		
54.35GHz						
.10000000E+1	.30252529E+0	.48244693E+0	.90664385E-1	-.23128135E-1	-.10770838E-1	-.18460047E-1
-.62392209E-1	-.17909504E-1	-.71230216E-1	-.19219201E-1	-.72046216E-1	-.368894	
-.65467236E-1	-.15288707E-1	-.52293727E-1	-.11634725E-1	.30640090E+0		
54.90GHz						
.10000000E+1	.30941458E+0	.48341050E+0	.92299544E-1	-.22275960E-1	-.10883022E-1	-.18866422E-1
-.62420361E-1	-.18297148E-1	-.71315535E-1	-.19642109E-1	-.72134340E-1	.370550	
-.65547707E-1	-.16177158E-1	-.52358718E-1	-.11890848E-1	.31094508E+0		
58.40GHz						
.10000000E+1	.35308196E+0	.49045007E+0	.10325306E+0	-.16426243E-1	-.11419156E-1	-.21478142E-1
-.62589478E-1	-.20771828E-1	-.71920656E-1	-.22359124E-1	-.72761442E-1	.380231	
-.66121005E-1	-.18416918E-1	-.52821356E-1	-.13537574E-1	.33915074E+0		
58.82GHz						
.10000000E+1	.35609849E+0	.49093359E+0	.10455210E+0	-.15689422E-1	-.11379821E-1	-.21687536E-1
-.62588423E-1	-.20965118E-1	-.71977381E-1	-.22576647E-1	-.72821073E-1	.381782	
-.66175624E-1	-.18596657E-1	-.52865533E-1	-.13669973E-1	.34249930E+0		
59.40GHz						
.10000000E+1	.36350373E+0	.49232587E+0	.10647667E+0	-.14636938E-1	-.11438722E-1	-.22135110E-1
-.62614921E-1	-.21385745E-1	-.72091956E-1	-.23042074E-1	-.72940414E-1	.383173	
-.66284733E-1	-.18980559E-1	-.52953556E-1	-.13952286E-1	.34700110E+0		

Table 7. Optical parameters (Marshall-Palmer 4mm/hr)

Freq(GHz)	$\bar{\omega}_1$	$\bar{\omega}_2$	$\bar{\omega}_3$	$\bar{\omega}_4$	$\bar{\omega}_5$
$\bar{\omega}_0$					
$\bar{\omega}_6$					
$\bar{\omega}_{12}$					
50.50GHz					
.10000000E+1	.27936264E+0	.47868582E+0	.88328239E-1	-.24241288E-1	-.10081960E-1
-.62247414E-1	-.16744915E-1	-.70998671E-1	-.17965453E-1	-.71809772E-1	-.17256226E-1
-.65251498E-1	-.14797347E-1	-.52120194E-1	-.10878014E-1	.36732478E+0	.369816
53.20GHz					
.10000000E+1	.31439899E+0	.48345105E+0	.96808990E-1	-.19747665E-1	-.10603425E-1
-.62374929E-1	-.18722155E-1	-.71433227E-1	-.20126367E-1	-.72259238E-1	-.19332913E-1
-.65662419E-1	-.16578150E-1	-.52451921E-1	-.12107055E-1	.39625288E+0	.378498
54.35GHz					
.10000000E+1	.32612371E+0	.48514269E+0	.10050889E+0	-.17703004E-1	-.10647098E-1
-.62398036E-1	-.19418578E-1	-.71607836E-1	-.20897355E-1	-.72441265E-1	-.20074386E-1
-.65328970E-1	-.17214275E-1	-.52586506E-1	-.12655095E-1	.40332036E+0	.382578
54.90GHz					
.10000000E+1	.33341119E+0	.48635588E+0	.10236397E+0	-.16685712E-1	-.10717014E-1
-.62422105E-1	-.19831740E-1	-.71712340E-1	-.21352967E-1	-.72549975E-1	-.20512442E-1
-.65928373E-1	-.17589948E-1	-.52666753E-1	-.12931337E-1	.41400230E+0	.384100
58.40GHz					
.10000000E+1	.37952617E+0	.49514276E+0	.11479302E+0	-.97308619E-2	-.10933021E-1
-.62549152E-1	-.22462198E-1	-.72440498E-1	-.24279620E-1	-.73319786E-1	-.23327733E-1
-.66632457E-1	-.20004832E-1	-.53234964E-1	-.14707537E-1	.44906141E+0	.392979
58.82GHz					
.10000000E+1	.38281677E+0	.49579889E+0	.11625385E+0	-.88608076E-2	-.10849801E-1
-.62538725E-1	-.22671747E-1	-.72518961E-1	-.24520767E-1	-.73394912E-1	-.23560197E-1
-.66701285E-1	-.20204411E-1	-.53290565E-1	-.14854617E-1	.45319758E+0	.394443
59.40GHz					
.10000000E+1	.39061003E+0	.49751566E+0	.11843982E+0	-.76143673E-2	-.10845311E-1
-.62554641E-1	-.23116588E-1	-.72657073E-1	-.25021294E-1	-.73540332E-1	-.24041975E-1
-.66834305E-1	-.20617730E-1	-.53397930E-1	-.15158714E-1	.45875429E+0	.395710

Table 8. Optical parameters (Marshall-Palmer 5mm/hr)

Freq(GHz)	$\bar{\omega}_1$	$\bar{\omega}_2$	$\bar{\omega}_3$	$\bar{\omega}_4$	$\bar{\omega}_5$
$\bar{\omega}_0$					
$\bar{\omega}_6$					
$\bar{\omega}_{12}$					
50.50GHz					
.1000000E+1	.29701037E+0	.48021003E+0	.95563178E-1	-.20293266E-1	-.10044780E-1
-.62249691E-1	-.17870954E-1	-.71252878E-1	-.19211714E-1	-.72075326E-1	-.18454867E-1
-.65494738E-1	-.15826063E-1	-.52317221E-1	-.11635461E-1	.45922777E+0	.3806662
53.20GHz					
.10000000E+1	.33364779E+0	.48594835E+0	.10490631E+0	-.15141100E-1	-.10397982E-1
-.62358333E-1	-.19949749E-1	-.71763088E-1	-.21503093E-1	-.72606076E-1	-.20657868E-1
-.65980102E-1	-.17715456E-1	-.52709038E-1	-.13024754E-1	.49359412E+0	.308807
54.35GHz					
.10000000E+1	.34602374E+0	.48803910E+0	.10896940E+0	-.12807568E-1	-.10356365E-1
-.62367490E-1	-.20686403E-1	-.71969109E-1	-.22328454E-1	-.72822629E-1	-.21452143E-1
-.66178294E-1	-.18396925E-1	-.52869161E-1	-.13526324E-1	.50785787E+0	.392676
54.90GHz					
.10000000E+1	.35362231E+0	.49848256E+0	.11101520E+0	-.11646018E-1	-.10382346E-1
-.62384982E-1	-.21119020E-1	-.72090750E-1	-.22811007E-1	-.72950132E-1	-.21916366E-1
-.66294909E-1	-.18795166E-1	-.52963325E-1	-.13819264E-1	.51456302E+0	.394098
58.40GHz					
.10000000E+1	.40164200E+0	.49986409E+0	.12472816E+0	-.37268698E-2	-.10271454E-1
-.62451675E-1	-.23865646E-1	-.72942269E-1	-.25910449E-1	-.73848994E-1	-.24900478E-1
-.67117426E-1	-.21355418E-1	-.53627074E-1	-.15702963E-1	.55574798E+0	.402370
58.82GHz					
.10000000E+1	.40514901E+0	.50068025E+0	.12632884E+0	-.27413713E-2	-.10144515E-1
-.62430076E-1	-.24087027E-1	-.73024837E-1	-.26171168E-1	-.73938141E-1	-.25152175E-1
-.67199188E-1	-.21571546E-1	-.53693186E-1	-.15862288E-1	.56068187E+0	.403769
59.40GHz					
.10000000E+1	.41324203E+0	.50268880E+0	.12874159E+0	-.13263080E-2	-.10076593E-1
-.62432216E-1	-.24549362E-1	-.73183344E-1	-.26700501E-1	-.74106899E-1	-.25662377E-1
-.67353649E-1	-.22009285E-1	-.53817836E-1	-.16184493E-1	.56707465E+0	.404943

Table 9. Optical parameters (Marshall-Palmer 10mm/hr)

Freq (Ghz)									
$\bar{\omega}_0$	$\bar{\omega}_1$	$\bar{\omega}_2$	$\bar{\omega}_3$	$\bar{\omega}_4$	$\bar{\omega}_5$	$\bar{\omega}_6$	$\bar{\omega}_7$	$\bar{\omega}_8$	$\bar{\omega}_9$
$\bar{\omega}_{12}$	$\bar{\omega}_{13}$	$\bar{\omega}_{14}$	$\bar{\omega}_{15}$	$\bar{\omega}_{10}$	$\bar{\omega}_{11}$	$\bar{\omega}_{12}$	$\bar{\omega}_{13}$	$\bar{\omega}_{14}$	$\bar{\omega}_{15}$
θ_{ext}	θ_{ext}	θ_{ext}	θ_{ext}	θ_{ext}	θ_{ext}	θ_{ext}	θ_{ext}	θ_{ext}	θ_{ext}
50.50GHz									
.1000000E+1	.36093848E+0	.48951757E+0	.12262584E+0	-.41402253E-2	-.87863093E-2	-.62023475E-1	-.21908425E-1	-.72427155E-1	-.73326414E-1
-.66641428E-1	-.19634845E-1	-.53245206E-1	-.14441571E-1	.89594238E+0	.411507	-.67435929E-1	-.21904511E-1	-.53086577E-1	.95218622E+0
54.35GHz									
.1000000E+1	.40258453E+0	.49938498E+0	.13527074E+0	.35405875E-2	-.82581204E-2	-.61953387E-1	-.24281849E-1	-.73234870E-1	-.74193661E-1
-.67755408E-1	-.22746010E-1	-.54152811E-1	-.16732326E-1	.97515409E+0	.421319	-.69250919E-1	-.26293969E-1	-.55352102E-1	.10510474E+1
54.90GHz									
.1000000E+1	.41701050E+0	.50317153E+0	.14071610E+0	.69763924E-2	-.77908956E-2	-.61818749E-1	-.25128965E-1	-.73562758E-1	-.74552777E-1
-.67755408E-1	-.22746010E-1	-.54152811E-1	-.16732326E-1	.97515409E+0	.421319	-.69250919E-1	-.26293969E-1	-.55352102E-1	.10510474E+1
58.40GHz									
.1000000E+1	.42556904E+0	.50556084E+0	.14349197E+0	.86890517E-2	-.75975237E-2	-.61818749E-1	-.25613949E-1	-.73749411E-1	-.74756449E-1
-.67952174E-1	-.23223385E-1	-.54303509E-1	-.17083881E-1	.98590729E+0	.422445	-.69250919E-1	-.26293969E-1	-.55352102E-1	.10510474E+1
58.92GHz									
.1000000E+1	.47941101E+0	.52235422E+0	.16208856E+0	.20282942E-1	-.59096163E-2	-.61453676E-1	-.28648382E-1	-.75027062E-1	-.76170676E-1
-.69250919E-1	-.26293969E-1	-.55352102E-1	-.19346159E-1	.10510474E+1	.429964	-.69385021E-1	-.26568938E-1	-.55460665E-1	.10585614E+1
59.40GHz									
.1000000E+1	.48359282E+0	.52382556E+0	.16421848E+0	.21703873E-1	-.55778436E-2	-.61273778E-1	-.28895100E-1	-.75153129E-1	-.76316263E-1
-.69385021E-1	-.26568938E-1	-.55460665E-1	-.19549150E-1	.10585614E+1	.430178	-.69385021E-1	-.26568938E-1	-.55460665E-1	.10585614E+1
59.40GHz									
.1000000E+1	.49258928E+0	.52697968E+0	.16749501E+0	.23761808E-1	-.52130210E-2	-.61273778E-1	-.29394957E-1	-.75384309E-1	-.76576783E-1
-.69624639E-1	-.27092870E-1	-.55654148E-1	-.19935314E-1	.10686644E+1	.431085	-.69624639E-1	-.27092870E-1	-.55654148E-1	.10686644E+1

Table 10. Optical parameters (Marshall-Palmer 15mm/hr)

Freq(GHz)	$\bar{\omega}_1$	$\bar{\omega}_2$	$\bar{\omega}_3$	$\bar{\omega}_4$	$\bar{\omega}_5$
$\bar{\omega}_0$					
$\bar{\omega}_6$					
$\bar{\omega}_{12}$					
50.50GHz					
.10000000E+1	.40489974E+0	.49945961E+0	.14234551E+0	.86064041E-2	-.67925486E-2
-.61550264E-1	-.24612721E-1	-.73452160E-1	-.27083996E-1	-.74456659E-1	-.26050576E-1
-.67679281E-1	-.22348698E-1	-.54084738E-1	-.16442662E-1	.13016306E+1	.427672
53.20GHz					
.10000000E+1	.44941841E+0	.51260638E+0	.15744676E+0	.18158680E-1	-.54152315E-2
-.61228033E-1	-.27113496E-1	-.74472934E-1	-.30122669E-1	-.75588129E-1	-.28993056E-1
-.60718405E-1	-.24075565E-1	-.54923391E-1	-.18304254E-1	.13745169E+1	.433375
54.35GHz					
.10000000E+1	.46502350E+0	.51773404E+0	.16391647E+0	.22405226E-1	-.45496123E-2
-.60992755E-1	-.2799760E-1	-.74885064E-1	-.31259461E-1	-.76058520E-1	-.30098638E-1
-.69151511E-1	-.25825809E-1	-.55273948E-1	-.19005093E-1	.14039892E+1	.436308
54.90GHz					
.10000000E+1	.47412169E+0	.52084973E+0	.16723198E+0	.24525361E-1	-.41537489E-2
-.60888758E-1	-.28501636E-1	-.75115838E-1	-.31895147E-1	-.76320671E-1	-.30716643E-1
-.69392801E-1	-.26356845E-1	-.55468810E-1	-.19396531E-1	.14177649E+1	.437277
58.40GHz					
.10000000E+1	.53119208E+0	.54246361E+0	.18944828E+0	.38836538E-1	-.10219920E-2
-.59998916E-1	-.31595389E-1	-.76670204E-1	-.35969580E-1	-.78124808E-1	-.34693194E-1
-.71057164E-1	-.29775847E-1	-.56813578E-1	-.21918156E-1	.15005644E+1	.442872
58.82GHz					
.10000000E+1	.53575518E+0	.54443285E+0	.19196726E+0	.40574308E-1	-.50714008E-3
-.59832359E-1	-.31841726E-1	-.76822736E-1	-.36340519E-1	-.78312949E-1	-.35059194E-1
-.71231725E-1	-.30091143E-1	-.56954878E-1	-.22151202E-1	.15100086E+1	.443992
59.40GHz					
.10000000E+1	.54524228E+0	.54842635E+0	.19588108E+0	.43107757E-1	.12957861E-3
-.59637668E-1	-.32340639E-1	-.77098336E-1	-.37031008E-1	-.78641636E-1	-.35736689E-1
-.71535870E-1	-.30674074E-1	-.57200889E-1	-.22581343E-1	.15227365E+1	.444759

Table 11. Optical parameters (Marshall-Palmer 20mm/hr)

Freq(GHz)	$\bar{\omega}_0$	$\bar{\omega}_1$	$\bar{\omega}_2$	$\bar{\omega}_3$	$\bar{\omega}_4$	$\bar{\omega}_5$
	$\bar{\omega}_6$	$\bar{\omega}_7$	$\bar{\omega}_8$	$\bar{\omega}_9$	$\bar{\omega}_{10}$	$\bar{\omega}_{11}$
	$\bar{\omega}_{12}$	$\bar{\omega}_{13}$	$\bar{\omega}_{14}$	$\bar{\omega}_{15}$	β_{ext}	$\bar{\omega}_v$
50.50GHz	.43903765E+0	.50919796E+0	.15845329E+0	.19427627E-1	-.45405659E-2	
.10000000E+1	-.26642226E-1	-.74357782E-1	-.29684166E-1	-.75489565E-1	-.28577161E-1	
-.60929717E-1	-.24520497E-1	-.54854059E-1	-.18045204E-1	.16838562E+1	.438371	
53.20GHz	.48552493E+0	.52507979E+0	.17558237E+0	.30517876E-1	-.23651432E-2	
.10000000E+1	-.29190570E-1	-.75539413E-1	-.32946217E-1	-.76842013E-1	-.31751468E-1	
-.60324405E-1	-.27240487E-1	-.55860447E-1	-.20056421E-1	.17703700E+1	.443509	
54.35GHz	.50193737E+0	.53132037E+0	.18289610E+0	.35426753E-1	-.11286337E-2	
.10000000E+1	-.30083094E-1	-.76012826E-1	-.34175223E-1	-.77404730E-1	-.32955103E-1	
-.59943299E-1	-.28284035E-1	-.56281167E-1	-.20820792E-1	.18051030E+1	.446229	
54.90GHz	.51139987E+0	.53503014E+0	.18665708E+0	.37881713E-1	-.54439316E-3	
.10000000E+1	-.30585711E-1	-.76275507E-1	-.34856000E-1	-.77715034E-1	-.33621469E-1	
-.59765431E-1	-.28857205E-1	-.56512977E-1	-.21243618E-1	.18213244E+1	.447094	
58.40GHz	.57062622E+0	.56053480E+0	.21184221E+0	.54429277E-1	.39097171E-2	
.10000000E+1	-.33632881E-1	-.78019590E-1	-.39212432E-1	-.79836925E-1	-.37910917E-1	
-.59321935E-1	-.32550020E-1	-.58104159E-1	-.23969665E-1	.19182941E+1	.452083	
58.92GHz	.57544649E+0	.56290685E+0	.21467965E+0	.56426206E-1	.45898001E-2	
.10000000E+1	-.33866531E-1	-.78188716E-1	-.39614499E-1	-.80059641E-1	-.38313001E-1	
-.58077910E-1	-.32897094E-1	-.58272864E-1	-.24226486E-1	.19292605E+1	.453141	
59.40GHz	.58525632E+0	.56756765E+0	.21911441E+0	.59351559E-1	.54706432E-2	
.10000000E+1	-.34347081E-1	-.78492791E-1	-.40350446E-1	-.80443102E-1	-.39043284E-1	
-.57774342E-1	-.33526580E-1	-.58561881E-1	-.24691468E-1	.19440731E+1	.453817	
-.73216236E-1						

Table 12. Optical parameters (Marshall-Palmer 25mm/hr)

Freq(GHz)	$\bar{\omega}_0$	$\bar{\omega}_1$	$\bar{\omega}_2$	$\bar{\omega}_3$	$\bar{\omega}_4$	$\bar{\omega}_5$
	$\bar{\omega}_6$	$\bar{\omega}_7$	$\bar{\omega}_8$	$\bar{\omega}_9$	$\bar{\omega}_{10}$	$\bar{\omega}_{11}$
	$\bar{\omega}_{12}$	$\bar{\omega}_{13}$	$\bar{\omega}_{14}$	$\bar{\omega}_{15}$	β_{ext}	$\bar{\omega}_v$
50.50GHz						
.1000000E+1	.46717341E+0	.51855401E+0	.17232806E+0	.28967801E-1	-.21834488E-2	
-.60213211E-1	-.28249534E-1	-.75167057E-1	-.31870951E-1	-.76444522E-1	-.30713691E-1	
-.69512096E-1	-.26358383E-1	-.55567630E-1	-.19402249E-1	.20475379E+1	.446257	
53.20GHz						
.10000000E+1	.51512995E+0	.53678951E+0	.19121456E+0	.41386600E-1	.74454173E-3	
-.59310646E-1	-.30798992E-1	-.76473803E-1	-.35312376E-1	-.77987951E-1	-.34080899E-1	
-.70938488E-1	-.29254492E-1	-.56720548E-1	-.21538707E-1	.21456670E+1	.450980	
54.35GHz						
.10000000E+1	.53214434E+0	.54398351E+0	.19925572E+0	.46867901E-1	.23248313E-2	
-.58779234E-1	-.31677175E-1	-.76992487E-1	-.36614429E-1	-.78630161E-1	-.35365866E-1	
-.71534858E-1	-.30361208E-1	-.57203110E-1	-.22356254E-1	.21848400E+1	.453545	
54.90GHz						
.10000000E+1	.54187524E+0	.54819744E+0	.20340130E+0	.49612777E-1	.30848575E-2	
-.58525542E-1	-.32170743E-1	-.77278504E-1	-.37330867E-1	-.78981650E-1	-.36072509E-1	
-.71861179E-1	-.30969705E-1	-.57467045E-1	-.22805463E-1	.22031279E+1	.454335	
58.40GHz						
.10000000E+1	.60267618E+0	.57697304E+0	.23114114E+0	.68094563E-1	.87656564E-2	
-.56522240E-1	-.35111644E-1	-.79154474E-1	-.41907684E-1	-.81372392E-1	-.40622032E-1	
-.74090999E-1	-.34892315E-1	-.59271781E-1	-.25703557E-1	.23119910E+1	.458881	
58.82GHz						
.10000000E+1	.60768566E+0	.57968392E+0	.23425180E+0	.70314457E-1	.95974271E-2	
-.56201517E-1	-.35325828E-1	-.79333503E-1	-.42333479E-1	-.81624045E-1	-.41053989E-1	
-.74328186E-1	-.35266001E-1	-.59464233E-1	-.25980347E-1	.23242168E+1	.459895	
59.40GHz						
.10000000E+1	.61772795E+0	.58489953E+0	.23913298E+0	.7357952E-1	.10703176E-1	
-.55789220E-1	-.35778189E-1	-.79655550E-1	-.43104248E-1	-.82053374E-1	-.41828149E-1	
-.74730993E-1	-.35934602E-1	-.59790541E-1	-.26474721E-1	.23407661E+1	.460503	

Table 13. Optical parameters (Marshall-Palmer 30mm/hr)

Freq(GHz)	$\bar{\omega}_0$	$\bar{\omega}_1$	$\bar{\omega}_2$	$\bar{\omega}_3$	$\bar{\omega}_4$	$\bar{\omega}_5$
	$\bar{\omega}_6$	$\bar{\omega}_7$	$\bar{\omega}_8$	$\bar{\omega}_9$	$\bar{\omega}_{10}$	$\bar{\omega}_{11}$
	$\bar{\omega}_{12}$	$\bar{\omega}_{13}$	$\bar{\omega}_{14}$	$\bar{\omega}_{15}$	β_{ext}	$\bar{\omega}_v$
50.50GHz						
.10000000E+1	.49121131E+0	.52749945E+0	.18465455E+0	.37577904E-1	.21290313E-3	
-.59429845E-1	-.29563076E-1	-.75897299E-1	-.33771160E-1	-.77335382E-1	-.32581278E-1	
-.70337569E-1	-.27966255E-1	-.56235893E-1	-.20590112E-1	.23960058E+1	.452447	
53.20GHz						
.10000000E+1	.54032236E+0	.54780426E+0	.20510323E+0	.51180242E-1	.38509139E-2	
-.58223647E-1	-.32082087E-1	-.77302182E-1	-.37361464E-1	-.79047595E-1	-.36114894E-1	
-.71925554E-1	-.31008134E-1	-.57320069E-1	-.22835439E-1	.25042903E+1	.456847	
54.35GHz						
.10000000E+1	.55780879E+0	.55583315E+0	.21379108E+0	.57168584E-1	.57554127E-2	
-.57541798E-1	-.32934138E-1	-.77854825E-1	-.38723578E-1	-.79759449E-1	-.37470051E-1	
-.72589742E-1	-.32176766E-1	-.58057924E-1	-.23699273E-1	.25473123E+1	.459291	
54.90GHz						
.10000000E+1	.56774825E+0	.56048550E+0	.21827777E+0	.60172031E-1	.66808367E-2	
-.57212071E-1	-.33412287E-1	-.78157829E-1	-.39469207E-1	-.80146923E-1	-.38211398E-1	
-.72951209E-1	-.32815977E-1	-.50350499E-1	-.24171486E-1	.25673943E+1	.460022	
58.40GHz						
.10000000E+1	.62976264E+0	.59208608E+0	.24827964E+0	.80379454E-1	.13509334E-1	
-.54649687E-1	-.36207737E-1	-.80123911E-1	-.44223746E-1	-.82770825E-1	-.42905301E-1	
-.75413257E-1	-.36938696E-1	-.60345256E-1	-.27219927E-1	.26865309E+1	.464224	
58.82GHz						
.10000000E+1	.63491862E+0	.59508925E+0	.25163174E+0	.82797327E-1	.14482012E-1	
-.54233648E-1	-.36398237E-1	-.80308177E-1	-.44668250E-1	-.83047480E-1	-.43442921E-1	
-.75676103E-1	-.37335561E-1	-.60558841E-1	-.27514176E-1	.26998304E+1	.465204	
59.40GHz						
.10000000E+1	.64513795E+0	.60078004E+0	.25590628E+0	.86364644E-1	.15797053E-1	
-.53734037E-1	-.36816316E-1	-.80640734E-1	-.45466419E-1	-.83516071E-1	-.44254764E-1	
-.76119073E-1	-.38038221E-1	-.60918238E-1	-.28034263E-1	.27178708E+1	.465761	

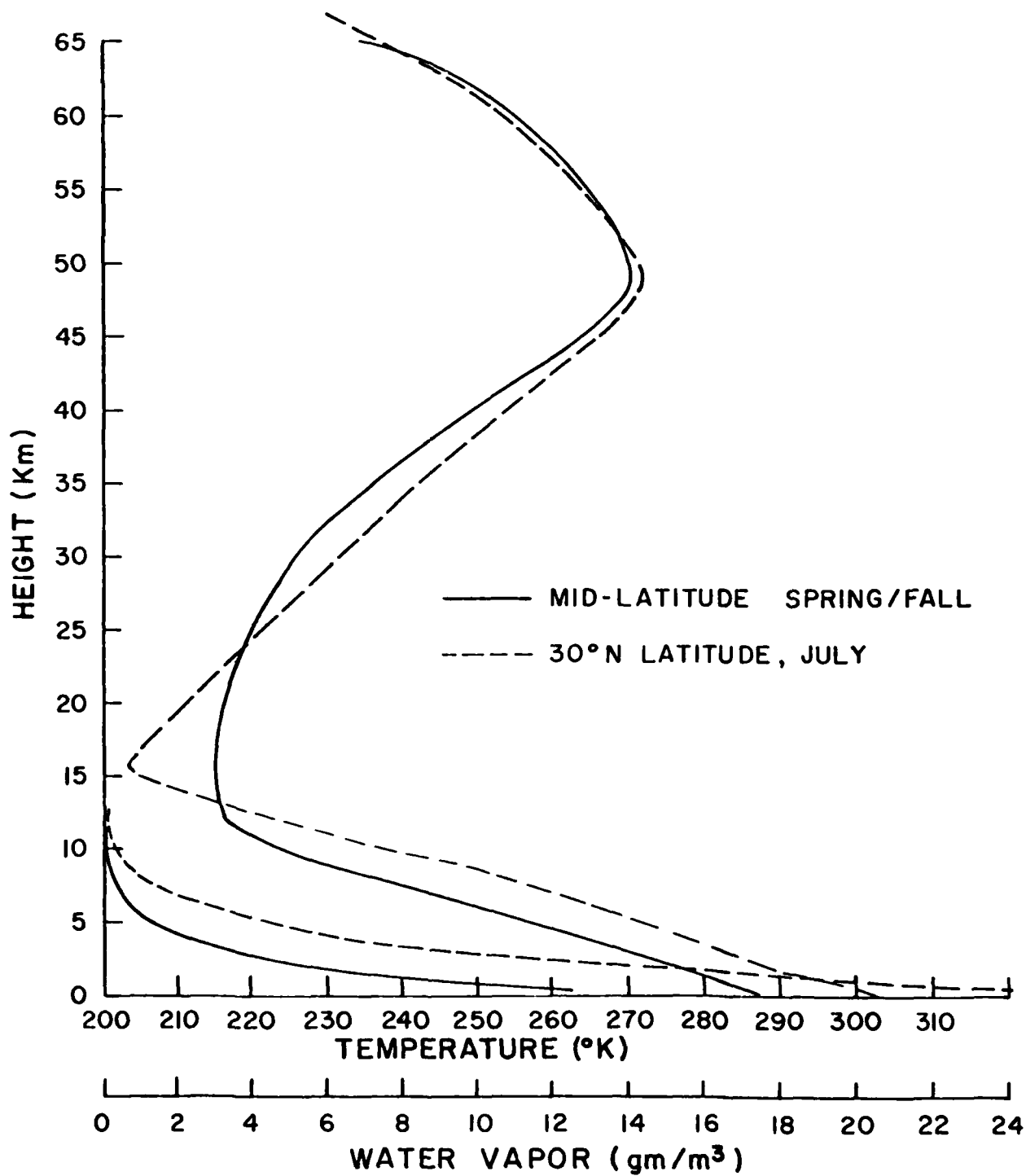


Figure 5. Climatological temperature and water vapor profiles for the northern hemispheric mid-latitude Spring/Fall and 30° north latitude, July.

$$V = \frac{4}{3} \pi \int_0^{\infty} r^3 N(r) dr . \quad (5.1)$$

Therefore the total liquid water content per unit volume of atmosphere is

$$W = \frac{4}{3} \pi \rho \int_0^{\infty} r^3 N(r) dr \text{ (gm cm}^{-3}\text{)} , \quad (5.2)$$

where r is in units of cm, $N(r)$ is in terms of cm^{-4} , and the density of liquid water ρ is 1 gm cm^{-3} . For the Deirmendjian L-Model we have

$$W = 1.16678(10^{-2}) \text{ (gm cm}^{-2} \text{ km}^{-1}\text{)} , \quad (5.3)$$

while for the Marshall-Palmer size distribution we find

$$W = 1.77883(10^{-2}) R^{.84} \text{ (gm cm}^{-2} \text{ km}^{-1}\text{)} . \quad (5.4)$$

Table 14 contains the total mass of liquid water within a column of 1 cm^2 cross section for the cloud models used.

Table 14. Total liquid water content ($10^{-2} \text{ gm cm}^{-2}$).

Thickness (km)	1	2	3	4	5
Deirmendjian	1.16678	2.33356	3.50034	4.66712	5.83390
Marshall-Palmer R(mm/hr)					
1	0.88941	1.77882	2.66823	3.55764	4.44705
2	1.59209	3.18418	4.77627	6.36835	7.96044
3	2.23812	4.47625	6.71437	8.95249	11.1906
4	2.84992	5.69984	8.54975	11.3997	14.2496
5	3.43745	6.87491	10.3124	13.7498	17.1873
10	6.15321	12.3064	18.4596	24.6129	30.7661
15	8.65005	17.3001	25.9502	34.0020	43.2503
20	11.0146	22.0291	33.0437	44.0583	55.0728
25	13.2853	26.5706	39.8559	53.1412	66.4266
30	15.4840	30.9681	46.5421	61.9361	77.4201

SECTION 4

SENSITIVITY ANALYSES ON THE EFFECTS OF CLOUD AND PRECIPITATION ON THE DMSP SSM/T CHANNELS

In this section, we present the results computed from the microwave radiative transfer program for cloudy atmospheres described in Section 2. The computations utilize the single-scattering parameters presented in Section 3 for input to the transfer program. In view of the weighting function diagram depicted in Figure 2, it is anticipated that only 50.50, 53.20 and 54.35 GHz (channels 1-3) whose peaks of the weighting function are below the tropopause will be affected by clouds and precipitation. However, in the following analyses, 54.90 GHz (channel 4) is also included in the computations. Effects of the rainfall rate and cloud layer thickness on the upwelling ($\mu=1$) brightness temperature over land and ocean surfaces are first discussed. Computational results due to variations of the cloud location and atmospheric profile are then reported. Physical interpretation of the results are given in the last Subsection.

4.1 Dependence on Layer Thickness and Rainfall Rate

To determine the effect of varying cloud thicknesses on the brightness temperatures reaching the top of the atmosphere, cloud models were inserted into the atmosphere with a constant cloud base of 1 km.

Thickness cases of 1, 2, 3, 4 and 5 km were examined. Variations of the brightness temperature over land with respect to cloud thickness and rainfall rate for channels 1 through 4 are displayed in Figures 6 through 10. The surface emissivity over land is taken to be 0.97 for all channels in this sensitivity study. The energy sensed by channels 5, 6 and 7 originates sufficiently high in the atmosphere that the brightness temperatures are insignificantly affected by the cloud models as pointed out previously.

The differences due to thickness for channels 3 and 4 may be misleading. It is more likely that a positional dependence for the cloud is being exhibited. Since the cloud base was fixed at 1 km, only the thicker clouds reached into the energy peak source regions for channels 3 and 4 (weighting functions peak at about 7 and 10.5 km, respectively). For example, a 5 km thick cloud with a 1 km base at a 2 mm/hr rainfall rate results in brightness temperatures for channels 3 and 4 of 237.3° K and 228.9° K. The Deirmendjian L-Model cloud used in the position study has liquid water content approximately equal to that of the 2 mm/hr rain model and yet a 2 km thick L-Model cloud with base at 4 km results in nearly the same brightness temperatures for channels 3 and 4 as the 5 km thick cloud discussed above (237.2° K and 228.7° K, respectively).

Channels 1 and 2 (peaking near the surface and 2 km) show decreasing brightness temperatures for increasing cloud thicknesses. The decrease is greater for clouds modeling higher rainfall rates. For channel 1 at a rainfall rate of 1 mm/hr the decrease ranges from 2° K for a 1 km thick cloud to 16° K for a 5 km thick cloud. For a

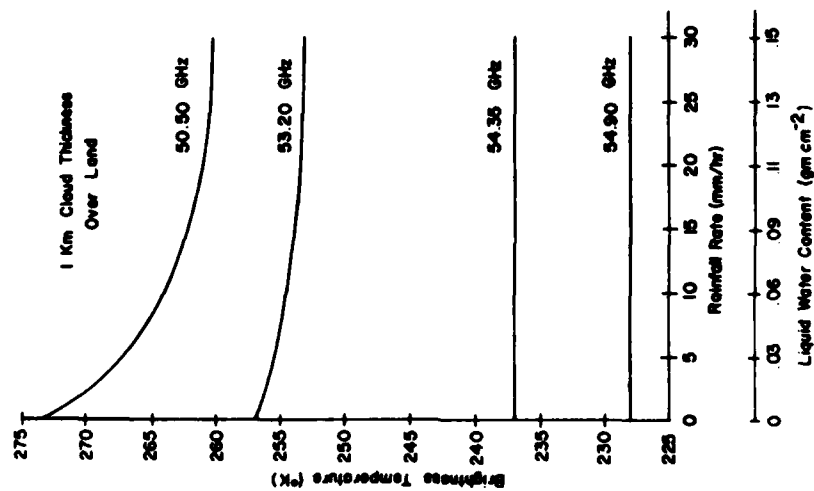


Figure 6. Brightness temperature versus rainfall rate (or liquid water content) for 1 km thick cloud over land (mid-latitude Spring/Fall profile).

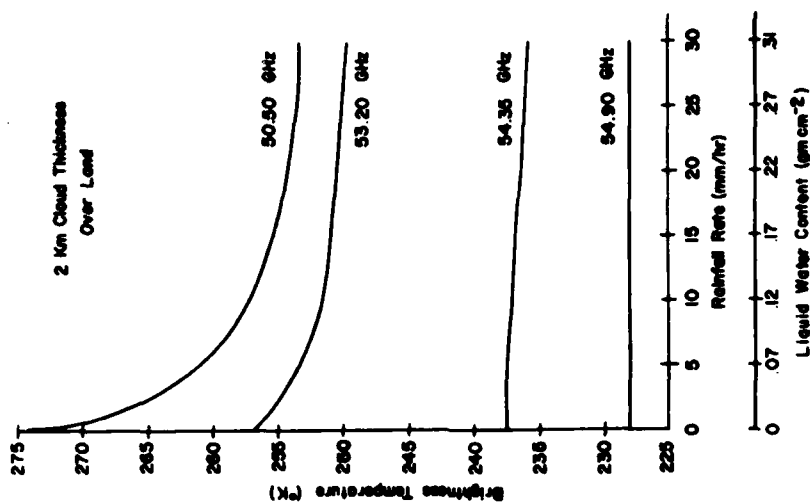


Figure 7. Brightness temperature versus rainfall rate (or liquid water content) for 2 km thick cloud over land (mid-latitude Spring/Fall profile).

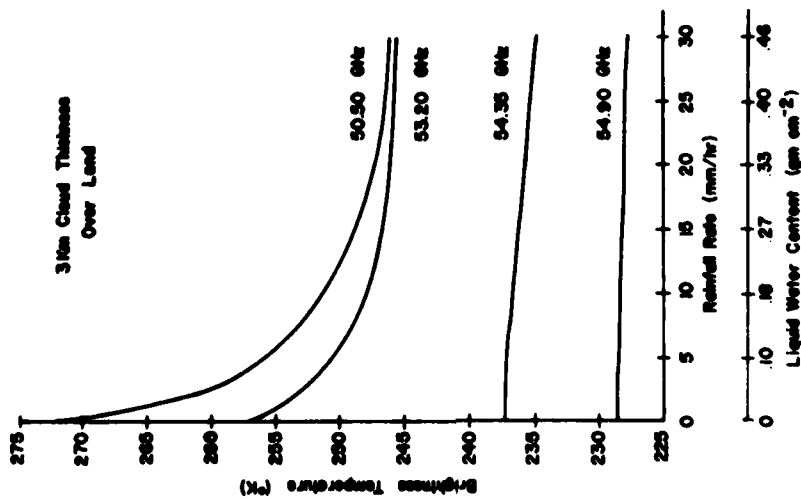


Figure 8. Brightness temperature versus rainfall rate (or liquid water content) for 3 km thick cloud over land (mid-latitude Spring/Fall profile).

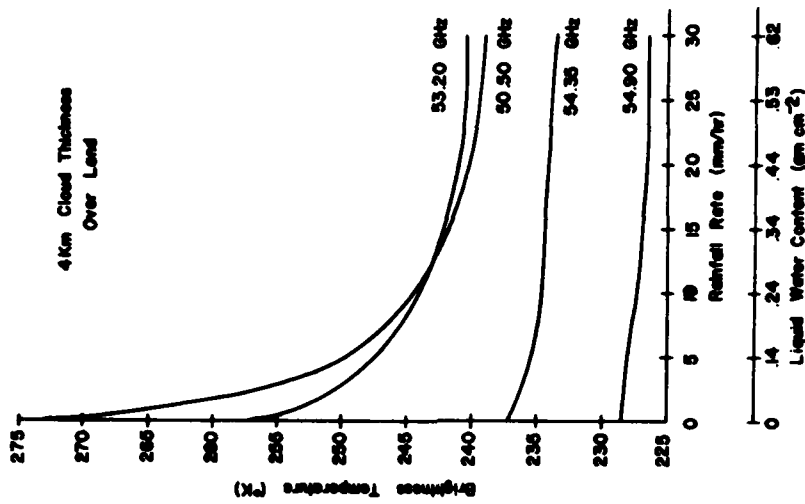


Figure 9. Brightness temperature versus rainfall rate (or liquid water content) for 4 km thick cloud over land (mid-latitude Spring/Fall profile).

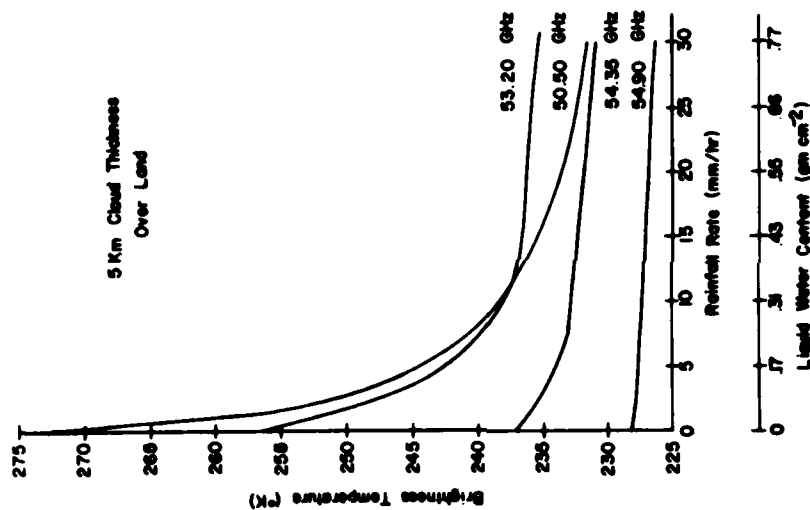


Figure 10. Brightness temperature versus rainfall rate (or liquid water content) for 5 km thick cloud over land (mid-latitude Spring/Fall profile).

rainfall rate of 30 mm/hr the channel 1 brightness temperature decreases 14° K for a 1 km thick cloud and 43° K for a 5 km thick cloud. Channel 2 displays the same trends as channel 1 but to a lesser degree. For a rainfall rate of 1 mm/hr the decrease in channel 2 brightness temperature as a function of cloud thickness ranges from less than 0.5° K to 5° K. For a rainfall rate of 30 mm/hr the range is 4° K to 22° K.

The results for channels 3 and 4 show relatively constant brightness temperatures unaffected by rainfall rates less than about 3 mm/hr. For rainfall rates greater than 3 mm/hr the brightness temperature decreases with increasing cloud thickness. The maximum decreases are 7° K for channel 3 and 2° K for channel 4; both at rainfall rates of 30 mm/hr and cloud thicknesses of 5 km.

Figures 11 through 14 depict the variation of brightness temperatures over ocean having an emissivity of 0.51 with respect to cloud thickness and liquid water content for channels 1-4. The results over the ocean are quite different from the results over land. Note that the only physical difference between these two cases is that the surface emissivity for all channels is taken to be 0.97 for a land surface and 0.51 for an ocean surface. The difference in the surface emissivity reveals that the surface contribution over ocean is approximately one half of the surface contribution over land, and that the atmospheric contribution reflected from the earth's surface is approximately 16 times greater over ocean.

Figures 11 and 12 show that channels 1 and 2 follow similar trends over the ocean but to different degrees. For all cloud thicknesses, an initial reduction in the brightness temperature is observed for clouds with low liquid water content. For channel 1 the maximum

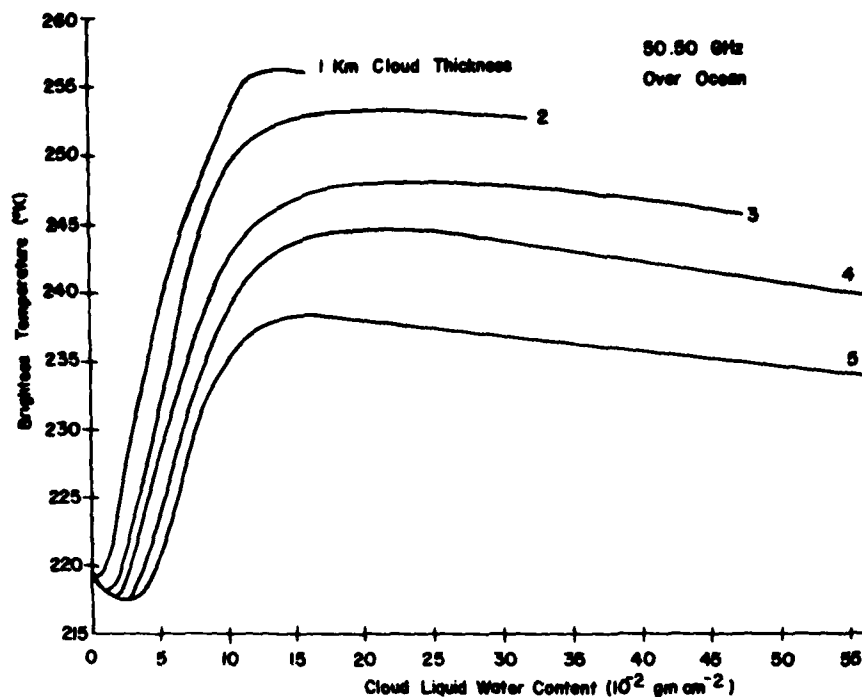


Figure 11. 50.50 GHz brightness temperature as a function of cloud thickness and liquid water content (mid-latitude Spring/Fall profile over ocean).

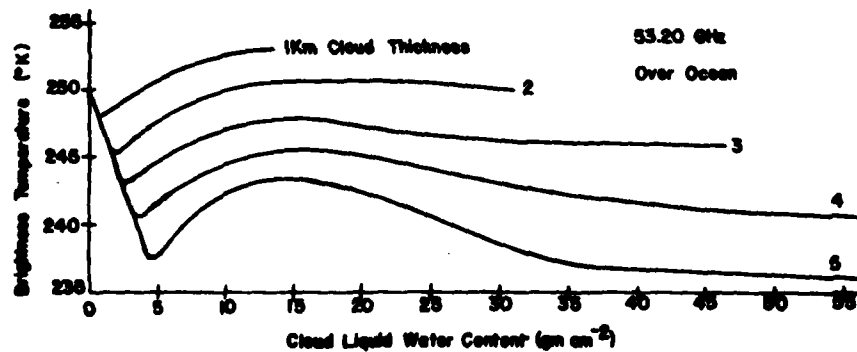


Figure 12. 53.20 GHz brightness temperature as a function of cloud thickness and liquid water content (mid-latitude Spring/Fall profile over ocean). Values of the cloud liquid water content in this figure should be multiplied by 10^{-2} .

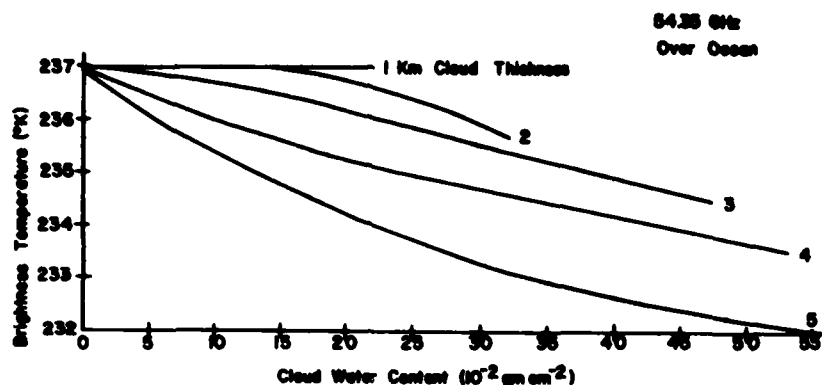


Figure 13. 54.35 GHz brightness temperature as a function of cloud thickness and liquid water content (mid-latitude Spring/Fall profile over ocean).

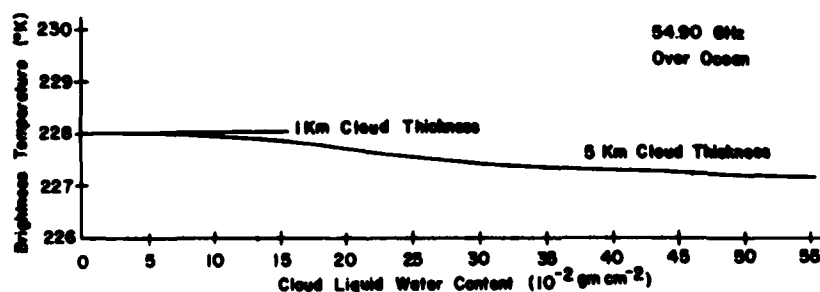


Figure 14. 54.90 GHz brightness temperature as a function of cloud thickness and liquid water content (mid-latitude Spring/Fall profile over ocean).

reduction is 2° K, while for channel 2 the maximum reduction is nearly 13° K. At some liquid water content less than $5 \times 10^{-2} \text{ gm cm}^{-2}$ the brightness temperature begins to increase. The turning point occurs at lower liquid water content for thinner clouds. The brightness temperature then increases until the liquid water content reaches about $15 \times 10^{-2} \text{ gm cm}^{-2}$. For channel 1 this increase results in a maximum brightness temperature more than 35° K higher than the clear column value. For channel 2 the maximum brightness temperature is about 3° K greater than the clear column value. For liquid water contents greater than $15 \times 10^{-2} \text{ gm cm}^{-2}$, the brightness temperature decreases for both channels and all cloud thicknesses. Note that the maximum brightness temperature observed over ocean is very nearly the same as the minimum brightness temperature observed over land (within about $4 - 7^{\circ}$ K).

Figures 13 and 14 indicate that channels 3 and 4 are relatively unaffected by a 1 km thick cloud with a cloud base at 1 km. Increasing the cloud thickness above 1 km causes steadily decreasing brightness temperatures. For a constant cloud thickness, increasing the liquid water content causes a decrease in the brightness temperature. Note that channel 4 is affected by less than 1° K for all of the cases studied.

4.2 Dependence on Layer Location and Atmospheric Profile

In this subsection, we first investigate the effect of the position of the cloud layer in the atmosphere on the upwelling brightness temperature. Figures 15 through 20 display the importance of cloud position relative to the peak of the weighting function for each channel over the land surface.

— Brightness Temperature
 --- Weighting Function

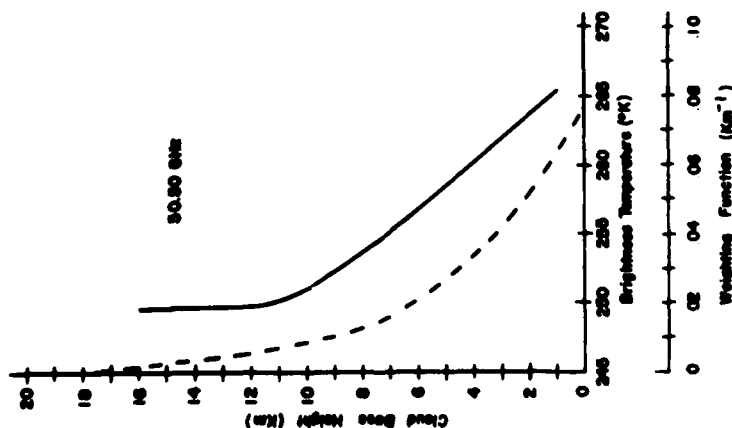


Figure 15. 50.50 GHz brightness temperature dependence on cloud base height for a 2 km thick Deirmendjian L-Model cloud (mid-latitude Spring/Fall profile over land).

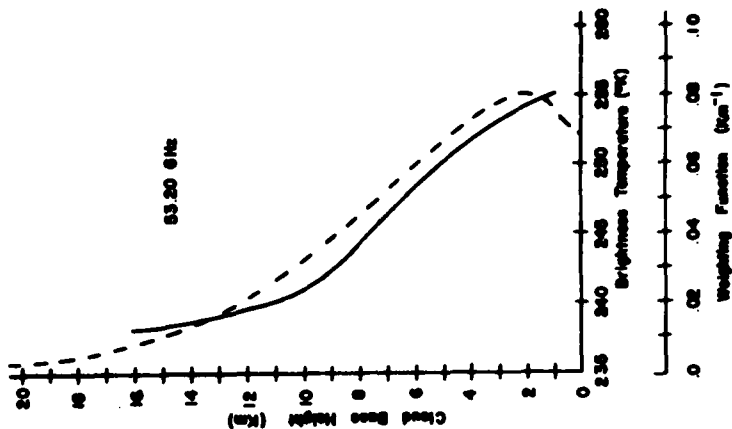


Figure 16. 53.20 GHz brightness temperature dependence on cloud base height for a 2 km thick Deirmendjian L-Model cloud (mid-latitude Spring/Fall profile over land).

— Brightness Temperature
 --- Weighting Function

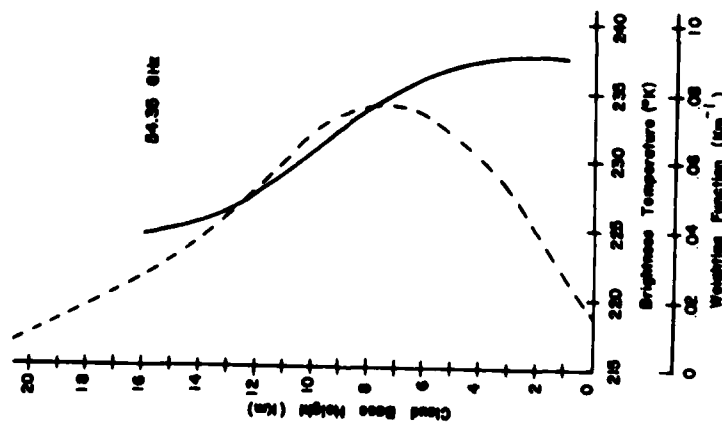


Figure 17. 54.35 GHz brightness temperature dependence on cloud base height for a 2 km thick Deirmendjian L-Model cloud (mid-latitude Spring/Fall profile over land).

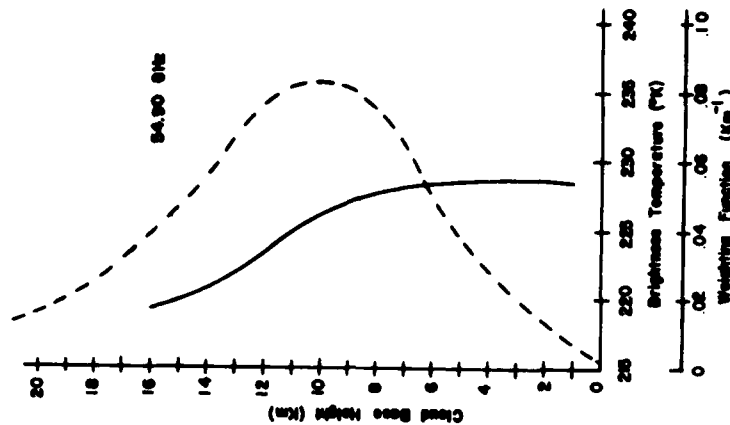


Figure 18. 54.90 GHz brightness temperature dependence on cloud base height for a 2 km thick Deirmendjian L-Model cloud (mid-latitude Spring/Fall profile over land).

— Brightness Temperature
 --- Weighting Function

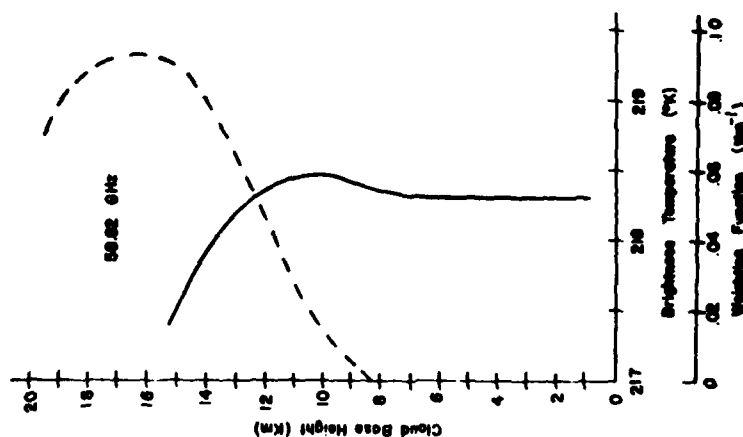


Figure 19. 58.82 GHz brightness temperature dependence on cloud base height for a 2 km thick Deirmendjian L-Model cloud (mid-latitude Spring/Fall profile over land).

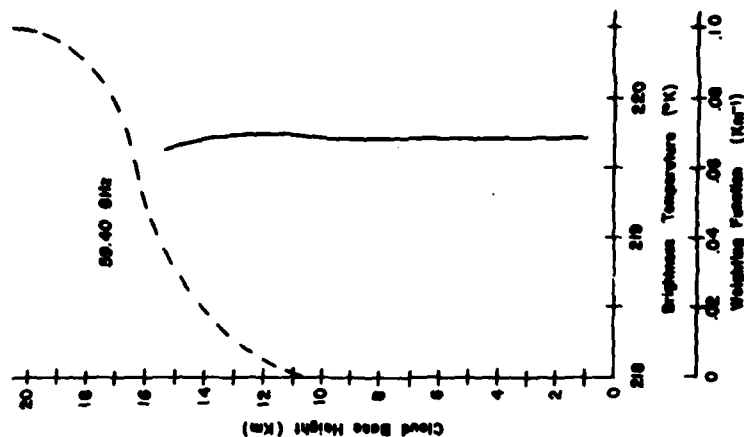


Figure 20. 59.40 GHz brightness temperature dependence on cloud base height for a 2 km thick Deirmendjian L-Model cloud (mid-latitude Spring/Fall profile over land).

Channels 1 and 2 show the effects on a channel when the cloud is at or above the peak of the weighting function, i.e., the prime energy source region for the channel. The brightness temperature decreases rapidly as the cloud is moved higher putting more and more of the energy source for the channel beneath it. Eventually, the brightness temperature becomes near constant as all of the significant source region for channel energy is below the cloud and therefore raising the cloud higher has only slight effects.

Channels 3 and 4 show the results of moving a cloud from below the prime energy source region up through the source region. Far below the energy source the brightness temperature remains nearly constant. Then, as the cloud moves into the source region, the brightness temperature decreases significantly, becoming near constant above the source region as for channels 1 and 2.

Channels 6 and 7 indicate that for channels peaking high enough in the atmosphere to be free of surface effects, brightness temperatures increase slightly as the cloud approaches the energy source region from below.

Sensitivity analyses on the positional dependence were also carried out over the ocean surface. It was found that the same positional dependence exists over the ocean as over the land with no change in trends for any of the channels.

Figures 21 through 28 display the results when the sensitivity analysis is duplicated for the 30° N latitude, July, profile. Although the different profile naturally results in different brightness temperatures, the trends noted for the northern hemispheric mid-latitude Spring/Fall profile for liquid water content greater than about $3 \times 10^{-2} \text{ gm cm}^{-2}$ are observed to persist.

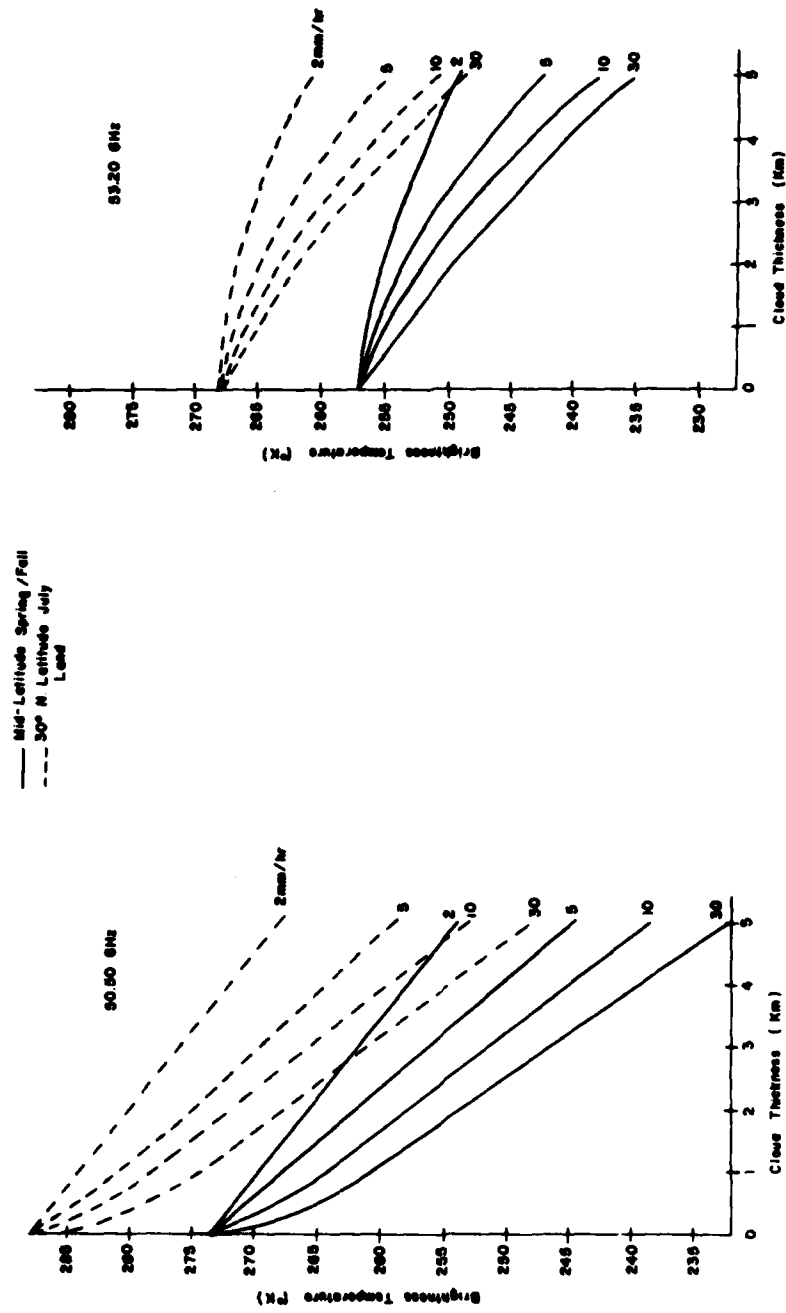


Figure 21. 50.50 GHz channel profile dependence over land.

Figure 22. 53.20 GHz channel profile dependence over land.

— Mid-Latitude Spring/Fall
 --- 30° North Latitude July
 Land

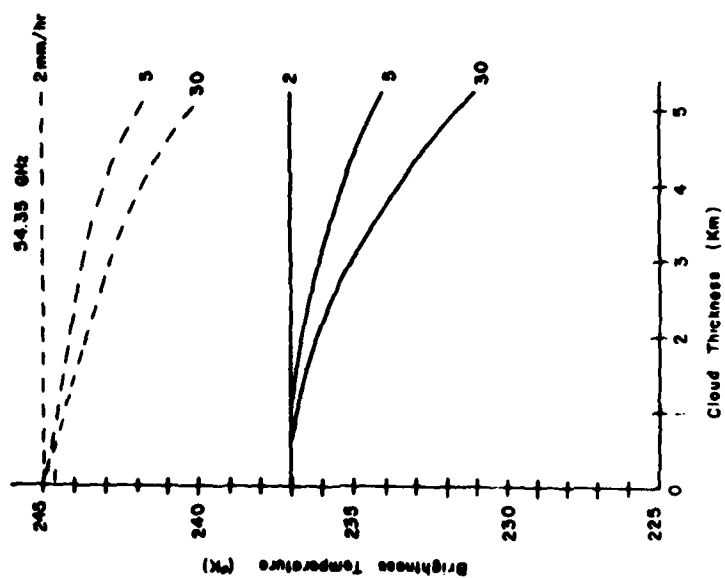


Figure 23. 54.35 GHz channel profile dependence over land.

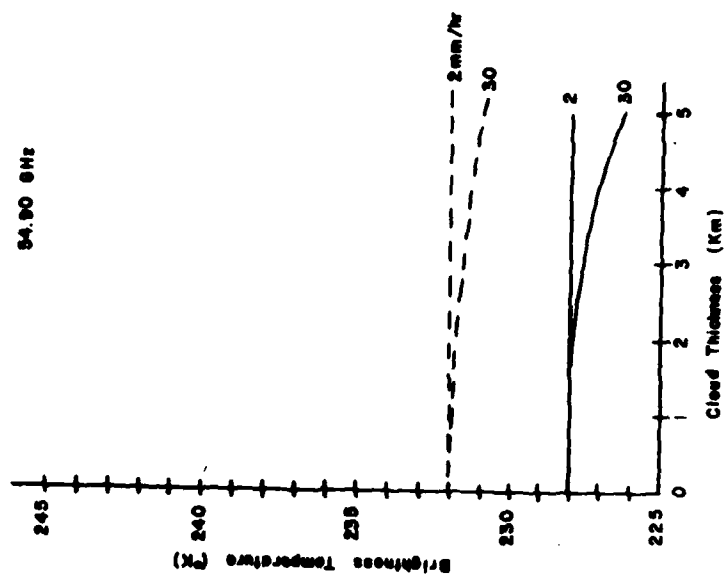


Figure 24. 54.90 GHz channel profile dependence over land.

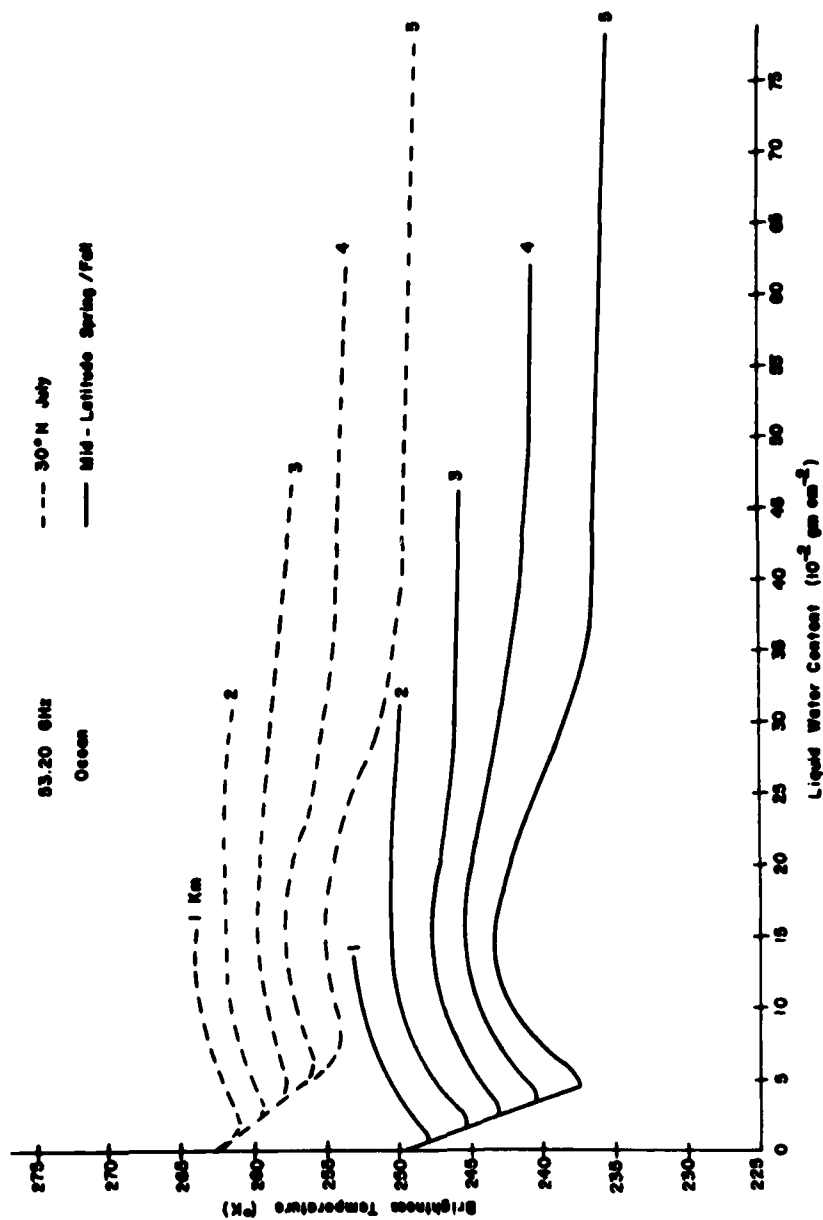


Figure 25. 53.20 GHz channel profile dependence over ocean.

50.50 GHz
Ocean

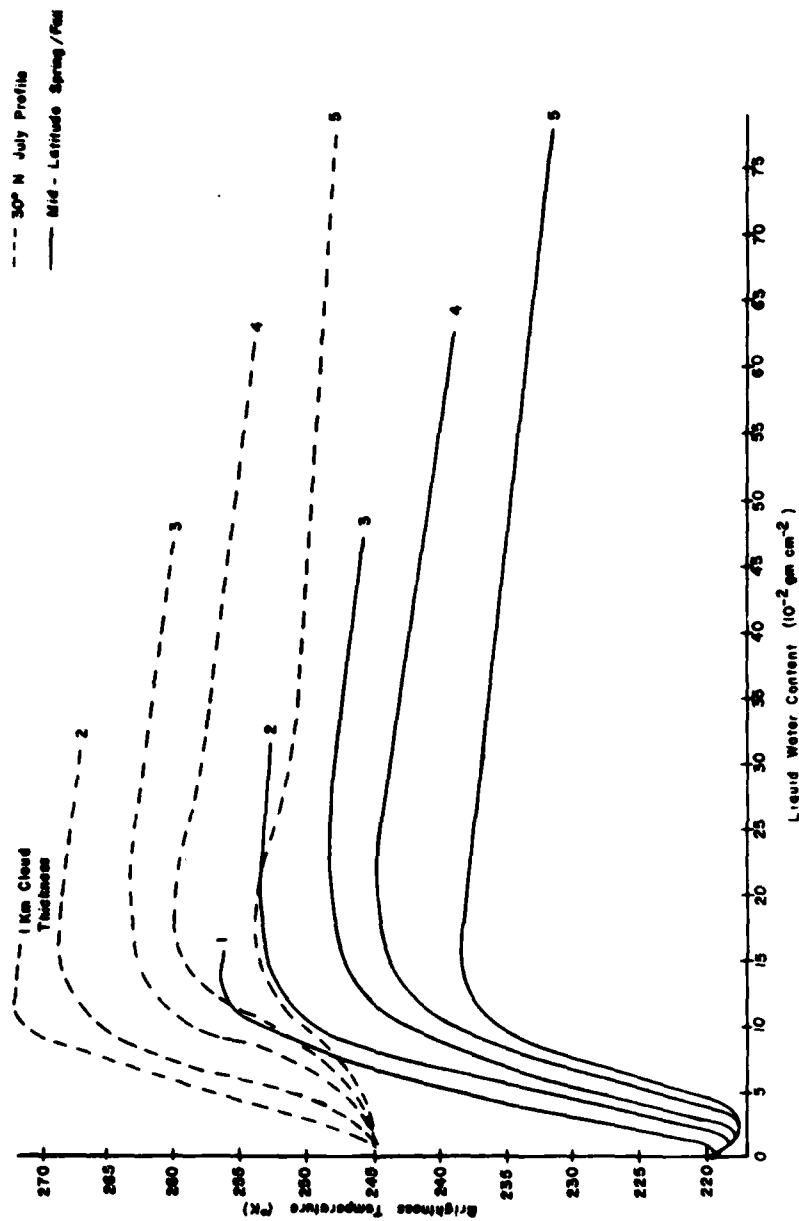


Figure 26. 50.50 GHz channel profile dependence over ocean.

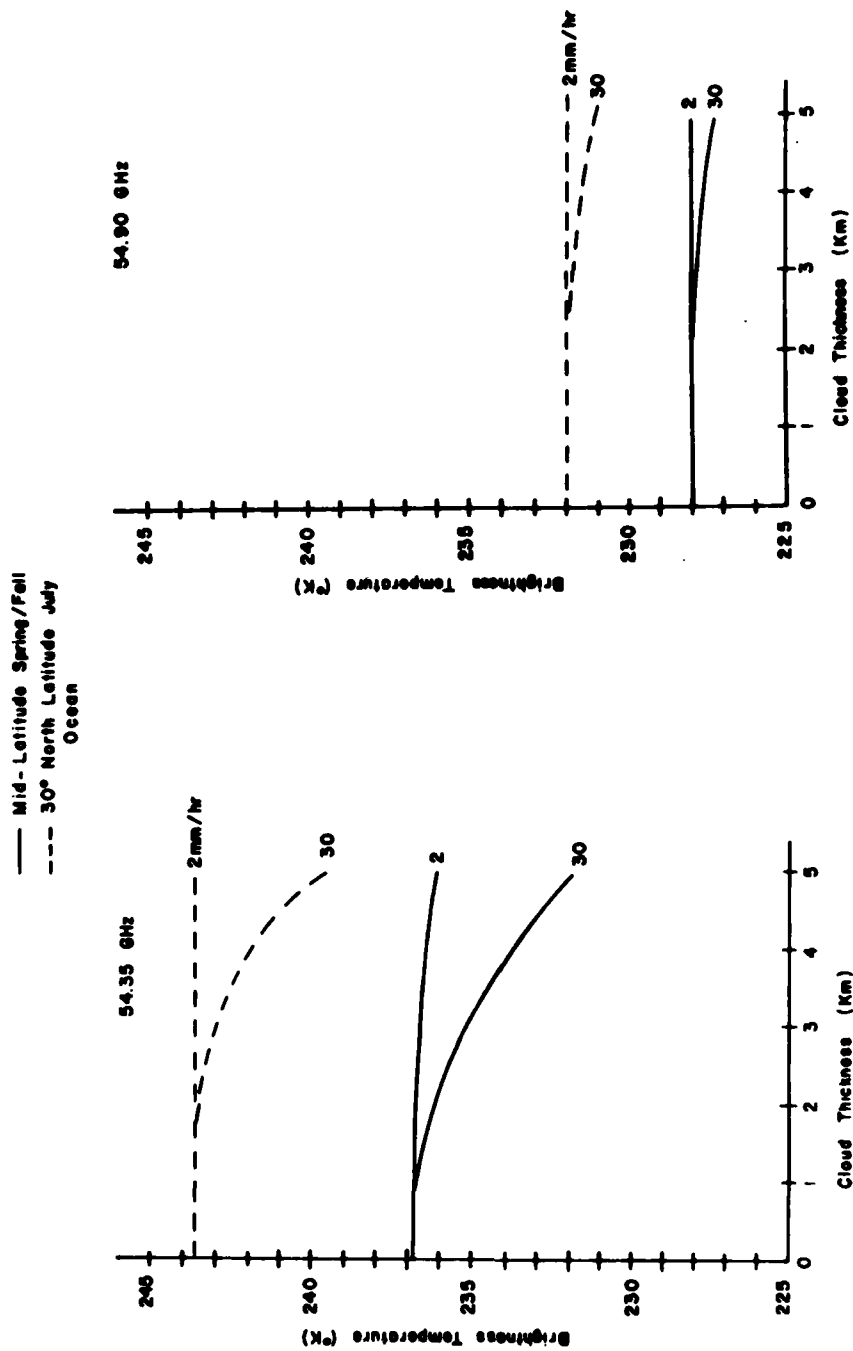


Figure 27. 54.35 GHz channel profile dependence over ocean.

Figure 28. 54.90 GHz channel profile dependence over ocean.

4.3 Interpretation of Results

In order to aid in the interpretation of results Figure 29 is included which graphically depicts the terms contributing to the cloud boundary conditions and the mechanisms for energy loss and gain within the cloud. There are three mechanisms for energy gain within the cloud caused by gaseous emission, droplet emission, and multiple scattering. On the other hand, there are three mechanisms for energy loss due to gaseous absorption, droplet absorption, and single scattering. The balance of these mechanisms determines whether the emergent energy at the cloud top and bottom is greater or less than the respective boundary condition at the cloud bottom and top. Of course, the boundary conditions themselves are an important contribution to the emergent energy at the opposite side of the cloud. The upper boundary condition is totally unaffected by the earth's surface. For the lower boundary condition, quite the opposite is true. When the emissivity (and hence the reflectivity) of the earth's surface varies, three of the four contributing terms for the lower boundary condition are influenced. These include (1) the surface emission term, (2) the atmospheric contribution reflected from the surface of the earth, and (3) the emergent energy from the cloud bottom reflected from the earth's surface. Finally, it is important to note that the emission terms are generally stronger near the cloud bottom due to the temperature gradient within the cloud; and throughout the cloud, emission by water droplets is greater than emission by atmospheric molecules.

Introducing a cloud into the atmosphere will then result in increased emission within the cloud layer and decreased transmission through the cloud layer. The three contributions to the cloud top

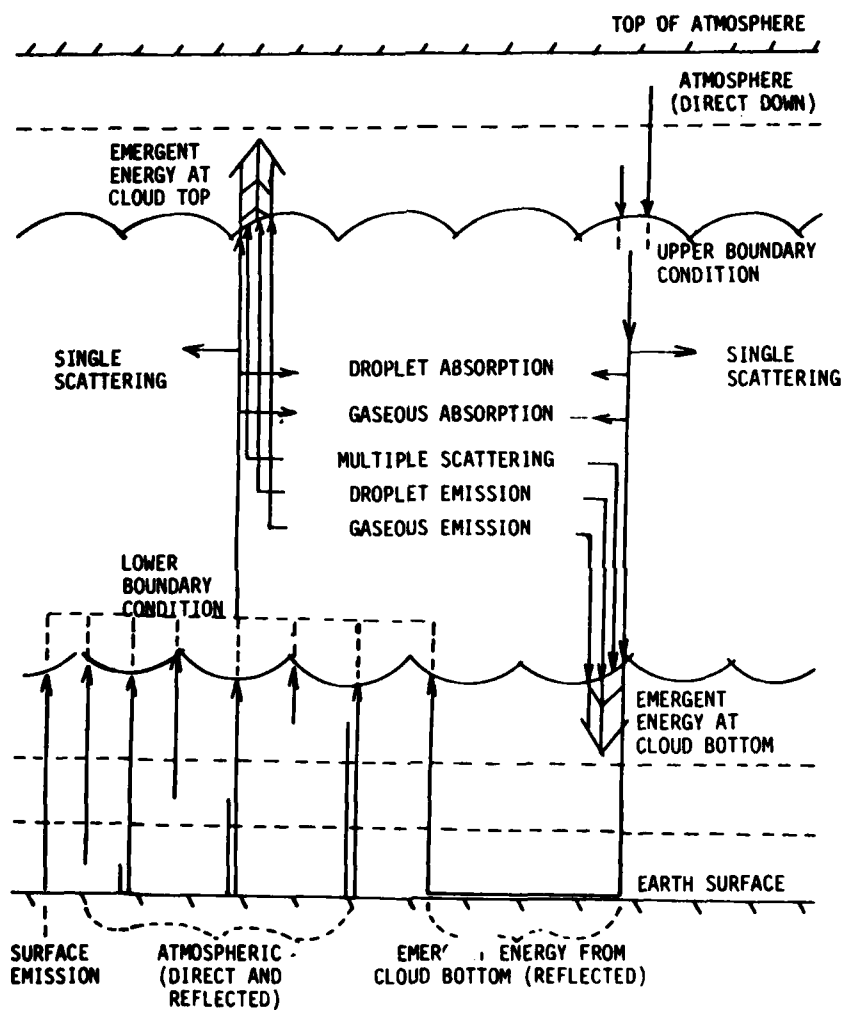


Figure 29. Radiative transfer through a cloud layer.

brightness temperature are emission by the earth's surface, atmospheric emission and emission from within the cloud layer.

To remove the surface term, atmospheric contribution, or cloud emission, we simply set the surface temperature, boundary conditions, or cloud temperature equal to zero, respectively. Each contribution can be analyzed separately by removing the other two.

In Figures 30 through 33 the effects of cloud thickness and liquid water content on the three contributions to the cloud top brightness temperature over the ocean are displayed. The solid line is the graphical sum of these components and the X's are computed cloud top temperatures directly from the microwave transfer program for randomly selected cases. Clearly, for channel 1 the addition of energy by droplet emission is greater than the energy loss due to droplet absorption and single scattering for liquid water contents greater than 0.01 gm cm^{-2} . This explains the region of increasing brightness temperature for channel 1 observed in Figure 11. For liquid water content greater than 0.20 gm cm^{-2} the droplet emission reaching the cloud top is nearly constant and represents almost all of the energy reaching the cloud top. Although the surface emission and atmospheric contributions to cloud top brightness temperature are small in this region, they are decreasing with increasing liquid water content so the cloud top brightness temperature decreases slightly also. This explains the slight decrease in the brightness temperature for channel 1 in this region. The fact that cloud emission accounts for almost all of the cloud top brightness temperature for water content greater than 0.20 gm cm^{-2} explains why the brightness temperatures over ocean and over land are nearly the same in this region. The cloud emission must be less than the reduction

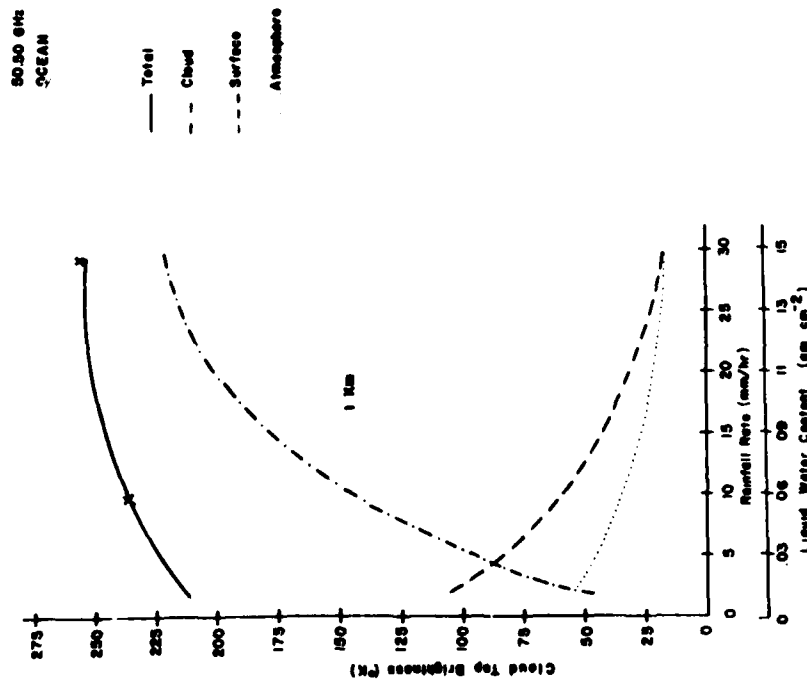


Figure 30. Component contributions to 50.50 GHz cloud top brightness for a 1 km thick cloud over ocean (mid-latitude Spring/Fall profile).

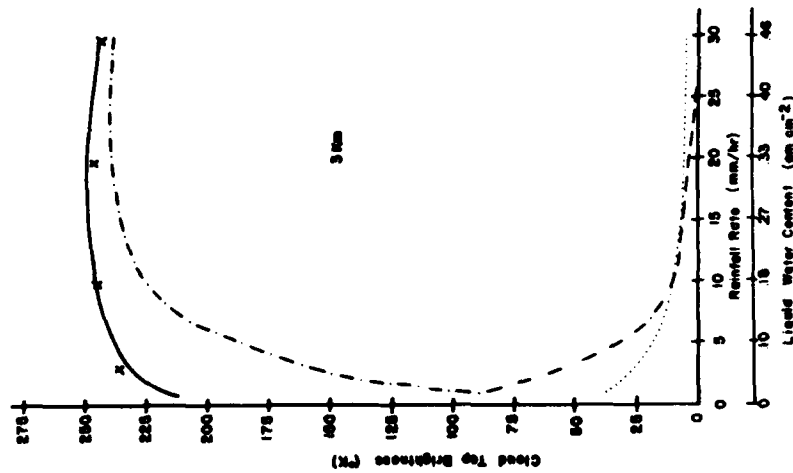


Figure 31. Component contributions to 50.50 GHz cloud top brightness for a 3 km thick cloud over ocean (mid-latitude Spring/Fall profile).

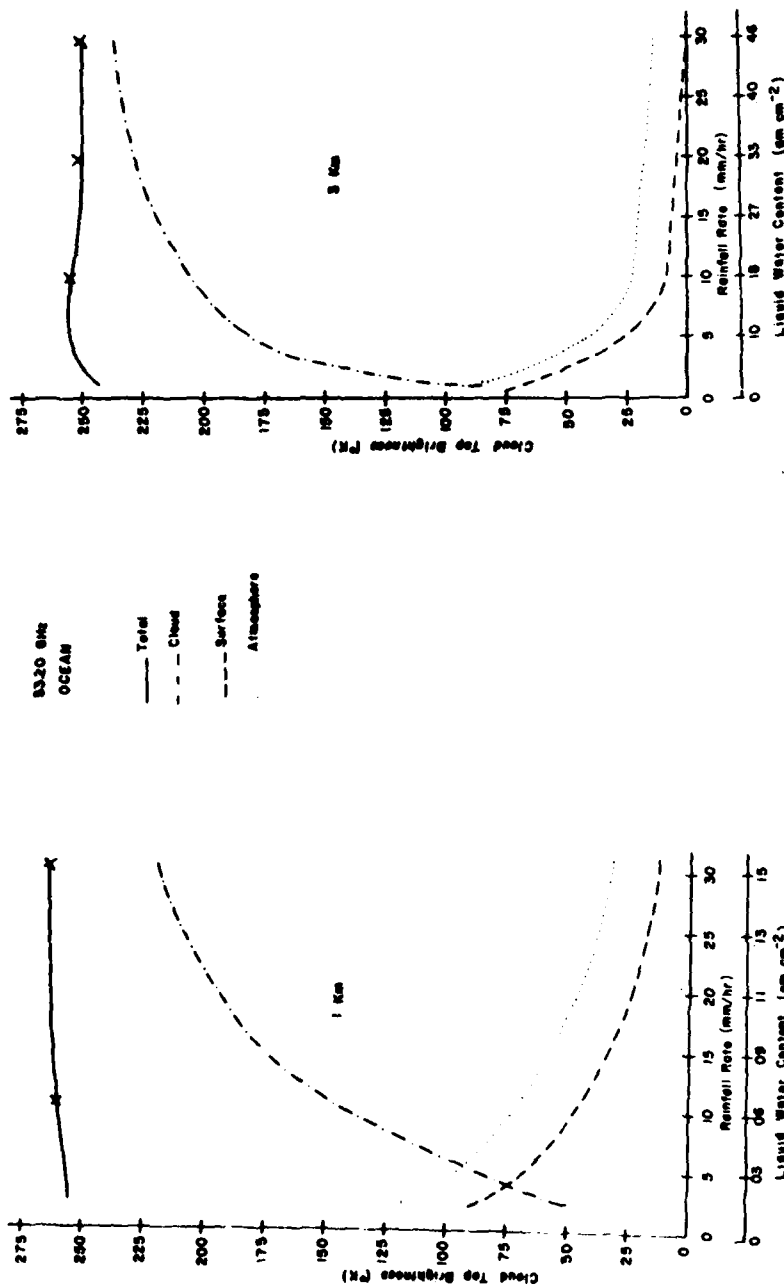


Figure 32. Component contributions to 53.20 GHz cloud top brightness for a 1 km thick cloud over ocean (mid-latitude Spring/Fall profile).

Figure 33. Component contributions to 53.20 GHz cloud top brightness for a 3 km thick cloud over ocean (mid-latitude Spring/Fall profile).

in the surface and atmospheric terms for liquid water content less than 0.01 gm cm^{-2} in order to explain the slight decrease in brightness temperatures in this region.

The cloud emission curves remain the same over land. As previously noted, the surface term is nearly twice as large over land while the reflected terms are only one sixteenth as large. The direct atmospheric contribution remains the same. The net result is that the upwelling brightness temperature at the cloud bottom over land is significantly larger than that over ocean. For example, our numerical experiments show that the lower boundary condition over land for a 1 km thick cloud ranges from 280° K to 285° K for rainfall rates of 1 to 30 mm/hr. Over ocean the range is 200° K to 240° K for the same cloud thickness and rainfall rates. Since the loss of energy within the cloud layer is proportional to the energy incident at the cloud base, a greater reduction takes place over land. This increased reduction is associated with the continuously decrease of brightness temperatures with increasing liquid water content over land for channel 1 as shown in Figures 6 through 10. In these cases, effects of the cloud transmittance is larger than the cloud emission.

For channel 2 the surface contribution to the cloud top brightness temperature over ocean is less than that for channel 1 due to the differences in transmittance as illustrated in Figures 30 and 32 and Figures 31 and 33 for 1 km and 3 km cases, respectively. The atmospheric terms are increased, however, due to the shape of the weighting function for this channel and the high reflectivity of the ocean surface. Although the cloud emission contribution for the 3 km case is slightly greater for channel 1 over channel 2, the net result is a higher cloud top brightness temperature for channel 2. The lower boundary condition is also higher for the same reasons, resulting in a greater reduction in the surface and atmospheric terms as they traverse

the cloud. Thus, for liquid water contents less than about 0.05 gm cm^{-2} (see Figure 12) at which the cloud emission is small, we observe the deepening of the reduction in the brightness temperature. The greater reduction in the surface and atmospheric terms also explains why the increase in brightness temperature for liquid water content of 0.05 through 0.15 gm cm^{-2} is not as large as the corresponding increases for channel 1. Note that the cloud emission is only slightly larger for channel 1 than channel 2 in the 3 km case and for the 1 km case it is approximately the same for both channels.

Based on the numerical experiments carried out for the study of the effects of the liquid water content on the surface, atmospheric, and cloud emission contributions to the cloud top brightness temperature over land, we find that nearly the same discussion applies for channel 2 over land as for channel 1. Thus, no plots similar to Figures 30 - 33 will be presented but rather some general discussions are made here. The surface terms increases due to the change in emissivity of the earth's surface. The increase is smaller than for channel 1, however, due to transmittance differences. The atmospheric terms decrease due to the change in reflectivity of the earth's surface. The decrease is larger than for channel 1 due to the shape of the weighting functions. The net result is that the lower boundary condition increases for channel 2 over land, but the increase is much reduced for channel 1. For example, for a 1 km thick cloud and rainfall rates of 1 and 30 mm/hr the lower boundary condition for channel 2 is nearly constant at 285° K over land but ranges from 250° K to 265° K over ocean. Once again, this increase in the lower boundary condition results in increased reduction through the cloud. Thus, we find that the brightness temperatures are continuously decreasing with increasing rainfall rates.

Channels 3 and 4 are relatively unaffected by the emissivity of the earth's surface. This is due to the fact that the transmittance from the earth's surface to the top of the atmosphere is only 0.023 for channel 3 and 0.003 for

channel 4 based on the results from the transmittance program. This fact explains why the brightness temperatures for channels 3 and 4 do not vary much between land and ocean surfaces. In this study the highest cloud top modeled was set at 6 km. This is just below the peak of the weighting function for channel 3 (7 km), but still well below the peak for channel 4 (10.5 km). As the cloud top moves into the energy source regions for these channels, emission by the cloud is lower since the temperatures are lower at higher altitudes. Furthermore, the transmittance from 6 km to the top of the atmosphere is less than 0.3 for both channels. Therefore, within the energy source region occupied by the upper portion of the cloud, the extinction mechanisms dominate over the mechanisms for the addition of energy, leading to slight decreases in the brightness temperature.

The interpretation of the cloud position dependence illustrated by Figures 15 through 20 is quite straightforward. Note that the lowest cloud base examined is at 1 km where the brightness temperatures over land for channels 1 and 2 are 265.5° K and 255.2° K, respectively. These values are higher than clear column values due to the fact that cloud emission exceeds extinction for a 2 km thick cloud with a base at 1 km for liquid water contents greater than 0.01 gm cm^{-2} (see Figures 30 through 33). The liquid water content for the 2 km thick L-Model cloud is 0.23 gm cm^{-2} . Thus, as the cloud is moved higher in the atmosphere, two changes take place. The emission by the cloud reduces due to the decrease in the cloud temperature. Meanwhile, the lower boundary condition increases because a greater portion of the energy source region for channels 1 and 2 is below the cloud base. This results in the increased energy reduction within the cloud layer. The net result is that brightness temperatures for channels 1 and 2 decrease as the cloud is moved higher up in the atmosphere. Eventually, the cloud is raised high enough that nearly all of the energy source regions for channels 1 and 2 are already below the cloud and cloud

temperatures are sufficiently low that emission by the cloud is very small. Then moving the cloud still higher reveals very little effect so that brightness temperatures are nearly constant. The same discussion applies overocean except that overocean even the 1 km cloud base shows a reduction in brightness temperatures as compared with the clear column values.

Channels 3, 4, 6 and 7 have weighting functions which peak higher in the atmosphere, reducing the effects of surface emissivity. Therefore, the effects of a 2 km thick cloud are basically the same over ocean and land surfaces. For clouds well below the peak of the weighting functions they show no effect on the brightness temperature. This is because there is no significant energy source below the cloud to be reduced by the cloud extinction. Also, the transmittance from the cloud top to the top of the atmosphere is sufficiently small that emission by the cloud may be negligible. As the cloud moves into the energy source region, an intricate trade-off of cloud emission and extinction takes place. For channels 3 and 4 this trade off holds the brightness temperatures very nearly constant until the cloud is well into the energy source region. For channels 6 and 7 a very slight increase in the brightness temperatures is noted. This increase is less than 0.2° K, however, which is below the noise level of the SSM/T. Finally, as the cloud is moved above the peaks of the weighting functions, brightness temperatures decrease, eventually becoming near constant as for channels 1 and 2.

SECTION 5

TEMPERATURE PROFILE RETRIEVAL EXERCISES

The main objective of the DMSP SSM/T microwave sounders has been to derive accurate temperature profiles in all weather conditions for operational use. The prime advantage of microwave temperature sounders over infrared sounders is that the longer microwaves are much less affected by clouds and precipitation. It seems therefore important to investigate the effects of clouds and precipitation on the temperature retrieval program. In this section, we first present the hypothetical temperature inversion in precipitating atmospheres using the brightness temperatures calculated from the microwave radiative transfer program. The retrieval program adopted in this study is the statistical method developed for the Air Force Global Weather Central intended for operational use. We also report a number of case studies using the real DMSP SSM/T data that we recently secured for clear, cloudy and precipitating cases.

5.1 Temperature Profile Retrieval Using Simulated Brightness Temperatures

The program that we use for temperature retrieval exercises is based on the statistical method described by Rigone and Stogryn (1977). The computer package for the statistical method was kindly provided to us by the Air Force Global Weather Central.

In the statistical method, the surface emissivity effect is

first removed so that the retrieval method could be applied to all surface conditions. For the purpose of outlining the method, we define

$$T_u(\nu) = \int_0^{\infty} T(z) \frac{\partial T_{\nu}(z, \infty)}{\partial z} dz, \quad (5.1)$$

and

$$T_d(\nu) = \int_{\infty}^0 T(z) \frac{\partial T_{\nu}(0, z)}{\partial z} dz, \quad (5.2)$$

so that Eq. (2.9) can be rewritten as follows:

$$T_B(\nu) = \epsilon_{\nu} T_s T_{\nu}(0) [1 - T_d(\nu)/T_s] + T_a(\nu), \quad (5.3)$$

where

$$T_a(\nu) = T_u(\nu) + T_d(\nu) T_{\nu}(0). \quad (5.4)$$

In Eq. (5.3), the second term in the right-hand side denotes the contribution to the upwelling brightness temperature caused by the atmosphere only, and the surface effects are contained in the first term. Since channel 1 centered at 50.5 GHz has a weighting function peak at the surface, it is utilized in the context of removing the surface contribution for other channels. Based on Eq. (5.3), we may define the contribution to the brightness temperature caused by the atmosphere only for channels 2-7 in the form

$$T_a(\nu_j) = T_B(\nu_j) - [T_B(\nu_1) - T_a(\nu_1)] a(\nu_j), \quad j=2,3,\dots,7, \quad (5.5)$$

where

$$a(v_i) = \frac{\epsilon_{v_i} T_s^{T_{v_i}}(p_s) [1 - T_d(v_i)/T_s]}{\epsilon_{v_1} T_s^{T_{v_1}}(p_s) [1 - T_d(v_1)/T_s]}, \text{ and } a(v_1) = 1.$$

In the statistical approach, it is generally assumed that the derivation of the predicted parameter T_i from the climatological mean \bar{T}_i may be expressed as a linear combination of the deviation of the measured data. Upon finding a linear operator \vec{D} which will yield a minimum mean square deviation of the predicted temperature profile \hat{T}_i from the true temperature profile T_i in a statistical sense, the predicted temperature profile may be obtained. The linear operator, called the predictor matrix, may be expressed in terms of a covariance matrix, which can be constructed experimentally by collecting coincidences of radiances derived from remote sounders with temperature values obtained from direct soundings. The measured data, in the present case is \hat{T}_a given by Eq. (5.5). Thus, we write

$$\begin{aligned} (\hat{T}_i - \bar{T}_i) &= \sum_j D_{ij} (\hat{T}_{aj} - \bar{T}_{aj}) \\ &= \sum_j D_{ij} [\hat{T}_{Bj} - (\hat{T}_{B1} - \hat{T}_{a1}) a_j - \bar{T}_{aj}] \\ &= \sum_{j=1} D_{ij} \hat{T}_{Bj} - \hat{T}_{B1} \sum_j D_{ij} a_j + \sum_j D_{ij} (\hat{T}_{a1} a_j - \bar{T}_{aj}). \end{aligned} \quad (5.6)$$

where \bar{T}_a denotes the mean measured value and we note that T_{B1} is not defined in Eq. (5.5), and so the first terms contain $j=2, \dots, 7$. In matrix notations, we find

$$\vec{\hat{T}} = \vec{D}' \vec{\hat{T}}_B + \vec{R}, \quad (5.7)$$

where

$$\vec{R} = \vec{T} + \hat{T}_a \vec{D} \vec{a} - \vec{D} \vec{T}_a,$$

and the linear operator \vec{D}' is a matrix whose first column is $-\vec{D} \vec{a}$ and whose remaining columns are the columns of \vec{D} . It is clear that the retrieval technique contains elements depending mainly on the atmosphere but not on the surface, and so it should be valid over land, water, or mixed surface conditions. As pointed out previously, the \vec{D} and \vec{R} may be determined from a large number of upper air soundings for a wide range of meteorological conditions which have been achieved over the years and the brightness temperatures calculated for a given atmosphere.

Shown in Figure 34 is an exercise of temperature retrieval using the statistical covariance method. The mid-latitude Spring/Fall profile of a standard atmosphere (solid curve) is used and the observed brightness temperatures used for the seven SSM/T channels are values theoretically calculated. The exercise has been carried out for cases over ocean and land. It is apparent that the procedure outlined above has very successfully removed surface effects from the temperature retrieval. Also shown are the temperature retrievals when a 2 km thick precipitation layer with a base height set at 1 km, having various rainfall rates, have been added to the atmosphere. It is seen that the surface temperature suffers increased degradation as the rainfall rate increases. Based on these analyses, it seems that large errors in the recovered temperature profile may be anticipated, even with microwave sounders, when the atmosphere within the satellite field-of-view contains precipitation and heavy clouds.

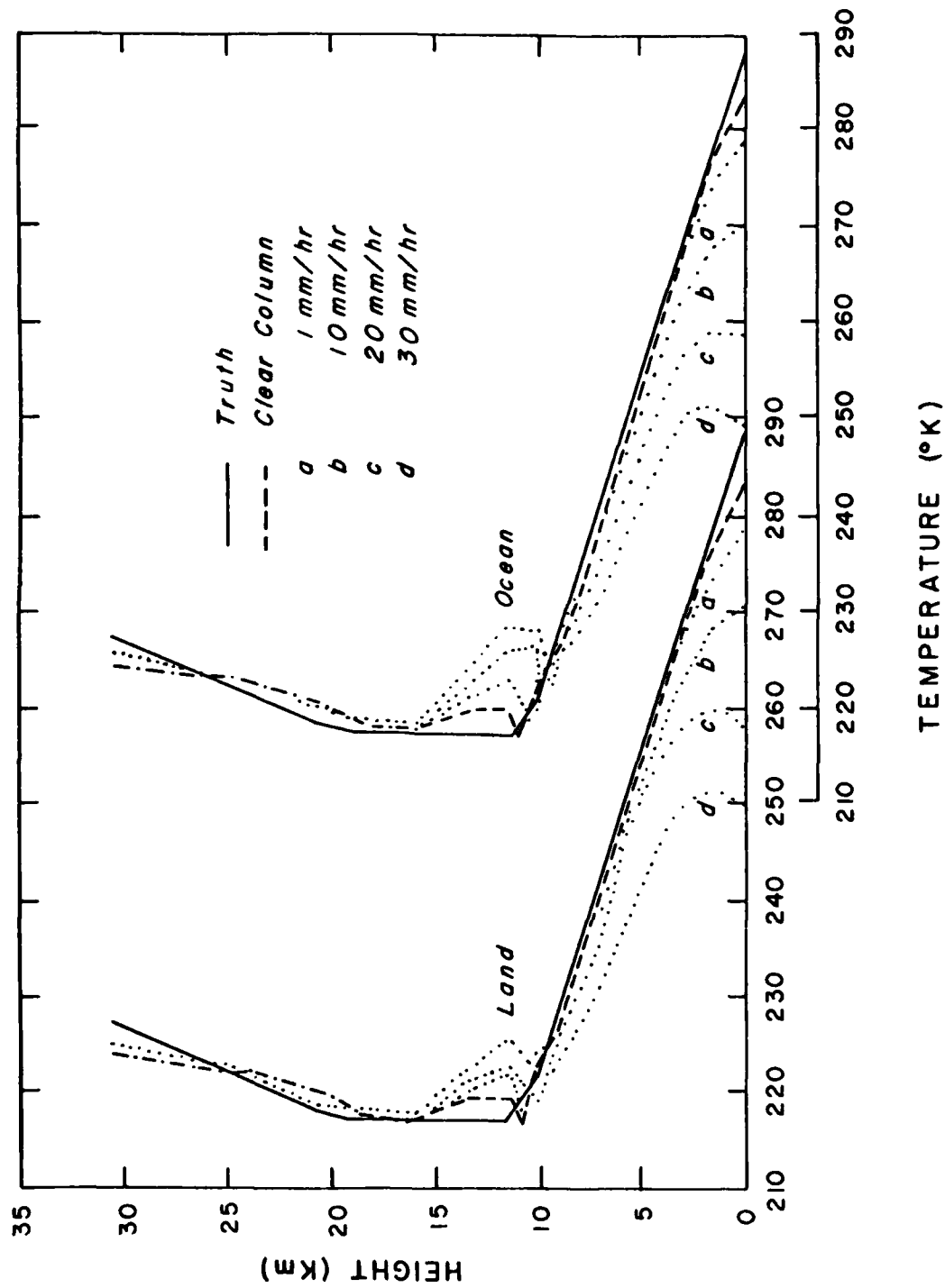


Figure 34. Hypothetical temperature retrieval exercise over land and ocean using the statistical covariance method.

5.2 Temperature Profile Retrieval Using DMSP SSM/T Data

The statistical method for temperature retrieval described in Section 5.1 was applied to a number of cases where SSM/T and radiosonde data were both available. Two days, 30 October 1979 and 23 November 1979, were chosen during which significant cloud and precipitation events were present over the continental United States. The cases selected are presented in Table 15 where the satellite pass times are the actual observation times of the SSM/T instrument. We find that the scan times are generally between 0000 Zulu and 0600 Zulu on these two days. Thus, the 0000 Zulu radiosonde observations were deemed most representative and were used in the comparisons.

Figure 35 shows the retrieved temperature profiles for the four clear cases; two on 30 October 1979 and two on 23 November 1979. Except for the Greensboro case, the retrieved temperature profiles when they are compared with nearby radiosonde data appear to be reasonably good in view of the statistical method used. The failure in the retrieval for the Greensboro case seems largely due to the fluctuated temperature profile that occurred in the atmosphere. Generally, we found that the statistical method is working properly when the actual profile is smooth and when no inversion is present.

The retrieved temperature profiles under cloudy conditions are illustrated in Figure 36. Apparently, the temperature profile retrieval program using the microwave frequencies in the 60 GHz oxygen band is affected insignificantly by non-precipitating clouds. Compared with the temperatures obtained from radiosondes, the retrieved patterns involving clouds are similar to those under clear conditions. Note that in each diagram, the percentage of cloud cover is depicted.

Table 15. Selected cases.

Station Name	Satellite Pass Time	Latitude (°N)	Longitude (°W)	Case Type
Centerville Alabama (AL)	0346 Z 30 Oct 79	32.54	87.15	Clear
Little Rock Arkansas (AR)	0346 Z 30 Oct 79	34.44	92.14	Clear
Greensboro North Carolina (NC)	0251 Z 23 Nov 79	36.03	79.57	Clear
Glasgow Montana (MT)	0435 Z 23 Nov 79	48.13	106.37	Clear
Bismark North Dakota (ND)	0350 Z 30 Oct 79	46.46	100.45	Cloudy
Medford Oregon (OR)	0531 Z 30 Oct 79	42.22	122.52	Cloudy
Green Bay Wisconsin (WI)	0253 Z 23 Nov 79	44.29	88.08	Cloudy
Monterrey Mexico (MEX)	0429 Z 23 Nov 79	25.52	100.12	Cloudy
Dodge City Kansas (KS)	0348 Z 30 Oct 79	37.46	99.58	Precipitating
Omaha Nebraska (NE)	0349 Z 30 Oct 79	41.22	96.01	Precipitating
Pittsburgh Pennsylvania (PA)	0252 Z 23 Nov 79	40.32	80.14	Precipitating
Spokane Washington (WA)	0435 Z 23 Nov 79	47.38	117.32	Precipitating

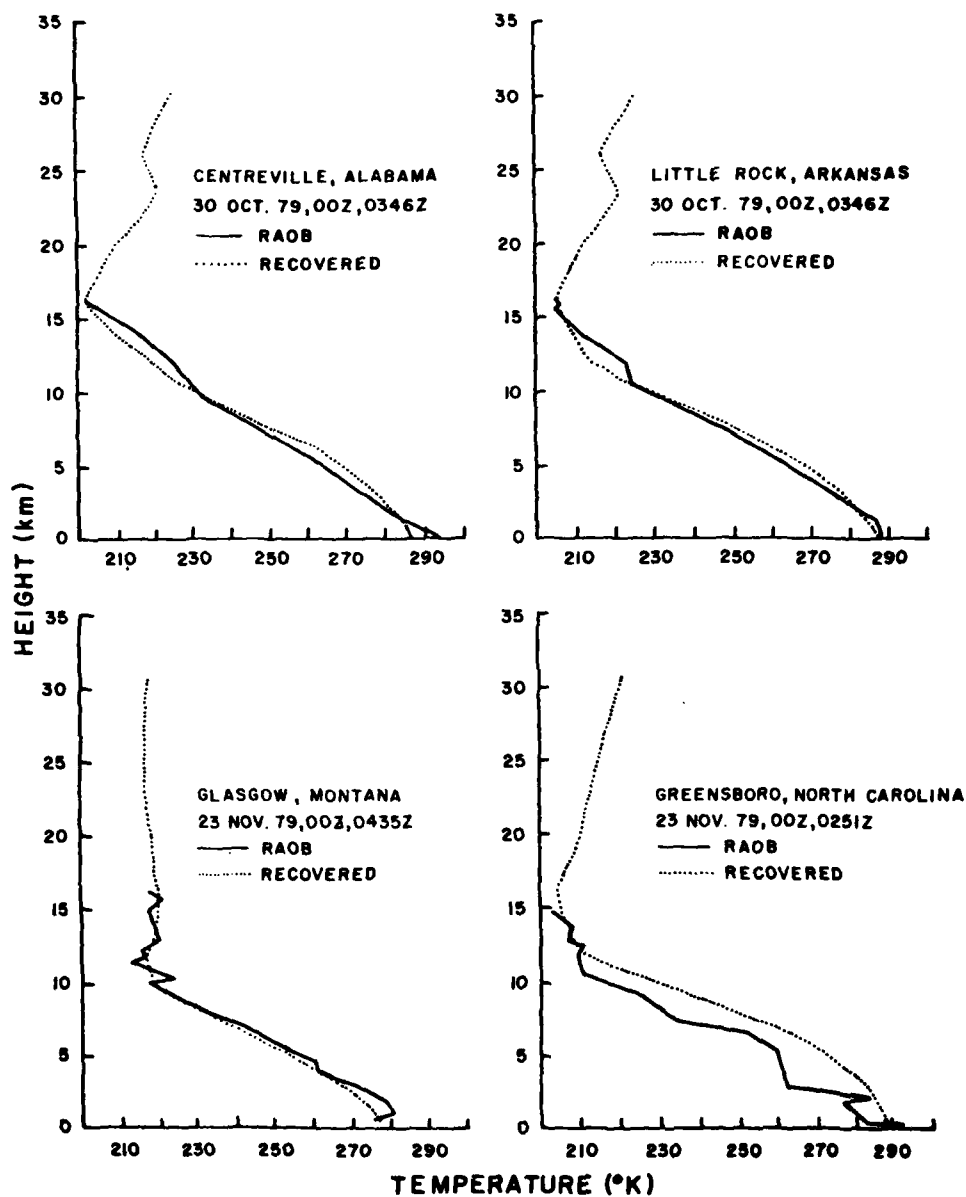


Figure 35. Comparisons of the retrieved temperature profiles (dots) with the radiosonde data (solid lines) for the four clear cases.

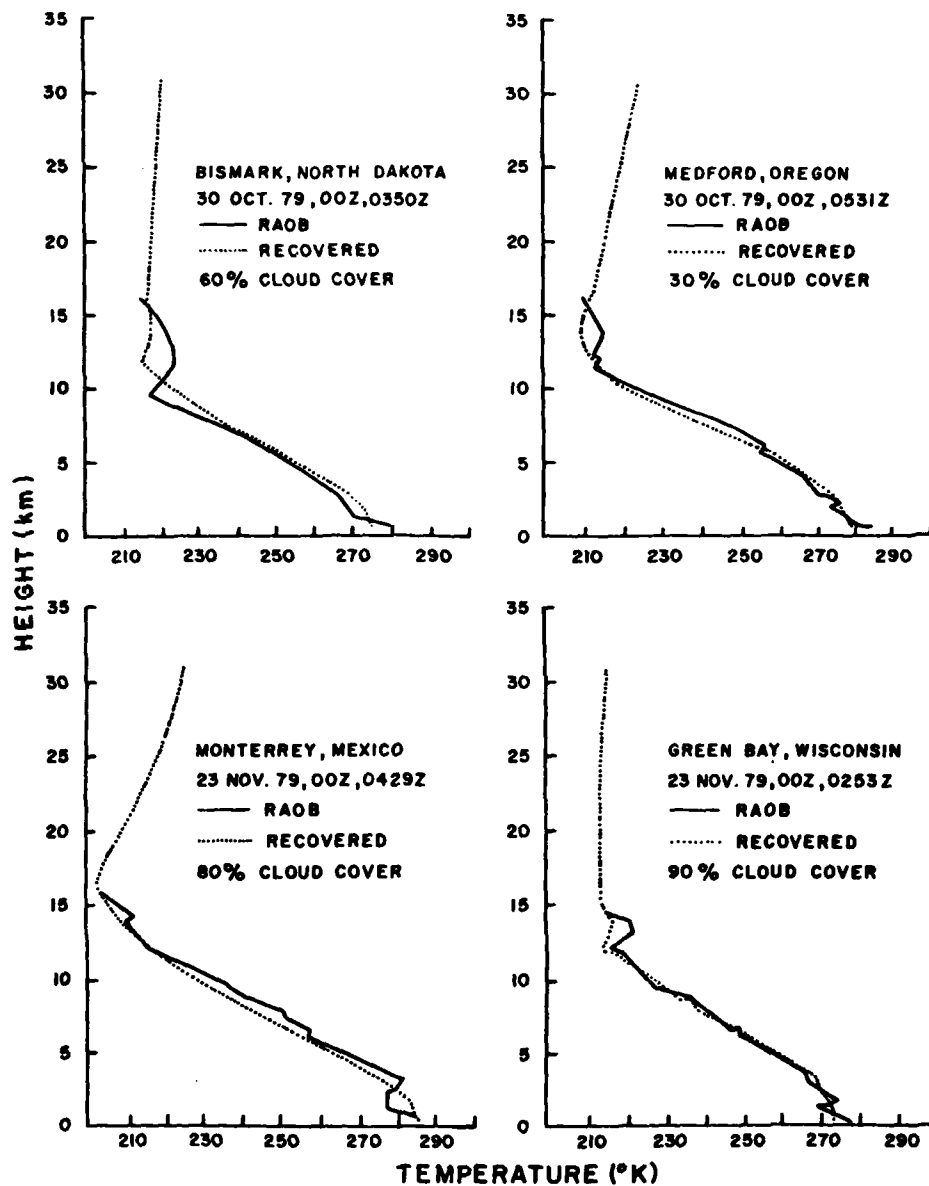


Figure 36. Comparisons of the retrieved temperature profiles (dots) with the radiosonde data (solid lines) for the four cloudy cases.

Generally, the retrieved and observed (radiosonde) temperatures in clear and cloudy conditions are within about 5° K.

In the final figure (Figure 37) we show the retrieved temperature profiles under precipitating conditions. Again, four cases are presented in this study. Two cases are selected from 30 October 1979; both have a 5 mm/hr rainfall rate with 30% cloud cover in the field of view of the SSM/T. In the other two cases, selected from 23 November 1979, both indicate a 1 mm/hr rainfall rate but with cloud covers varying from 50% to 80%. The most distinct feature in the retrieved temperature profiles using the statistical covariance method for precipitating cases is the significant and consistent deviation from the radiosonde data in the lower boundary layer where precipitation takes place. In the moderate 5 mm/hr rainfall rate cases, the differences between the retrieved and radiosonde temperature profiles near the surface are as large as $10 - 15^{\circ}$ K. It should be noted that precipitation in these two cases cover only about 30% within the field of view of SSM/T. As for the cases involving 1 mm/hr rainfall rate, about $5 - 10^{\circ}$ K differences near the ground are observed. The findings for these precipitation cases using the real SSM/T data are in general agreement with those described previously in the hypothetical temperature retrieval exercises. As shown in Figure 34 the surface temperatures suffer increased degradation from 10° to 20° K as the rainfall rate increases from 1 mm/hr to 20 mm/hr. These hypothetical analyses, of course, assume that the cloud covers the entire field of view.

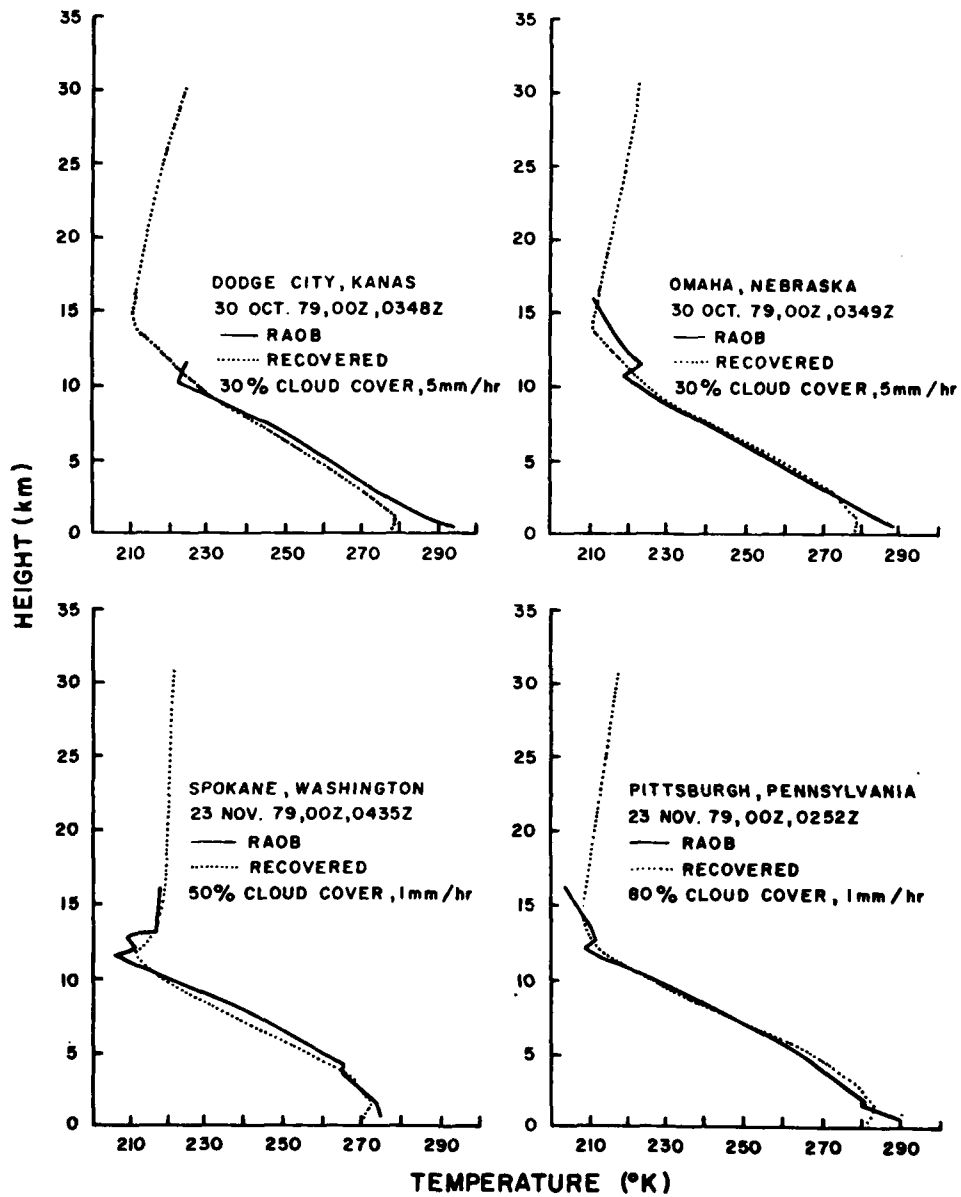


Figure 37. Comparisons of the retrieved temperature profiles (dots) with the radiosonde data (solid lines) for the four precipitation cases.

SECTION 6

CONCLUSION

In this study, we have developed a microwave radiative transfer program for cloudy atmospheres applicable to the DMSP SSM/T channels. The transfer program takes into account the simultaneous effects of multiple scattering and absorption by hydrometeors and absorption due to molecular oxygen and water vapor in the atmosphere.

Sensitivity analyses were carried out to investigate the effects of the rainfall rate, cloud thickness, and cloud location on the upwelling brightness temperature over land and ocean surfaces for two atmospheric profiles. The effects of precipitation on SSM/T channels 1 and 2 are shown to be significant depending on the cloud liquid water content (or rainfall rate), thickness, and surface emissivity. Over the land surface, increasing the cloud liquid water content and thickness reduces the upwelling brightness temperature for channels 1 and 2. For channels 3 and 4, unless high rainfall rates are involved, the reduction in the brightness temperature is normally insignificant. Over the ocean surface, however, the increase of the liquid water content and thickness may increase or decrease the upwelling brightness temperature relative to the clear column value. A significant variation for the brightness temperature is shown for channel 1. This conclusion is in general agreement with that

described by Wilheit et al. (1977) using a simple scattering program for a frequency of 37 GHz. Moreover, sensitivity analyses also reveal the importance of the position of the cloud layer in the atmosphere and the atmospheric temperature profile on the upwelling brightness temperature values.

In addition, investigation of the effects of precipitation on the temperature profile retrieval using both the theoretically simulated values and real data was carried out utilizing the seven SSM/T channels. The retrieval method adopted in this study is the statistical method developed at the Air Force Global Weather Central in which the surface effect is removed in the recovery program. The hypothetical retrieval exercises show that the temperatures close to the surface suffer increased degradation as the rainfall rate increases. This finding is also supported by the analysis employing the real SSM/T data for a number of case studies in which temperature profiles from radiosondes are available for comparison. Furthermore, the latter study indicates that nonprecipitating clouds have an insignificant affect on the microwave temperature retrieval with accuracy generally on the same order as in clear conditions.

Although the current study employs only four precipitating cases in the analysis and may not be conclusive in view of the limited sample used, it appears that the effect of precipitation on the temperature profile retrieval using microwave frequencies is substantial and significant. Of course, the reliability of the statistical method for the temperature profile retrieval in clear atmospheres should be examined comprehensively and completely utilizing the data that are

available in different localities and seasons. In addition, in order to derive the temperature field over the global space, further studies concerning the influence of precipitating clouds on the temperature retrieval in the microwave region seem also warranted.

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APPENDIX

LISTING OF MICROWAVE TRANSFER PROGRAM

```

C*****
C   MAIN PROGRAM
C*****

PARAMETER NUMCHN=6, NUMLV=40, NG=16, NMOST=NUMLV+10
DIMENSION T(NMOST), W(NMOST), TAUM(NMOST, NUMCHN), TAUS(NMOST, NUMCHN),
*      EMIS(NUMCHN), TANT(NUMCHN)
COMMON /ANGLE/ UM(NG), A(NG), PI
COMMON /SFCMS/ EMIS
COMMON /FREQ/ FMU(NUMCHN)
COMMON /HANDU/ H(NMOST), U(NMOST, NUMCHN)
COMMON /TAUVAL/ TAUM
COMMON /TANDW/ T, W
COMMON /PRES/ P(NMOST)
COMMON /NEWLEV/ ADDHT(10), NUMNEW, LVNUM
REAL*8 SFC(2)
DATA SFC/'LAND','WATER'/
DATA KTAU1, KTAU2/'TAU(','GHZ)'/
CALL PROFIL
CALL TRANMW(TAUM, TAUS, 0., 1, NUMCHN, 1, 1.)
WRITE(6, 900) (KTAU1, FMU(J), KTAU2, J=1, NUMCHN)
DO 20 I=1, LVNUM
WRITE(6, 901) P(I), T(I), W(I), (TAUM(I, J), J=1, NUMCHN)
20 CONTINUE
READ(5, 903) (EMIS(I), I=1, NUMCHN)
60 CALL ANTEMP(TAUM, TAUS, T, T(LVNUM), EMIS, TANT, LVNUM, 1, NUMCHN)
WRITE(6, 902) SFC(1), EMIS, TANT
CALL BUFF(P, T, W, H, U)
900 FORMAT('1 P(MB) T(K.) W(G/KG)', 2X, 7(2X, A4, F5.2, A4), '/')
901 FORMAT(2(2X, F6.1), 2X, E8.3, 5X, 6(E10.5, 5X), E10.5)
902 FORMAT(1H0, A0, '... EMISSIVITY', 4X, 6F15.3, /, 1H0, 12X,
*      'ANTENNA TEMP.', 1X, 7F15.3)
903 FORMAT(' ')
STOP
END

```

SUBROUTINE PROFIL

```

C*****
C THIS SUBROUTINE IS TO READ OBSERVED PROFILES (TEMPERATURE K, MIXING
C RATIO G/KG, PRESSURE MB, AND HEIGHT KM) OF ATMOSPHERE, AND
C INTERPOLATE THE PROFILES TO THE DESIRED PRESSURE LEVELS (40 STANDARD
C LEVELS).
C THE LEVELS CAN BE ADDED UP TO 10 LEVELS OF YOUR CHOICE.
C*****

PARAMETER NUMLV=40, NMOST=NUMLV+10
DIMENSION T(NMOST), W(NMOST), H(NMOST), OBPR(40)
*      , OBH(40), OBTEMP(40), OBW(40)
COMMON /PRES/ P(NMOST)
COMMON /HANDU/ H, U
COMMON /TANDW/ T, W
COMMON /NEWLEV/ ADDHT(10), NUMNEW, LVNUM
200 FORMAT(' ')
201 FORMAT(1X, 'NLV OBS TO HIGH. MAX VALUE EQUALS 40. ')

```

```

      READ(5,200) NLV OBS
      IF(40-NLV OBS)202,203,204
202 PRINT 201
      RETURN 0
203 IF(NLV OBS)105,105,204
204 N=NLV OBS
      READ(5,200)(OBTEMP(I),I=N,1,-1)
      READ(5,200)(OBW(I),I=N,1,-1)
      READ(5,200)(OBPR(I),I=N,1,-1)
      READ(5,200)(OBH(I),I=N,1,-1)
      WRITE(6,301) OBTEMP
      WRITE(6,301) OBW
      WRITE(6,301) OBPR
      WRITE(6,301) OBH
301 FORMAT(10F8.3)
302 FORMAT(1X,'PROFIL USED CLIMO DATA (MID LAT SPRING/FALL) ABOVE ',
      * F7.2,' Mh')
      K=1
      DO 100 I=1,NUMLV
      IF (OBPR(K).EQ.0.0) GO TO 100
10 IF (OBPR(K).EQ.P(I)) GO TO 90
      IF (OBPR(K).LT.P(I)) GO TO 101
      IF (OBPR(K).GT.P(I)) GO TO 89
      IF (K.EQ.1) GO TO 100
      DELT=OBTEMP(K)-OBTEMP(K-1)
      DELW=OBW(K)-OBW(K-1)
      DLNP=ALOG(OBPR(K)/OBPR(K-1))
      T(I)=DELT/DLNP*ALOG(P(I)/OBPR(K-1))+OBTEMP(K-1)
      W(I)=DELT/DLNP*ALOG(P(I)/OBPR(K-1))+OBW(K-1)
      CALL HINTHP(OBPR(K),OBPR(K-1),P(I),OBH(K),OBH(K-1),H(I))
      GO TO 100
89 K=K+1
      IF (K.EQ.2) PRINT 302,P(I)
      GO TO 10
90 T(I)=OBTEMP(K)
      W(I)=OBW(K)
      H(I)=OBH(K)
99 K=K+1
100 CONTINUE
101 CONTINUE
105 READ(5,200) NUMNEW
      NEW=NUMNEW
      LVNUM=NUMLV+NUMNEW
      IF (NUMNEW.EQ.0) GO TO 210
      READ(5,200)(ADDHT(I),I=1,NEW)
      NTOP=NUMLV
      K=1
      DO 206 I=NUMLV,1,-1
110 IF(H(I).LE.ADDHT(K)) GO TO 206
      J=I+1
      DO 175 L=NTOP,J,-1
      P(L+1)=P(L)
      T(L+1)=T(L)
      W(L+1)=W(L)
      H(L+1)=H(L)
175 CONTINUE
      NTOP=NTOP+1
      H(J)=ADDHT(K)
      CALL NEWLVL(J)
      K=K+1
      IF(K.GT.NUMNEW) GO TO 210
      GO TO 110
206 CONTINUE
210 CONTINUE
      RETURN
      END

```

```

SUBROUTINE HINTRP(SPTOP,SPBOT,SPP,SZTOP,SZBOT,SZZ)
C*****
C THIS SUBROUTINE IS TO INTERPOLATE THE OBSERVED HEIGHTS TO THE HEIGHTS
C OF DESIGNATED PRESSURE LAYERS.
C *****

```

```

IMPLICIT DOUBLE PRECISION (A-H,O-Z)
DATA TOL/1.0E-4/RADKM/6.37122D3/N/11/
REAL SPTOP,SPBOT,SPP,SZTOP,SZBOT,SZZ
DIMENSION ZMATRX(50,50),F(50),ZMAT(2500),COLF(50)
DIMENSION ARKAY(12)/1.00D-1,1.25D-1,1.40D-1,1.60D-1,1.55D-1,
* 1.70D-1,1.55D-1,1.47D-1,1.42D-1,1.32D-1,1.21D-1,1.315D-1/
ENTER=1
1 ZHOT=DNLE(SZBOT)
ZTOP=DNLE(SZTOP)
PHOT=DNLE(SPBOT)
PTOP=DNLE(SPTOP)
PP=DBLE(SPP)
ZTST=ZHOT
IF(ZTST.GT.ZTOP) ZTST=ZTOP
IF(1.00D.LE.ZTST) GO TO 2
CONS=(SZTOP-SZBOT)/ALOG(SPTOP/SPBOT)
SZZ=ALOG(SPP/SPBOT)*CONS+SZHOT
GO TO 99
2 KK=(PTOP-PBOT)/(1.00/(ZTOP+RADKM)-1.00/(ZHOT+RADKM))
KNUM=1.00-RADKM*(PP-PHOT)/RK-RADKM/(ZHOT+RADKM)
RDEN=(PP-PHOT)/KK+1.00/(ZHOT+RADKM)
IF(ENTER.EQ.1) SZZ=SNGL(KNUM/RDEN)
IF(ZTST-1.00)99,5,5
5 LVLA=0
ZTST=DNLE(SZZ)
IF(ZTST.GT.6.00D1) LVLA=LVLA+1
IF(ZTST.GT.4.75D1) LVLA=LVLA+1
IF(ZTST.GT.4.25D1) LVLA=LVLA+1
IF(ZTST.GT.3.75D1) LVLA=LVLA+1
IF(ZTST.GT.3.25D1) LVLA=LVLA+1
IF(ZTST.GT.2.45D1) LVLA=LVLA+1
IF(ZTST.GT.2.15D1) LVLA=LVLA+1
IF(ZTST.GT.1.65D1) LVLA=LVLA+1
IF(ZTST.GT.1.45D1) LVLA=LVLA+1
IF(ZTST.GT.1.25D1) LVLA=LVLA+1
IF(ZTST.GT.6.50D) LVLA=LVLA+1
IF(ZTST.GT.0.00D) LVLA=LVLA+1
A=ARKAY(LVLA)
ZINC=(ZTOP-ZBOT)/FLOAT(N-1)
DO 10 I=1,N
ZMATRX(I,1)=1.
10 CONTINUE
DO 20 I=1,N
ZMATRX(I,2)=1./(ZBOT+FLOAT(I-1)*ZINC)
DO 15 J=3,N
ZMATRX(I,J)=ZMATRX(I,J-1)*ZMATRX(I,2)
15 CONTINUE
ZZ=1./ZMATRX(I,2)
GRAD=ZZ+RADKM
F(I)=DEXP(-A*ZZ)/(GRAD*GRAD)
COEF(I)=F(I)
20 CONTINUE
K=0
DO 23 I=1,N
DO 21 J=1,N
K=K+1
ZMAT(K)=ZMATRX(I,J)
21 CONTINUE
23 CONTINUE
CALL SIMU(ZMAT,COLF,N,KS)
SUMBOT=COEF(1)*ZRLT+COEF(2)*NLOG(ZBOT)
SUMTOP=COEF(1)*ZTOP+COEF(2)*NLOG(ZTOP)
HOT=1./ZHOT
TOP=1./ZTOP

```

```

      DO 60 J=3,N
      COEF(J)=-COEF(J)/FLOAT(J-2)
      SUMHOT=SUMHOT+COEF(J)*ROT
      BOT=BOT/ZHOT
      SUMTOP=SUMTOP+COEF(J)*TOP
      TOP=TOP/ZTOP
60  CONTINUE
      AINT=SUMTOP-SUMHOT
      RK=(PHOT-PTOP)/AINT
      GO TO (N1,96),IENTEN
61  CONS=(PHOT-PP)/RK+SUMHOT
      FI=CONS-SUMHOT
      FT=CONS-SUMTOP
62  ZMEAN=(ZHOT+ZTOP)/2.
      IF (ABS(ZTOP-ZBOT)-TOL)95,95,65
65  SUMEAN=COEF(1)*ZMLAN+COEF(2)*DLOG(ZMEAN)
      EAN=1./ZMEAN
      DO 66 J=3,N
      SUMEAN=SUMEAN+COEF(J)*EAN
      EAN=EAN/ZMEAN
66  CONTINUE
      FMEAN=CONS-SUMEAN
      IF (FMEAN)70,95,80
70  ZTOP=ZMEAN
      FI=FMEAN
      GO TO 62
80  ZBOT=ZMEAN
      FR=FMEAN
      GO TO 62
95  SZZ=SINGL(ZMEAN)
96  ZZ=DHLE(SZZ)
      SUMTOP=COEF(1)*ZZ+COEF(2)*DLOG(ZZ)
      EZZ=1./ZZ
      DO 97 J=3,N
      SUMTOP=SUMTOP+COEF(J)*EZZ
      EZZ=EZZ/ZZ
97  CONTINUE
      AINT=SUMTOP-SUMHOT
      PF=-RK*AINT+PHOT
      SPP=SINGL(PP)
      GO TO 99
      ENTRY PINTRP(SPTOP,SPROT,SPP,SZTOP,SZHOT,SZZ)
      IENTEN=2
      GO TO 1
99  CONTINUE
      RETURN
      END

```

```

SUBROUTINE SIMQ (A,B,N,KS)
C *****
C OBTAINS SOLUTION OF A SET OF SIMULTANEOUS LINEAR EQUATIONS ...
C A - MATRIX OF COEFFICIENTS STORED COLUMNWISE. THESE ARE
C DESTROYED IN THE COMPUTATIONS. THE SIZE OF MATRIX 'A'
C IS N X N.
C B - VECTOR OF ORIGINAL CONSTANTS (LENGTH N) WHICH IS REPLACED
C BY FINAL SOLUTION VALUES, VECTOR X.
C N - NO. OF EQUATIONS AND VARIABLES.
C KS - OUTPUT DIGIT
C 0 FOR NORMAL SOLUTION
C 1 FOR A SINGULAR SET OF EQUATIONS
C *****

DOUBLE PRECISION A,B,BIGA,TOL,SAVE
DIMENSION A(1),h(1)
TOL=0.0001
KS=0
JJ=-N
DO 60 J=1,N
  JY=J+1
  JJ=JJ+N+1
  BIGA=0.0001
  IT=JJ-J
  DO 20 I=J,N
    IJ=IT+1
    IF ((DAHS(BIGA)-DAHS(A(IJ))) 10,20,20
10  HIGA=A(IJ)
    IMAX=I
20  CONTINUE
    IF (DAHS(RIGA)-TOL) 30,30,40
30  KS=1
    RETURN
40  I1=J+N*(J-2)
    I1=IMAX-J
    DO 50 K=J,N
      I1=I1+N
      I2=I1+IT
      SAVE=A(I1)
      A(I1)=A(I2)
      A(I2)=SAVE
50  A(I1)=A(I1)/RIGA
      SAVE=h(IMAX)
      h(IMAX)=h(J)
      h(J)=SAVE/BIGA
      IF (J=N) 60,90,60
60  IQS=N*(J-1)
      DO 60 IX=JY,N
        IXJ=IQS+IX
        IT=J-IX
        DO 70 JX=JY,N
          IXJX=N*(JX-1)+IX
          JJX=IXJX+IT
70  A(IXJX)=A(IXJX)-(A(IXJ)*A(JJX))
80  h(IX)=h(IX)-(h(J)*A(IXJ))
90  NY=N-1
      IT=N*N
      DO 100 J=1,NY
        IA=IT-J
        IB=N-J
        IC=N
        DO 100 K=1,J
          B(IB)=B(IB)-A(IA)*B(IC)
          IA=IA-N
100  IC=IC-1
      RETURN
END

```

```

SUBROUTINE NEWLVL(K)
PARAMETER LEV=40, INT=6, NMOST=LEV+10
COMMON/PRES/PRE(NMOST)
COMMON/HANDU/ALT(1,NMOST),11(NMOST)
COMMON/TANDW/TEMP(NMOST),H2O(NMOST)
SPTOP=PRE(K+1)
SPHOT=PRE(K-1)
SZTOP=ALT(K+1)
SZBOT=ALT(K-1)
SZZ=ALT(K)
TESI=SZHOT
IF (TEST.GT.SZTOP) TEST=SZTOP
IF (TEST.GT.1) GO TO 10
RHOG=(SPTOP-SPHOT)/(SZTOP-SZHOT)
PRE(K)=SPHOT+RHOG*(SZZ-SZHOT)
GO TO 11
10 CALL PINTRP(SPTOP,SPHOT,PRE(K),SZTOP,SZHOT,SZZ)
11 SPP=PRE(K)
TEMP(K)=(TEMP(K+1)+TEMP(K-1))*ALOG(SPP/SPBOT)/ALOG(SPTOP/SPBOT)
* +TEMP(K-1)
H2O(K)=(H2O(K+1)+H2O(K-1))*ALOG(SPP/SPBOT)/ALOG(SPTOP/SPBOT)
* +H2O(K-1)
RETURN
END

```

```

SUBROUTINE TRANMW(TAW,TAWSQ,Z,NU1,NU2,KOUNT,UMQ)
C*****
C COMPUTE TRANSMISSIVITIES TO THE TOP OF THE ATMOSPHERE FROM EACH LEVEL
C OF THE PROFILE FOR EACH CHANNEL (ANTENNA GAIN CHARACTERISTICS INCLUDED)
C*****

```

```

PARAMETER NUMCHN=0,NUMLV=40,NG=16, NMOST=NUMLV+10
DOUBLE PRECISION TAU,TAX,TOW,T1,T2
COMMON/INPUT/ALT(NMOST),TEMP(NMOST),PRE(NMOST),H2O(NMOST),
* THINS(4),NTHICK,NHASE,TKN(NMOST,NUMCHN),N*LEV(10)
COMMON /STOKAL/ ALFA
COMMON /TANDW/ T,W
COMMON/HANDU/H(NMOST),U(NMOST,NUMCHN)
COMMON/NEWLEV/ADDT(10),NUMNEW,LVNUM
DIMENSION T(NMOST),w(NMOST),ALFA(NMOST,NUMCHN),TAX(NMOST,NUMCHN),
* TAXSQ(NMOST,NUMCHN),GSUM(NUMCHN)
DIMENSION DLS(126),TOW(127),SA(127),SZ(127),G(127,NUMCHN)
COMMON/GAIN/GANT(127,NUMCHN)
COMMON/PRES/ P(NMOST)
LEV1=LVNUM
NF=NUMLV
ZLAST=999.
DTK=.0174533
NANG=127
ISW=1
CRH=.17268

```

```

      IF (ISW) 10,10.1
1  I,INI=NANG/2
   MN=INIT
   MM=INIT+1
   CA=1.
   A=0.
   DO 2 I=1,INIT
     A=A+1.
     CH=LOS(A*DTR)
     DC=CA-CH
     DJ=.5*DC
     DCS(1+MM)=DD
     DCS(MM-I)=DD
2  CA=CB
   DO 7 KK=NU1,NU2
     DO 3 I=n4,127
3  G(I,KK)=GANT(I,KK)*GANT(I,KK)
     DO 4 I=1,63
4  G(I,KK)=G(128-I,KK)
     GS=0.
     G1=G(1,KK)
     DO 6 I=2,127
6  G2=G(I,KK)
     GS=GS+(G1+G2)*DCS(I-1)
     G1=G2
     GSUM(KK)=GS
7  CONTINUE
   SMAX=ASIN(1./CRH)/DTR
   ISW=0
10 IF (Z.EQ.ZLAST) GO TO 20
   DO 12 K=1,127
   DU=FLOAT(K-MM)
   SA(K)=AHS(Z+DU)
   SZ(K)=1.
   IF (SA(K).GT.SMAX) GO TO 12
   A=CHH*SIN(DTR*SA(K))
   OU=ASIN(A)
   SZ(K)=1./COS(OU)
12 CONTINUE
   ZLAST=Z
20 IF (2.LE.KOUNT) GO TO 30
   CALL RASEU(LVINUM)
30 DO 120 J=NU1,NU2
   TAUNU=0.
   DO 110 I=1,LEV1
   UU=U(I,J)/UM0
100 IF (UU) 101,101,102
101 TAU=1.
   GO TO 103
102 IF (UU.GT..89E1) GO TO 53
   GO TO 54
53 UU=.40E1+TAUNU
   TAUNU=TAUNU+.2
54 TAU=EXP(-UU)
103 DO 104 K=1,127
   TAX=TAU
   IF (SA(K).GT.SMAX) GO TO 104
   IF (TAX.LT..0001) GO TO 104
   IF (TAX.GT..9999) GO TO 104
   IF (SZ(K).EQ.1.) GO TO 104
   TAX=TAX**SZ(K)
104 TOW(K)=TAX
   GS=0.
   G1=G(1,J)
   T1=TOW(1)

```

```

      GO 106 K=2,127
      G2=G(K,J)
      T2=TOW(K)
      IF(SA(K).GT.SMAX) GO TO 105
      GS=GS+(G1*T1+G2*T2)*DCS(K-1)
105  G1=G2
106  T1=T2
      TAW(I,J)=GS/GSUM(J)
      TAU2=TAW(I,J)
      TAWR=.5*(TAU1+TAU2)
      TAWSG(I,J)=TAWR*TAWH
      TAU1=TAU2
110  CONTINUE
120  CONTINUE
      RETURN
      END

```

```

SUBROUTINE PASEU(LEV1)
  PARAMETER NUMCHN=6,NUMLV=40, NMOST=NUMLV+10
  DIMENSION KERNEL(NMOST,NUMCHN),GAMT(NMOST),RHOW(NMOST)
  REAL KERNEL
  COMMON/FRQ/FRGHZ(NUMCHN)
  COMMON/HANDU/H(NMOST),U(NMOST,NUMCHN)
  COMMON/PRES/P(NMOST)
  COMMON/TANDW/T(NMOST),W(NMOST)
  COMMON/TRANSM/TRANS(NMOST,NUMCHN)
  COMMON/INFUT/ALT(NMOST),TEMP(NMOST),PRE(NMOST),H2O(NMOST),
  * THINS(9),NTHICK,NBASE,TRN(NMOST,NUMCHN),NEWLEV(10)
  REAL NBASE
  DATA RSURV/4.615E6/RSURD/2.87E6/

C
C P = PRESSURE (MR)
C T = TEMPERATURE (DEG K.)
C W = MIXING RATIO (GM/KGM)
C H = HEIGHT (KM)
C RHOW = WATER VAPOR DENSITY (GM/M**3)
C KERNEL = WEIGHTING FUNCTION (TRANSMITTANCE/KM)
C GAMT = ATTENUATION IN LAYER I TO I+1 (DB/KM)
C FRGHZ = CHANNEL FREQUENCIES IN GHZ
C
      PRINT 16,(FRGHZ(I),I=1,NUMCHN) @ ECHO CHECK
      PRINT 17,LEV1,P(LEV1),P(1) @ ECHO CHECK
C
C*****
C COMPUTE WATER VAPOR DENSITY (RHOW) FROM PRES,TEMP AND MIXING RATIO.
C*****
      DO 30 I=1,LEV1
      RSUBH=(1./(1.+W(I)*1.E-3))*(W(I)*1.E-3*RSURV+RSURD)
      PCGS=P(I)*1.E3 @ DYNES/CM**2
      RHOM=PCGS/(RSUBH*T(I)) @ MOIST AIR DENSITY (GM/CM**3)
      RHOW(I)=RHOM*W(I)*1.E-3/(1.+W(I)*1.E-3) @ H2O VAPOR DENSITY (GM/CM**3)
      RHOW(I)=RHOW(I)*1.E6 @ H2O VAPOR DENSITY (GM/M**3)
30  CONTINUE

```

```

C
C*****
C COMPUTE ATTENUATION COEFFICIENTS
C*****
  PRINT 18
  DO 35 I=LEV1,1,-1
35 PRINT 19,P(I),H(I),T(I),RHOW(I)
  NUMLAY=LEV1-1      @ NUMBER OF LAYERS IN VERTICAL PROFILE
  DO 500 J=1,NUMCHN
    B=1.0              @ TRANSMITTANCE AT EARTHS SURFACE.
    A=0.0
    DO 215 I=2,LEV1
      TRANS(I-1,J)=B
      U(I-1,J)=A
      PTORR=P(I)*760./1013.25 @ CONVERT PRESSURE TO TORR (MM HG)
      GAM1=GAM02(FRGHZ(J),T(I),PTORR) @ ATTENU. AT PTORR,T(I) BY 02 (DB/KM)
      GAM2=GAMH20(FRGHZ(J),T(I),PTORR,RHOW(I)) @ ATTENU. AT PTORR,T(I),RHO(I)
      GAMT(I)=GAM1+GAM2 @ TOTAL ATTENUATION BETWEEN LEVEL I AND I+1 (DB/KM)
C FOR GROUND UP USE  DH=H(I+1)-H(I)
C FOR TOP DOWN USE  DH=H(I)-H(I+1)
      DH=H(I-1)-H(I) @ THICKNESS OF LAYER (KM)
      IF(H(I-1).EQ. H(I)) DH=H(I-1)-H(I+1) @FOR PRICIPITATING CASE
C*****
C LET ADB=GAMT(1)*DH (FLUX ATTENUATION IN DB)
C BUT DB=10LOG10(I(1)/I(2)) WHERE I(1).GT.I(2)
C THEREFOR ADB=GAMT(1)*DH=10LOG10(I(1)/I(2))
C ADB/10=LOG10(I(1)/I(2)), OR 10**(ADB/10)=I(1)/I(2)
C (ADB/10)LN(10)=LN(I(1)/I(2)), OF (LN(10)/10)*ADB=LN(I(1)/I(2))
C SO (.230259)*GAMT(1)*DH=LN(I(1)/I(2))=-LN(I(2)/I(1))=A
C SINCE TRANSMITTANCE, TNU2=I(2)/I(1)
C THEREFOR (.230259)*GAMT(1)*DH=-LN(TNU2)=A, AND EXP(-A)=TNU2
C
C SIMILARLY FOR LEVEL 3, ADB=GAMT(2)*DH
C AND (.230259)*GAMT(2)*DH=-LN(I(3)/I(2))
C AND (.230259)*GAMT(2)*DH+A=A YIELDS
C -LN(I(3)/I(2))-LN(I(2)/I(1))=-LN(I(3)/I(1))=A
C THEN -A=LN(I(3)/I(1))=LN(TNU3), THUS EXP(-A)=TNU3
C
C SIMILARLY FOR ANY LEVEL N, EXP(-A)=TNUN
C*****
      A=A+.230259*GAMT(I)*DH @ -LN(TNUN)
      WATEN=(B-EXP(-A))/DH @ (DELTA TNU)/(DELTA HT)
      KERNEL(I,J)=WATEN @ WEIGHTING FUNCTION
      B=EXP(-A) @ TRANSMITTANCE AT LOWER LVL OF NXT LAYER.
215 CONTINUE
      TRANS(LEV1,J)=B
      U(LEV1,J)=A
500 CONTINUE
1000 CONTINUE
16 FORMAT(1H0,'CHANNEL FREQS(GHZ) ',(10F7.2,2X))
17 FORMAT(1H0,'PROFILE CONTAINS ',I3,' LEVELS FROM ',F7.2,' MB TO',
  * F7.2,' MB')
18 FORMAT(1H0,'ECHO CHECK',T15,'PRES(MB)',T35,'HEIGHT(KM)',T55,
  * 'TEMP(DEG K)',T75,'H2O VAPOR DENSITY(GM/M**3)',/)
19 FORMAT(1X,T15,F8.3,T35,F7.3,T55,F6.2,T75,E9.4)
  RETURN
  END

```

```

      FUNCTION GAMU2(FREQ,T,P)
C*****
C COMPUTES OXYGEN ATTENUATION USING MEEKS AND LILLEY PARAMETERS
C*****
C
C FREQ = FREQUENCY (GHZ)
C NU = FREQUENCY (HZ)
C T = TEMPERATURE (DEGREES KELVIN)
C P = PRESSURE (TORR)
C
      REAL NU
      REAL NUPL(45),NUMI(45)
      IF (I1.EQ.10396) GO TO 2
C
C OXYGEN ABSORPTION LINES
C
      NUPL(1)=56.2648E9
      NUPL(3)=58.4466E9
      NUPL(5)=59.5910E9
      NUPL(7)=60.4348E9
      NUPL(9)=61.1506E9
      NUPL(11)=61.6002E9
      NUPL(13)=62.4112E9
      NUPL(15)=62.9980E9
      NUPL(17)=63.5685E9
      NUPL(19)=64.1272E9
      NUPL(21)=64.6779E9
      NUPL(23)=65.2240E9
      NUPL(25)=65.7626E9
      NUPL(27)=66.2978E9
      NUPL(29)=66.8313E9
      NUPL(31)=67.3627E9
      NUPL(33)=67.8923E9
      NUPL(35)=68.4205E9
      NUPL(37)=68.9478E9
      NUPL(39)=69.4741E9
      NUPL(41)=70.0000E9
      NUPL(43)=70.5244E9
      NUPL(45)=71.0497E9
      NUMI(1)=118.7505E9
      NUMI(3)=62.4863E9
      NUMI(5)=60.3061E9
      NUMI(7)=59.1642E9
      NUMI(9)=58.3239E9
      NUMI(11)=57.6125E9
      NUMI(13)=56.9682E9
      NUMI(15)=56.3634E9
      NUMI(17)=55.7839E9
      NUMI(19)=55.2214E9
      NUMI(21)=54.6728E9
      NUMI(23)=54.1294E9
      NUMI(25)=53.5960E9
      NUMI(27)=53.0695E9
      NUMI(29)=52.5458E9
      NUMI(31)=52.0259E9
      NUMI(33)=51.5091E9
      NUMI(35)=50.9949E9
      NUMI(37)=50.4830E9
      NUMI(39)=49.9730E9
      NUMI(41)=49.4648E9
      NUMI(43)=48.9582E9
      NUMI(45)=48.4530E9
      PM=267.41
      I1=10396
2 CONTINUE
      NU=1.E9*FREQ

```

```

C
C COMPUTE LINEWIDTH
C
      B=.25
      IF (P.LT.PM) B=.25*(1.+(ALOG(PM/P))/1.323)
      IF (P.LT.(18.957)) B=.75
      DNU1 = 1.95E+0*P *(0.21+0.78*B)*(300./T )**.85
      DNU12=DNU1**2
C
C GAMO2 = ATTENUATION COEFFICIENT (DB/KM)
C
      SUM=0.0
      DO 1103 N=1,45,2
      FN=N
      SUM=SUM
      1+((1./((NUPL(N)-NU)**2+DNU12)
      2+1./((NUPL(N) + NU)**2 + DNU12))
      3*(FN*(2.0*FN + 3.0)/(FN + 1.0))
      4+(1./((NUMI(N) - NU)**2 + DNU12)
      5+1./((NUMI(N) + NU)**2 + DNU12))
      6*(FN + 1.0)*(2.0*FN - 1.0)/FN
      7+1./((NU**2 + DNU12)
      8*2.0*(FN**2 + FN +1.0) *(2.0*FN +1.0)/(FN**2 + FN))
      9*EXP(-2.0684*FN*(FN+1.0)/T)
1103 CONTINUE
      GAMO2 =2.67*2E-9*P/T**3*NU**2*SUM*DNU1
      RETURN
      END

```

```

      FUNCTION GAMH20(FREQ,T,PTORR,RHO)
C
C NU = FREQUENCY (HZ)
C T = TEMPERATURE (DEGREES KELVIN)
C PTORR= PRESSURE (TORR)
C RHO = WATER VAPOR DENSITY (G/M3)
C GAMH20 = ATTENUATION COEFFICIENT (DB/KM)
C
      REAL NU,NUN
      DATA NUN/22.235E9/
      P=PTORR
      IF (FREQ.GT.60.) GO TO 300
C*****
C COMPUTES WATER VAPOR ATTENUATION USING BARRETT AND CHUNG FORMULA.
C*****
C
C BARRETT AND CHUNG LINELWIDTH.
C
      DNU2= (2.62E+9*P/7*NU.)/(T/318.])**.625*(1.0+.012 *RHO*T/P)
C
C COMPUTE WATER VAPOR ATTENUATION COEFFICIENT GAMH20.
C
      NU=FREQ*1.E9
      GAMH20 =4.56E-23*3.35E16*RHO *NU**2/T**1.5
      1*(EXP(-644./T)/T*(DNU2/((NU - NUN)**2 + DNU2**2)
      2 + DNU2/((NU + NUN)**2 + DNU2**2)) + 7.23E-24*DNU2)
      RETURN

```

AD-A087 434

UTAH UNIV. SALT LAKE CITY DEPARTMENT OF METEROLOGY
DEVELOPMENT OF THE MICROWAVE RADIATIVE TRANSFER PROG...ETC.(U)
SEP 79 LIQU, KUD-NAN F19628-78-C-0144

F/G 20/13

UNCLASSIFIED

PROJ. 7670

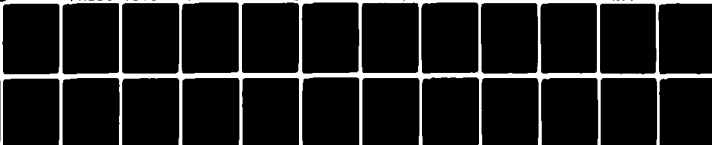
TASK 13

AFGL

TR-80-0051

N/I

2 of 2
24 2 13



END
DATE
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2-81
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```

C
  300 CONTINUE
C*****
C COMPUTES WATER VAPOR ABSORPTION FROM GAUT LINE MODEL FOR 183 GHZ
C LINE
C*****
  PMH=1013.25/760.*P/1000.
  CALL ALPHA5(FREQ,PMH,T,RHO,RES,RESNON)
  GAMM20=RES+RESNON
  RETURN
END

```

```

      SUBROUTINE ALPHA5 (FREQ,P,T, RHORAR,RES,RESNON)
C*****
C COMPUTES ATTENUATION FROM 183 GHZ WATER VAPOR LINE PLUS
C NONRESONANT BACKGROUND
C ORIGINAL PROGRAM BY N. GAUT
C*****
C
C FREQ = FREQUENCY (GHZ)
C P    = PRESSURE (MILLIBARS)
C T    = TEMPERATURE (DEGREES KELVIN)
C
  SORTF(X)=SORT(X)
  EXPF(X)=EXP(X)
  CONV=1000.
  FU93=323.756
  F164=183.310
  DELNSQ=(3.561*10.**(-7)*FREQ)**2*T /18.0
  DNU093=2.79*(P/1013.25)*(300./T)**(0.619)*(1.+0.0066*RHORAR)
  DNU093=SORTF(DELNSQ+LNU093**2)
  DNUWVW=1.04*0.280*RHORAR*(T/300.)**(0.1)
  DNUWVW=SORTF(DNUWVW**2+DELNSQ)
  DNUW=0.66*0.019*(300./T)**0.62
  DNUW=SORTF(DNUW**2+DELNSQ)
  DNU164=2.68*(P/1013.25)*(300./T)**(0.70)*(1.+0.0064*RHORAR)
  DNU164=SORTF(DNU164**2+DELNSQ)
  S1=1./((F164-FREQ)**2+DNU164**2)+1./((F164+FREQ)**2+DNU164**2)
  S2=1./((F093-FREQ)**2+DNU093**2)+1./((F093+FREQ)**2+DNU093**2)
  GAM1=0.978*RHORAR*FREQ**2/T**2.5*EXPF(-197.3/T)*DNU164*S1
  GAM2=0.849*RHORAR*FREQ**2/T**2.5*EXPF(-454./T)*DNU093*S2
  GAM3=2.55*FREQ**2/F164**3*(DNUWVW+DNUW)*(300./T)**(1.5)/1000.*RHO
  1HAR=4.165*T*0.760/1013.25
  RES=GAM1*CONV
  RESNON=(GAM2+GAM3)*CONV
  RETURN
END

```

```

      SUBROUTINE ANTEMP(TAU,TAUSQ,T,TSFC,EMIS,TA,ISFC,NU1,NU2)
      C*****
      C CALCULATE CLEAR COLUMN BRIGHTNESS VALUES FOR EACH CHANNEL (ANTENNA
      C TEMPERATURES).
      C*****

```

```

      PARAMETER NUMCHN=6,NUMLV=40, NMOST=NUMLV+10
      DIMENSION TAU(NMOST,NUMCHN),TAUSQ(NMOST,NUMCHN),T(NMOST),
      *      EMIS(NUMCHN),TA(NUMCHN)
      DIMENSION TAU1(NUMCHN),TAU2(NUMCHN),F(NUMCHN)
      B1=T(1)
      DO 3 J=NU1,NU2
      TAUS=TAU(ISFC,J)
      REFL=1.-EMIS(J)
      F(J)=TAUS*TAUS*REFL
      TAU1(J)=TAU(1,J)
3 TA(J)=0.
      DO 5 I=2,ISFC
      B2=T(I)
      DO 4 J=NU1,NU2
      TAU2(J)=TAU(I,J)
      TA(J)=TA(J)+0.5*(B1+B2)*(TAU1(J)-TAU2(J))*(1.+F(J)/TAUSQ(I,J))
4 TAU1(J)=TAU2(J)
5 B1=B2
      DO 6 J=NU1,NU2
      RS=TSFC*EMIS(J)
      TAUS=TAU(ISFC,J)
6 TA(J)=TA(J)+RS*TAUS
      RETURN
      END

```

```

      SUBROUTINE HUFF(P,T,h,H,U)
      C*****
      C INVERT PROFILES FOR COMPATIBILITY WITH REMAINING ROUTINES.
      C MAKE THE OPTION FOR CLEAR, CLOUDY, OR PRECIPITATING ATMOSPHERES.
      C*****

```

```

      PARAMETER LEV=40,NUMCHN=6, NMOST=LEV+10
      DIMENSION P(NMOST),T(NMOST),W(NMOST),H(NMOST),TRN(NMOST,NUMCHN)
      *      ,U(NMOST,NUMCHN),TAUABS(NMOST,NUMCHN)
      COMMON/INPUT/ALT(NMOST),TEMP(NMOST),PRE(NMOST),H2O(NMOST),
      *      THKNS(9),NTHICK,NBASE,TRN,NEWLEV(10),TAUABS
      COMMON/TAUVAL/TRANS(NMOST,NUMCHN)
      COMMON/SFCMS/EMIS(NUMCHN)
      COMMON/NEWLEV/ADDHT(10),NUMNEB,LVNUM
      J=LVNUM+1
      DO 100 I=1,LVNUM
      K=J-1
      PHE(K)=P(I)
      TEMP(K)=T(I)
      H2O(K)=W(I)
      ALT(K)=H(I)
      DO 50 L=1,NUMCHN
      TRN(K,L)=TRANS(I,L)
      TAUABS(K,L)=U(I,L)
50 CONTINUE
100 CONTINUE

```

```

      READ (5,200) ICHOSE
      GO TO (125,150,175,225),ICHOSE
125  RETURN
150  CALL DRIVER
      RETURN
175  CALL PFILE
      RETURN
225  CALL SIMPLR
      RETURN
200  FORMAT( )
      END

```

SUBROUTINE DRIVER

```

C *****
C THIS IS DRIVING ROUTINE FOR CALCULATION OF BRIGHTNESS TEMPERATURES
C WITH INTERVENING CLOUD LAYER.
C EXECUTIVE PROGRAM FOR CLOUD AND PRECIPITATION--A SOLUTION OF THE
C RADIATIVE TRANSFER EQUATION FOR A PLANE-PARALLEL CLOUDY ATMOSPHERE
C USING THE DISCRETE-ORDINATE METHOD. THIS PROGRAM WILL CALCULATE THE
C RADIANCE (INTENSITY) ABOVE AND BELOW NON-ISOTHERMAL CLOUDS AT THE
C WAVELENGTHS OF SCAMS OF NIMBUS 6. CLOUD TEMPERATURE GRADIENTS
C IN THE VERTICAL ARE APPROXIMATED BY DIVIDING THE CLOUD DEPTH INTO
C 'N' LAYERS, EACH LAYER ASSIGNED A RESPECTIVE TEMPERATURE ACCORDING
C TO THE INPUT ATMOSPHERIC PROFILE. A SYSTEM OF EQUATIONS FOR THE
C RADIANCE IS DERIVED FOR THE 16 DISCRETE RAYS.
C *****
C
C INPUT PARAMETERS. (UNITS)
C **ATMOSPHERIC CONDITIONS**
C ALT ALTITUDE OF ATMOSPHERIC LAYER ABOVE THE SURFACE KM
C TEMP TEMPERATURE CORRESPONDING TO 'ALT' DEG K
C PHE PRESSURE CORRESPONDING TO 'ALT' MILLIBARS
C H2O WATER VAPOR CONCENTRATION CORRESPONDING TO 'ALT' GM CM-3
C LEV NUMBER OF ATMOSPHERIC PROFILE LAYERS NONE
C
C **CLOUD PARAMETERS**
C NTHICK NUMBER OF CLOUD THICKNESS CASES PER RUN (10 MAX) NONE
C THICKS ARRAY OF CLOUD THICKNESS (MAX OF 9 VALUES) KM
C NBASE ALTITUDE OF CLOUD BASE (REAL NUMBER) KM
C NLYM NUMBER OF LAYERS WITHIN THE CLOUD (DEFAULT=3) NONE
C
C **EXTINCTION PROPERTIES**
C PINI LEGENDRE POLYNOMIAL EXPANSION OF PHASE FUNCTION NONE
C PT SINGLE SCATTERING ALBEDO ACCORDING TO 'NU' NONE
C NU WAVE NUMBER (LIMITED TO 'INT'=7) CM-1
C BEXT EXTINCTION COEFFICIENTS ACCORDING TO 'NU' NONE
C LV LINE CONTRIBUTION TO ABSORPTION COEFFICIENT NONE
C CK1 CONTINUUM CONTRIBUTION TO ABSORPTION COEFF TAKEN NONE
C CK2 FROM PAPER BY K J BIGNELL (1970) NONE
C

```

```

REAL NU, LV, NBASE
DOUBLE PRECISION PINO, UM, A, Z
PARAMETER LEV=40, INT=6, MZ=25, NG=16, NG2=8, NMOST=LEV+10
COMMON/INPUT/ALT(NMOST), TEMP(NMOST), PRE(NMOST), H2O(NMOST),
* THKNS(9), NTHICK, NBASE, TRN(NMOST, INT), NWLEV(10),
* TAUABS(NMOST, INT)
COMMON /ANGLE/ UM(NG), A(NG), PI
COMMON/LGMULT/UBC, ATMD, ATMR, ATMA, SFC, THRSV(INT), THRUPT(INT), ITER
* , ECB, IPCHAN(INT), CLDEM
COMMON /POLY/ PINO(NG)
COMMON /WINDOW/ UP(NG2, INT), DW(NG2, INT), THETA(NG), NU(INT), LV(INT),
ICK1(INT), CK2(INT), LUVERT(INT), CDVERT(INT)
COMMON /KADATA/ PINI(NG, INT), PT(INT), BEXT(INT)
COMMON/NEWLEV/ADDH1(10), NUMNEW, LVNUM
COMMON /FREQ/FGHZ(INT)
COMMON Z(1300)
COMMON/FLAGS/IFLAG(2)
DIMENSION NE=LV(10)
DATA SPOLGT/2.997929E10/
9 FORMAT( )
10 READ 9, LAST
IF (LAST, NE, 0) GO TO 100
C ***** INPUT CLOUD PARAMETERS *****
READ 9, (IPCHAN(I), I=1, INT)
READ 9, UBC, ATMD, ATMR, ATMA, SFC, ECB, CLDEM, ITER
READ 9, NTHICK, NBASE, NLYR
NTH=INT(NTHICK)
READ 9, (THKNS(I), I=1, NTH)
READ 9, (PT(I), I=1, INT)
READ 9, (BEXT(I), I=1, INT)
DO 12 J=1, INT
12 READ (5, 1900) (PINI(I, J), I=1, 16)
1900 FORMAT(5F16.8)
READ(5, 9) (IFLAG(J), J=1, 2)
C
C IFLAG(1)= 0 DO NOT INCLUDE SFC REFLECTION OF BRIGHTNESS CONTRIBUTION
C FROM ABOVE SCATTERING LAYER WITH LOWER BOUNDARY VALUE.
C
C 1 INCLUDE SFC REFLECTION OF BRIGHTNESS CONTRIBUTION
C FROM ABOVE SCATTERING LAYER WITH LOWER BOUNDARY VALUE.
C ( ASSUME TOP/DOWN THROUGHPUT EQUALS 1.)
C
C IFLAG(2)= 0 DO NOT WRITE SCATTERING PARAMETERS TO FILE 20.
C
C .NE.0 WRITE SCATTERING PARAMETERS TO FILE 20.
C (IFLAG(2)=RAINFALL RATE IN MM/HR 100 YIELDS RFR=0
C
DO 15 I=1, INT
15 NU(I)=(FGHZ(1)*1.E9)/SPOLGT @ CHANNEL WAVE NUMBERS.
K=1
KK=1
TEST=NBASE
DO 50 I=1, LVNUM
IF (ABS(ALT(I)-ADDH1(KK))-1.E-6) 35, 35, 38
35 NEWLV(KK)=I
KK=KK+1
38 IF (ABS(ALT(I)-TEST)-1.E-6) 40, 40, 50
40 NWLEV(K)=I
K=K+1
TEST=NBASE+THKNS(K-1)
IF (KK.GT.NUMNEW) GO TO 60
50 CONTINUE
60 PRINT 2000

```

```

2000 FORMAT(1H0,'THE LEVELS ADDED TO THE PROFILE ARE - ',/,5X,
* 'PRESS(MB)',/10X,'HEIGHT(KM)',/10X,'TEMP(DEG K)',/10X,
* 'MIX RAT(GM/KGM)')
DO 81 I=1,NUMNEW
J=NEWLV(I)
PRINT 2001,PHE(J),ALT(J),TEMP(J),H2O(J)
2001 FORMAT(5X,F9.2,10X,F10.2,10X,F11.2,10X,E15.5)
81 CONTINUE
KSTOP=0

```

```

C
C CHECK FOR LAST DATA CASE
C DETERMINE EXTINCTION COEFFICIENTS ONCE PER DATA CASE USING 'KSTOP'
C IN SUBSEQUENT ROUTINE.
C EXECUTE ENTIRE CODE ONCE FOR EACH INPUT CLOUD THICKNESS (THKNS)
C
DO 90 KTHK=1,NTHICK
85 CALL BNDHY (KSTOP,KTHK,NLYR)
CALL RAD (KTHK,NLYR,KSTOP,S85)
90 CONTINUE
100 CONTINUE
RETURN
END

```

```

SUBROUTINE BNDRY(KSTOP,KTHK,NLYR)
C*****
C CALCULATE UPWARD AND DOWNWARD BRIGHTNESS TEMPERATURES AT THE CLOUD
C BOTTOM AND TOP (BOUNDARY CONDITIONS)
C*****
C
  REAL LV,NU
  PARAMETER LEV=40, INT=6, MZ=25, N6=16, N62=8, NMOST=LEV+10
  DOUBLE PRECISION UM,A,CCRAD
  REAL NBASE
  COMMON/LBMULT/UBC,ATMO,ATMR,ATMA,SFC,THRUSV(INT),THRUPT(INT),INTER
  * ,EC0,IPCHAN(INT),CLOEM
  COMMON/INPUT/ALT(NMOST),TEMP(NMOST),PRE(NMOST),H2O(NMOST),
  * THKNS(9),NTHICK,NBASE,TRN(NMOST,INT),NWLEV(10)
  * ,TAUABS(NMOST,INT)
  COMMON /ANGLE/ UM(N6),A(N6),PI
  COMMON /WINDOW/ UP(N62,INT),DW(N62,INT),THETA(N6),NU(INT),LV(INT),
  * CK1(INT),CK2(INT),CUVERT(INT),CDVERT(INT)
  COMMON/NWLEV/ADUHT(10),NUMNEW,LVNUM
  COMMON /SFCEMS/ EMIS,NB
  COMMON /NADATA/ PINI(N6,INT),PT(INT),BEXT(INT)
  COMMON/FLAGS/IFLAG(2)
  DIMENSION TAUSQ(NMOST,INT)
  DIMENSION TAUM(NMOST,INT),TAU(NMOST,INT)
  DIMENSION TAUMOD(NMOST,INT),EMIS(INT),EUP(INT),EDW(INT)
  DIMENSION CUHORZ(INT),CDHORZ(INT),CCRAD(INT,N6)
  PI=3.1415926536
  KSTOP=KSTOP+1
  IF (KSTOP.GT.1) GO TO 40
  WRITE (6,160) NBASE,THKNS(KTHK)
  WRITE (6,180)
  DO 30 I=1,INT
    IF (IPCHAN(I).EQ.0) GO TO 30
    WRITE (6,190) (NU(I),PT(I),BEXT(I),(PINI(J,I),J=1,N6))
  30 CONTINUE
  40 CONTINUE
C
C DETERMINE GAUSSIAN QUADRATURE WEIGHTS FOR DISCRETE ORDINATE METHOD
C N6=NUMBER OF GAUSS POINTS (16)
C UM=COSINE(THETA)
C A=QUADRATURE WEIGHT VALUES
C
  CALL GAUSS (N6,-1.000,1.000)
C
C LOOP FOR ANGLES
C
  KOUNT=2
  DO 102 J=1,N62
    JJ=J+N62
    UM0=SNGL(UM(J))
    UM(JJ)=-UM(J)
    A(JJ)=A(J)
    UM1=SNGL(UM(J))
    UM2=SNGL(UM(JJ))
    THETA(J)=ACOS(UM1)*180./PI
    THETA(JJ)=ACOS(UM2)*180./PI
    CALL TRANMW(TAUM,TAUSQ,0.,1,INT,KOUNT,UM0)
    JJ=LVNUM+1
    DO 70 K=1,LVNUM
      L=JJ-K
      DO 60 M=1,INT
        TAU(L,M)=TAUM(K,M)
      60 CONTINUE
      NB=NWLEV(1)
      NTOP=NWLEV(KTHK+1)
    70 CONTINUE
  102 CONTINUE

```

```

C
C LOOP FOR CHANNELS
C
  DO 80 I=1,INT
    THRUPT(I)=1.0
    IF(IPCHAN(I).EQ.0) GO TO 80
    DO 75 M=NTOP,LVNUM
      IF(TAU(M,I).NE.0.) GO TO 74
      TAUMOD(M,I)=0.
    GO TO 75
    74 TAUMOD(M,I)=TAU(NTOP,I)/TAU(M,I)
    75 CONTINUE
    80 CONTINUE
    DO 101 I=1,INT
      IF(IPCHAN(I).EQ.0) GO TO 101
C
C COMPUTE DOWNWARD INTENSITY AT THE CLOUD TOP
C
    EDW(I)=0.
    NTOP1=NTOP+1
    DO 95 M=NTOP1,LVNUM
      TBAR=(TEMP(M-1)+TEMP(M))*0.5
      EDW(I)=EDW(I)+TBAR*(TAUMOD(M-1,I)-TAUMOD(M,I))
    95 CONTINUE
    DW(J,I)=EDW(I)*UBC
    CCRAU(I,J)=DBLE(DW(J,I))
    DO 97 K=1,LVNUM
      TAUMOD(K,I)=TAU(1,I)/TAU(K,I)
    97 CONTINUE
C
C COMPUTE UPWARD INTENSITY AT BOTTOM OF CLOUD
C
    BSFC=0.
    DATM=0.
    RATMB=0.
    RATMU=0.
    REFL=1.-EMIS(I)
    IF(NB.EQ.1) GO TO 98
    DO 100 K=2,NB
      TBAR=(TEMP(K)+TEMP(K-1))*0.5
      IF(TAU(NB,I).NE.0.) GO TO 90
      DATM=DATM
    GO TO 96
    90 DATM=DATM+TBAR*(TAU(K,I)-TAU(K-1,I))/TAU(NB,I)*ATMD
    96 RATMB=RATMB+TBAR*(TAUMOD(K-1,I)-TAUMOD(K,I))*REFL*TAUMOD(NB,I)
    * ATMD
    100 CONTINUE
    98 BSFC=TEMP(1)*EMIS(I)*TAUMOD(NB,I)*SFC
    RATMU=EDW(I)*TAUMOD(NB,I)*REFL*TAUMOD(NB,I)*ATMA*THRUPT(I)
    IF(ATMA.EQ.0)
      * RATMU=CDVENT(I)*TAUMOD(NB,I)*REFL*TAUMOD(NB,I)*ECB
    IF(IFLAG(1).EQ.0) RATMU=0.
    EUP(I)=DATM+RATMB+RATMU
    UP(J,I)=EUP(I)+BSFC
    CCRAU(I,J)=DBLE(UP(J,I))
    101 CONTINUE
    102 CONTINUE
C
C EXTRAPOLATE INTENSITIES TO THETA=0, AND 90 DEG.
C
    ICCRAD=1
    IEMIS=0
    DO 110 JAY=1,INT
      IF(IPCHAN(JAY).EQ.0) GO TO 110
      JOUM=INT
      IUPJ=0

```

```

CALL EXTRP0(THETA,CCRAD,TVERT,JAY,JJUM,IUPDOWN)
CALL EXTRP9(THETA,CCRAD,THORIZ,JAY,JJUM,IUPDOWN)
CUVERT(JAY)=TVERT
CUHORZ(JAY)=THORIZ
IUPDOWN=1
CALL EXTRP0(THETA,CCRAD,RVERT,JAY,JJUM,IUPDOWN)
CALL EXTRP9(THETA,CCRAD,RHORIZ,JAY,JJUM,IUPDOWN)
CUVERT(JAY)=RVERT
CUHORZ(JAY)=RHORIZ
110 CONTINUE
WRITE (6,200) (THKNS(KTHK),(THETA(J),J=1,N62))
DO 120 I=1,INT
IF(IPCHAN(I).EQ.0) GO TO 120
WRITE (6,210) (NU(I),CUVERT(I),(UP(J,I),J=1,N62),CUHORZ(I))
120 CONTINUE
N621=N62+1
WRITE (6,220) (THETA(J),J=N621,N6)
DO 130 I=1,INT
IF (IPCHAN(I).EQ.0) GO TO 130
WRITE (6,210) (NU(I),CUVERT(I),(DW(J,I),J=1,N62),CDHORZ(I))
130 CONTINUE
RETURN
160 FORMAT (//,54X,'CLOUD STRUCTURE',//,22X,'CLOUD TEMPERATURES = (DERI
IVED FROM THE ATMOS. TEMP. PROFILE + CLOUD THICKNESS)',//,22X,'CLOUD
2 BASE HEIGHT =',F5.2,'KM',//,28X,'CLD THKNS =',F4.1,'KM (MAXIMU
3M OF 10 VALUES)',//,35X,'INPUT ABSORPTION COEFF ',
4//,44X,'NU',9X,'LV',8X,'CK1',6X,'CK2')
180 FORMAT (/,37X,'SINGLE SCATTERING PROPERTIES OF WATER DROPLETS',//,
111X,'NU',9X,'PT',7X,'BEXT',5X,'.....P
2INI.....')
190 FORMAT (9X,F6.4,2(4X,F6.3),4X,8(F8.5,1X)//,39X,8(F8.5,1X))
200 FORMAT(1H1,37X,' BRIGHTNESS TEMPERATURE AT THE CLOUD TOP AND ',
* 'BASE',//,48X,'CLOUD THICKNESS ',F5.3,' KM',//,49X,'THE UPWARD ',
* 'BRIGHTNESS',//,4X,'(THETA) (0,0000)',3X,8('(',F7.4,')',3X),
* '(90,0000)',//,5X,'WAVE NO')
210 FORMAT(5X,F6.4,F10.3,1X,9(F10.3,2X),F10.3)
220 FORMAT (/,48X,'THE DOWNWARD BRIGHTNESS',//,4X,'(THETA) (180,000)',
*3X,8('(',F7.3,')',3X),'( 90,000)',//,5X,'WAVE NO')
END

```

```

      SUBROUTINE GAUSS(N,XL,XU)
C *****
C      SUBROUTINE 'GAUSS' DETERMINES THE GAUSSIAN QUADRATURE VARIABLES
C      ('R'='UM' = COSINE(THETA), 'W'='A' = WEIGHTING FACTOR) FOR THE
C      DISCRETE ORDINATE METHOD.
C *****
C
      IMPLICIT DOUBLE PRECISION (A-H,O-Z)
      COMMON /ANGLE/ R(16),W(16),PI
      DIMENSION Z(50),PA(50)
      TOL=1.0D-16
      PI=3.141592653589793D+00
      AA=2.0D+00/PI**2
      AB=-62.0D+00/(3.0D+00*PI**4)
      AC=15116.0D+00/(15.0D+00*PI**6)
      AD=-12554474.0D+00/(105.0D+00*PI**8)
      PA(1)=1.0D+00
      EN=N
      NP1=N+1
      U=1.0D+00-(2.0D+00/PI)**2
      D=1.0D+00/DSQRT((LN+0.5D+00)**2+U/4.0D+00)
      DO 10 I=1,N
      SM=1
      AZ=4.0D+00*SM-1.0D+00
      AE=AA/AZ
      AF=AB/AZ**3
      AG=AC/AZ**5
      AH=AD/AZ**7
10  Z(I)=0.25D+00*PI*(AZ+AE+AF+AG+AH)
      DO 20 K=1,N
      X=DCOS(Z(K)*U)
20  PA(K)=X
      DO 30 NN=3,NP1
      ENN=NN-1
30  PA(NN)=((2.0D+00*ENN-1.0D+00)*X*PA(NN-1)-(ENN-1.0D+00)*PA(NN-2))/
      1ENN
      PNP=EN*(PA(N)-X*PA(NP1))/(1.0D+00-X*X)
      XI=X-PA(NP1)/PNP
      XU=DAHS(XI-X)
      XDD=XD-TOL
      IF (XDD) 50,50,40
40  X=XI
      GO TO 20
50  R(K)=X
60  W(K)=2.0D+00*(1.0D+00-X*X)/(EN*PA(N))**2
      RETURN
      END

```

```

      SUBROUTINE EXTRP0(THETA,INTENS,TVERT,JAY,JDUM,IUPDWN)
C*****
C THIS SUBROUTINE IS TO EXTRAPOLATE THE INTENSITIES TO THETA=0, AND 90
C BY APPLYING LEAST-SQUARE POLYNOMIAL.
C*****
C
      IMPLICIT DOUBLE PRECISION (A-H,O-Z)
      REAL THETA,TVERT
      DOUBLE PRECISION INTENS
      DIMENSION THETA(16),INTENS(JDUM,16)
      K1=1+A*IUPDWN
      XX=FLOAT(IUPDWN*180)
5    K2=K1+1
      K3=K1+2
      T1=THETA(K1)
      T2=THETA(K2)
      T3=THETA(K3)
      R1=INTENS(JAY,K1)
      R2=INTENS(JAY,K2)
      R3=INTENS(JAY,K3)
      C3=(R2-R1)/((T2-T1)*(T2-T3))+ (R1-R3)/((T3-T1)*(T2-T3))
      C2=((R3-R1)-C3*(T3-T1)*(T3+T1))/(T3-T1)
      C1=R1-C2*T1-C3*T1*T1
      TVEN1=C1+C2*XX+C3*XX*XX
      GO TO 50
C
C THE FOLLOWING ENTRY IS FOR THETA=90.
C
      ENTRY EXTRP9(THETA,INTENS,TVERT,JAY,JDUM,IUPDWN)
      XX=90.
      K1=6+A*IUPDWN
      GO TO 5
50 CONTINUE
      RETURN
      END

```

```

      SUBROUTINE RAD (KTHK,NLYR,KSTOP,S)
C *****
C FOR A GIVEN CLOUD DEPTH ('KTHK') THIS ROUTINE APPLIES THE DISCRETE
C ORDINATE METHOD TO THE RADIATIVE TRANSFER EQUATION FOR EACH OF THE
C SPECIFIED WAVE NUMBERS ('NU').
C *****
C
      PARAMETER LEV=40, INT=6, MZ=25, NG=16, NG2=8, NMOST=LEV+10
      DOUBLE PRECISION PINO,UM,A
      COMMON /ANGLE/ UM(NG),A(NG),PI
      COMMON/LBMULT/UBC,ATMO,ATMR,ATMA,SFC,THRUSV(INT),THRUPT(INT),ITER
      * ,ECB,IPCHAN(INT)
      COMMON /POLY/ PINO(NG)
      COMMON /HADATA/ PINI(NG,INT),PT(INT),BEXT(INT)
C
C  LOOP THRU MAIN CALCULATIONS FOR EACH WAVE NUMBER (INDEX='JAY')
C
      DO 30 JAY=1,INT
        IF(IPCHAN(JAY).EQ.0) GO TO 30
        DO 10 L=1,NG
          PINO(L)=UBLE((PINI(L,JAY))*PT(JAY))
        10 CONTINUE
        IF (KSTOP.GT.1) GO TO 20
        CALL IGEN (JAY)
        20 CALL TGRAD (JAY,KTHK,NLYR)
        30 CONTINUE
        DO 40 I=1,INT
          IF(IPCHAN(I).EQ.0) GO TO 40
          IF(ABS(THRUSV(I)-THRUPT(I)).GT.0.1) RETURN 4
        40 CONTINUE
        RETURN
      END

```

```

      SUBROUTINE IGEN (JAY)
C *****
C SUBROUTINES 'IGEN', 'SIANT', 'FILLH', 'FILLC', 'FILLD', 'FILLS', 'GMPRD',
C 'LEP', 'UPRHM', AND 'UPQFB' DETERMINE THE EIGENVALUES RT(I)
C ACCORDING TO THE WAVE NUMBER (LIOU, 1973, APPENDIX)
C *****
C
      PARAMETER NG=16, NG2=8, INT=6
      DOUBLE PRECISION PINO,Z,VALUES
      COMMON Z(1300)
      COMMON /EIGEN/ VALUES(INT,NG2)
      COMMON /POLY/ PINO(NG)
      NDIM=1300
      NX=NG
      N=NG
      MDIM=3*N*(N+1)
      IF (MDIM.GT.NDIM) GO TO 10
      ND2=(NDIM-N)/2
      NN=N**2
      NS=3*NN+1
      CALL START (Z(1),Z(NN+1),Z(2*NN+1),PINO,Z(1),Z(NS+2*N),N,7(2*N+1),
      1Z(3*N+1),JAY)
      GO TO 20
10  WRITE (6,30) N,NDIM,MDIM
      PRINT 40
20  CONTINUE
      RETURN
C
30  FORMAT (J15)
40  FORMAT ('PROGRAM MUST BE REDEMENSIONED AT MDIM=')
      END

```

```

      SUBROUTINE START (B,C,P,T,D,S,N,COMPLE,POL,JAY)
C*****
C SUBROUTINE 'START' IS TO COMPUTE EIGENVALUES 'VALUES(JAY,I)' AND
C RETURNS THE EIGENVALUES THROUGH 'IGEN' TO 'TGRAG' WHERE THE
C RADIANCE CALCULATIONS ARE MADE.
C*****
C
      PARAMETER NG2=8, NG=16, INT=6
      DOUBLE PRECISION UM,A,P,C,B,S,D,T,DMU,COMPLE,POL,G,SGROOT,PROOT,
      1VALUE,VALUES
      COMMON /ANGLE/ UM(NG),A(NG),PI
      COMMON /EIGEN/ VALUES(INT,NG2)
      DIMENSION B(N,N),C(N,N),P(N,N),T(N),D(N),S(N)
      DIMENSION COMPLE(I),POL(N),G(NG),DMU(30),VALUE(NG2)
C
      NL1=N-1
      DO 10 I=1,N
10    CALL LEP (P(1,I),UM(I),NL1)
      DO 20 I=1,N
      DO 20 J=1,N
      CALL FILLC (T,P(1,I),P(1,J),A(J),C(1,J),N)
20    CONTINUE
      CALL FILLH (H,C,UM,N,NL1)
      CALL FILLR (S,R,C,P,N)
      CALL FILLD (D,S,N)
      IF (ABS(T(1))-1.0).LE.1.E-10) D(N)=0.
      DO 30 I=1,N
      DMU(I)=1.00/UM(I)
30    CONTINUE
      N2=N/2+1
      IH=N+2
      G(N2)=1.00
      DO 40 I=2,N/2
      J=(IH-1)/2
      G(J)=D(I)
40    CONTINUE
      CALL DPHRM (G,N2,VALUE,COMPLE,POL,IR,IER)
      DO 60 I=1,IR
      SGROOT=USQRT(DABS(VALUE(I)))
      PROOT=SGROOT
      DO 50 J=1,NL1
50    PROOT=(PROOT+D(J))*SGROOT
      PROOT=PROOT+D(N)
      VALUE(I)=SGROOT
      IF (PROOT.GT.1.00-02) WRITE (6,80) PROOT
60    CONTINUE
      DO 70 I=1,NG2
      VALUES(JAY,I)=VALUE(I)
70    CONTINUE
      RETURN
C
80    FORMAT ('0//////// EIGENVALUE ERROR =',D18.8/)
      END

```

```

      SUBROUTINE LEP (Y,X,N)
C
C ***** LEGENDRE POLYNOMIAL EXPANSION *****
C
      DOUBLE PRECISION Y,X,G
      DIMENSION Y(1)
      Y(1)=1.
      IF (N) 10,10,20
10  RETURN
20  Y(2)=X
      IF (N-1) 10,10,30
30  DO 40 I=2,N
      G=X*Y(I)
40  Y(I+1)=G-Y(I-1)+G-(G-Y(I-1))/FLOAT(I+1)
      RETURN
      END

```

```

      SUBROUTINE FILLC (T,PI,PJ,AJ,C,N)
C
C COMPUTE COEFFICIENT C(I,J)
C SEE L100, 1973 (APPENDIX)
C
      DOUBLE PRECISION T,PI,PJ,AJ,C
      DIMENSION T(1),PI(1),PJ(1)
      C=0.
      DO 10 L=1,N
10  C=C+T(L)*PI(L)*PJ(L)
      C=C*AJ/2.
      RETURN
      END

```

```

      SUBROUTINE FILLH (B,C,UM,N,NL1)
C
C SEE L100, 1973 (APPENDIX)
C
      DOUBLE PRECISION B,C,UM
      DIMENSION B(N,N),C(N,N),UM(N)
      DO 20 I=1,NL1
        IP1=I+1
        DO 10 J=IP1,N
          H(I,J)=C(I,J)/UM(I)
10      H(J,I)=C(J,I)/UM(J)
20      H(I,I)=(C(I,I)-1.00)/UM(I)
          H(N,N)=(C(N,N)-1.00)/UM(N)
      RETURN
      END

```

```

      SUBROUTINE FILLS (S,B,C,P,N)
C
C S(N) IS TRACE OF R(N,N)
C SEE L100, 1973 (APPENDIX)
C
      DOUBLE PRECISION S,B,P,C
      DIMENSION S(1),H(N,N),P(N,N),C(N,N)
      DO 10 I=1,N
        S(I)=0.
        S(I)=S(I)+B(I,I)
        DO 10 J=1,N
10      C(I,J)=B(I,J)
          NH=N
          IF (MOD(N,2).EQ.0) NH=N-1
          IF (N.LE.2) GO TO 50
          DO 40 I=2,NH,2
            CALL GMPKD (H,C,P,N,N,N)
            DO 20 J=1,N
20          S(I)=S(I)+P(J,J)
            CALL GMPKD (B,P,C,N,N,N)
            IP1=I+1
            DO 30 J=1,N
30          S(IP1)=S(IP1)+C(J,J)
40          CONTINUE
50          IF (NH.EQ.N) RETURN
            CALL GMPKD (B,C,P,N,N,N)
            DO 60 J=1,N
60          S(N)=S(N)+P(J,J)
      RETURN
      END

```

```

      SUBROUTINE GMPRO (A,B,R,N,M,L)
C
C THE SUBROUTINE IS TO COMPUTE GENERAL MATRIX PRODUCT.
C
      DOUBLE PRECISION K,A,B
      DIMENSION A(1),R(1),R(1)
      IR=0
      IK=-M
      DO 10 K=1,L
      IK=IK+M
      DO 10 J=1,N
      IR=IR+1
      JI=J-N
      IH=IK
      R(IH)=0
      DO 10 I=1,M
      JI=JI+N
      IR=IR+1
10  R(IK)=R(IR)+A(JI)*B(IB)
      RETURN
      END

```

```

      SUBROUTINE FILLD (D,S,N)
C
C D(I) ARE COEFFICIENTS OF THE EXPANDED POLYNOMIAL OF EIGENVALUES.
C SEE LIOU, 1973 (APPENDIX)
C
      DOUBLE PRECISION D,S,X
      DIMENSION S(1),D(1)
      D(1)=-S(1)
      DO 20 I=2,N
      IL1=I-1
      X=0.
      DO 10 J=1,IL1
10  X=X+D(1-J)*S(J)
20  D(I)=-X+S(1)/FLOAT(I)
      RETURN
      END

```

```

      SUBROUTINE DPKRM (C,IC,RR,RC,POL,IR,IER)
C*****
C SUBROUTINES 'DPKRM' AND 'DPQFB' ARE FOR THE CALCULATION OF ROOTS
C (EIGENVALUES) ACCORDING TO POLYNOMIAL EXPANSIONS
C*****
      DIMENSION C(1),RR(1),RC(1),POL(1),Q(4)
      DOUBLE PRECISION C,RR,RC,POL,Q,EPS,A,R,H,Q1,Q2
C
C TEST ON LEADING ZERO COEFFICIENTS
C
      EPS=1.0D-6
      LIM=100
      IR=1C+1
10  IR=IR-1
      IF (IR-1) 420,420,20
      20 IF (C(IR)) 30,10,30
C
C WORK UP ZERO ROOTS AND NORMALIZE REMAINING POLYNOMIAL
C
      30 IER=0
      J=IR
      L=0
      A=C(IR)
      DO 40 I=1,IR
      IF (L) 40,40,70
      40 IF (C(I)) 60,50,60
      50 RR(1)=0.00
      RC(I)=0.00
      POL(J)=0.00
      J=J-1
      GO TO 80
      60 L=1
      IST=I
      J=0
      70 J=J+1
      C(I)=C(I)/A
      POL(J)=C(I)
      CALL OVERFL (H)
      IF (H=2) 420,80,80
      80 CONTINUE
C
C START RAIRSTON ITERATION
C
      Q1=0.00
      Q2=0.00
      90 IF (J-2) 330,100,140
C DEGREE OF POLYNOMIAL IS EQUAL TO ONE
      100 A=POL(1)
      RR(1ST)=-A
      RC(1ST)=0.00
      IR=1K-1
      Q2=0.00
      IF (IR-1) 130,130,110
      110 DO 120 I=2,IR
      Q1=Q2
      Q2=POL(I+1)
      120 POL(I)=A+Q2+Q1
      130 POL(IR+1)=A+Q2
      GO TO 340
C DEGREE OF POLYNOMIAL IS GREATER THAN ONE
      140 DO 220 L=1,10
      N=L
      150 Q(1)=Q1
      Q(2)=Q2
      CALL DPQFB (POL,J,Q,LIM,I)
      IF (I) 160,240,230
      160 IF (Q1) 180,170,160
      170 IF (Q2) 180,210,160
      180 GO TO (190,200,190,210), N
      190 Q1=-Q1
      N=N+1
      GO TO 150

```

```

200 Q2=-Q2
    N=N+1
    GO TO 150
210 Q1=1.00+Q1
220 Q2=1.00-Q2
    IER=3
    Ik=Ik-J
    RETURN
230 IER=1
240 Q1=Q(1)
    Q2=Q(2)
    H=0.00
    A=0.00
    I=J
250 H=-Q1+R-Q2+A+POL(I)
    POL(I)=H
    B=A
    A=H
    I=I-1
    IF (I-2) 260,260,250
260 POL(2)=H
    POL(I)=A
    L=Ik-1
    IF (J-L) 270,270,290
270 DO 280 I=J,L
280 POL(I-1)=POL(I-1)+POL(I)*Q2+POL(I+1)*Q1
290 POL(L)=POL(L)+POL(L+1)*Q2+Q1
    POL(IR)=POL(IR)+Q2
    H=-.500*Q2
    A=H+Q1
    H=DSQRT(DABS(A))
    IF (A) 300,300,310
300 RR(IST)=H
    RC(IST)=B
    IST=IST+1
    RR(IST)=H
    RC(IST)=-B
    GO TO 320
310 B=H+DSIGN(B,H)
    RR(IST)=Q1/B
    RC(IST)=0.00
    IST=IST+1
    RR(IST)=H
    RC(IST)=0.00
320 IST=IST+1
    J=J-2
    GO TO 90
330 IR=IR-1
340 A=0.00
    DO 340 I=1,IR
    G1=C(I)
    Q2=POL(I+1)
    POL(I)=Q2
    IF (G1) 350,360,350
350 Q2=(G1-Q2)/G1
360 Q2=UARS(Q2)
    IF (Q2-A) 380,380,370
370 A=Q2
380 CONTINUE
    I=Ik+1
    POL(I)=1.00
    RR(I)=A
    RC(I)=0.00
    IF (IER) 390,390,410
390 IF (A-EPS) 410,410,400
400 IER=-1
410 RETURN
420 IER=2
    IR=0
    RETURN
END

```

```

SUBROUTINE DPMFH (C,IC,G,LIM,IER)
DIMENSION C(1),U(1)
DOUBLE PRECISION A,H,AA,RR,CA,CB,CC,CD,A1,B1,C1,H,HH,Q1,Q2,QQ1,
1QQ2,QQQ1,QQQ2,DQ1,DQ2,EPS,EPS1,C,Q
IER=0
J=IC+1
10 J=J-1
IF (J-1) 400,400,20
20 IF (C(J)) 30,10,30
30 A=C(J)
IF (A-1.00) 40,60,40
40 DO 50 I=1,J
C(I)=C(I)/A
CALL OVERFL (N)
IF (N-2) 400,50,50
50 CONTINUE
60 IF (J-3) 410,380,70
70 EPS=1.0-14
EPS1=1.0-6
L=0
LL=0
Q1=U(1)
Q2=U(2)
QQ1=0.00
QQ2=0.00
AA=C(1)
BH=C(2)
CH=UARS(AA)
CA=UARS(RR)
IF (CH-CA) 80,90,100
80 CC=CH+CR
CH=CH/CA
CA=1.00
GO TO 110
90 CC=CA+CA
CA=1.00
CR=1.00
GO TO 110
100 CC=CA+CA
CA=CA/CR
CB=1.00
110 CD=CC+.100
120 A=0.00
Q=A
A1=A
H1=H
I=J
QQQ1=Q1
QQQ2=Q2
DQ1=HH
DQ2=H
130 H=Q1+B-Q2+A+C(I)
CALL OVERFL (N)
IF (N-2) 420,140,140
140 B=A
A=H
I=I-1
IF (I-1) 180,150,160
150 H=0.00
160 H=Q1+B1-Q2+A1+H
CALL OVERFL (N)
IF (N-2) 420,170,170
170 C1=B1
B1=A1
A1=H
GO TO 130
180 H=CA+DABS(A)+CB+DABS(R)
IF (LL) 190,190,390
190 L=L+1

```

```

      IF (DABS(A)-EPS*DABS(C(1))) 200,200,210
200 IF (DABS(B)-EPS*DABS(C(2))) 390,390,210
210 IF (H-CC) 220,220,230
220 AA=A
      BH=B
      CC=C
      Q1=Q1
      Q2=Q2
230 IF (L-LIM) 280,280,240
240 IF (H-CD) 430,430,250
250 IF (G(1)) 270,260,270
260 IF (G(2)) 270,420,270
270 G(1)=0.00
      Q(2)=0.00
      GO TO 70
280 HH=UMAX1(DABS(A),DABS(B),DABS(C))
      IF (HH) 420,420,290
290 A1=A/HH
      B1=B/HH
      C1=C1/HH
      H=A1+C1-B1*B1
      IF (H) 300,420,300
300 A=A/HH
      B=B/HH
      HH=(H-A1-A*B1)/H
      H=(A+C1-B*B1)/H
      G1=G1+HH
      Q2=Q2+H
      IF (DABS(HH)-EPS*LAHS(Q1)) 310,310,330
310 IF (DABS(H)-EPS*DABS(Q2)) 320,320,330
320 LL=1
      GO TO 120
330 IF (L-1) 120,120,340
340 IF (DABS(HH)-EPS1*DABS(Q1)) 350,350,120
350 IF (DABS(H)-EPS1*LAHS(Q2)) 360,360,120
360 IF (DABS(Q1+HH)-DABS(Q1+Q1)) 370,440,440
370 IF (DABS(Q2+H)-DABS(Q2+Q2)) 120,440,440
380 Q(1)=C(1)
      Q(2)=C(2)
      Q(3)=0.00
      Q(4)=0.00
      RETURN
390 Q(1)=Q1
      Q(2)=Q2
      Q(3)=A
      Q(4)=B
      RETURN
400 IER=-1
      RETURN
410 IER=-2
      RETURN
420 IER=-3
      GO TO 440
430 IEN=1
440 G(1)=Q1
      G(2)=Q2
      G(3)=AA
      G(4)=BH
      RETURN
      END

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SUBROUTINE TGRAD (JAY,KTHK,NLYR)
C *****
C SUBROUTINE 'TGRAU' FORMS THE MATRIX OF LINEAR EQUATIONS RESULTING
C FROM THE LAYERED CLOUD STRUCTURE AND BOUNDARY CONDITIONS IMPOSED
C BY THE SURROUNDING NON-SCATTERING MOLECULAR ATMOSPHERE IN LOCAL
C THERMODYNAMIC EQUILIBRIUM. THE PRINCIPLE CALCULATIONS ARE
C PERFORMED IN THIS ROUTINE WITH THE RESULTANT OUTPUT OF THE SPATIAL
C DISTRIBUTION OF RADIANT INTENSITY ACCORDING TO THE DISCRETE RAYS
C ('NG2') ABOVE AND BELOW THE CLOUD. VALUES FOR THE ZENITH ANGLES
C 0 AND 90 DEG FOR THE TRANSMITTED AND REFLECTED RADIANCE ARE
C EXTRAPOLATED FROM A LEAST-SQUARES POLY. FIT TO THE CALCULATIONS.
C *****
C
PARAMETER LEV=40, NT=6, MZ=25, NG=16, NG2=8, NGLM=48
* ,NMOST=LEV+10 ,INT2=2*INT
REAL NU, LV, NBASE
DOUBLE PRECISION AX, DIFF, DIFNB, ENTENS, FNOR, FUNC, INTENS, LE, LI, ORT,
OSADA, PINO, POLI, RT, SADA, TAU, W, W1, W2, XPL, XPR, YPL, YPR, UM, A,
ZVALUES
COMMON/INPUT/ALT(NMOST), TEMP(NMOST), PRE(NMOST), H2O(NMOST),
* THKNS(9), NTHICK, NBASE, TRN(NMOST,INT), NEWLEV(10)
* ,TAUABS(NMOST,INT)
COMMON /ANGLE/ UM(NG), A(NG), PI
COMMON/LBMULT/UBC,ATMD,ATMR,ATMA,SFC,THRUSV(INT),THRUPT(INT),ITER
* ,ECB,IPCHAN(INT),CLUEM
COMMON /POLY/ PINO(NG)
COMMON /BINDOW/ UP(NG2,INT),DW(NG2,INT),THETA(NG),NU(INT),LV(INT),
ICK1(INT),CK2(INT),CUVERT(INT),COVERT(INT)
COMMON /CLOU/ TEMPC(10)
COMMON /EIGEN/ VALUES(INT,NG2)
COMMON/SFCEMS/EMIS(INT),NB
COMMON/NEWLEV/ADDH(10),NUMNEW,LVNUM
COMMON/FLAGS/IFLAG(2)
COMMON /HADATA/ PINI(NG,INT),PT(INT),BEXT(INT)
DIMENSION YPL(NG),YPR(NG),INTENS(3,NG),POLI(NG,NG)
DIMENSION LE(NGLM),W1(NGLM,NGLM),W2(NGLM,NGLM),ACOSUM(NG)
DIMENSION SADA(NG,NG),OSADA(NG,NG),OPTH(9),W(NG,NG),LI(NGLM)
DIMENSION BINTC(3),TAU(3),ENTENS(3,NG),RT(NG),SAVE(INT2)
C
C DETERMINE EXTINCTION COEFFICIENTS
C
DO 10 I=1,NG2
II=I+NG2
RT(I)=VALUES(JAY,I)
RT(II)=-RT(I)
10 CONTINUE
C
C COMPUTE LEGENDRE POLYNOMIAL
C
DO 30 I=1,NG
DO 20 L=1,NG
LI=L-1
XPL=UM(I)
CALL LEP (YPL,XPL,LI)
POLI(I,L)=YPL(L)
20 CONTINUE
30 CONTINUE

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      DO 40 I=1,NG
      UMI=SNGL(UM(I))
      ACOSUM(I)=ACOS(UMI)
40    CONTINUE
C
C  COMPUTE SADA
C
      DO 60 K=1,NG
      DO 50 L=1,NG
      LI=L-1
      XXR=RT(K)
      CALL DSADA (YXR,XXK,LI)
      SADA(K,L)=YXR(L)
50    CONTINUE
60    CONTINUE
C
C  FIND RIGHT W FOR THE SMALLEST ROOT
C
      IF (NG.LT.10) GO TO 170
      DO 160 NG5=NG2,NG2
      ORT=RT(NG5)
      IS=1
      COUNT=0.0
70    CONTINUE
      COUNT=COUNT+1.0
      RT(NG5)=ORT
      DO 90 L=1,NG
      LI=L-1
      XXR=RT(NG5)
      CALL DSADA (YXR,XXK,LI)
      SADA(NG5,L)=YXR(L)
      IF (IS.EQ.1) GO TO 80
      SADA(NG5,L)=USADA(NG5,L)
80    CONTINUE
90    CONTINUE
      FUNC=0.000
      DO 110 I=1,NG
      W(I,NG5)=0.000
      DO 100 L=1,NG
      W(I,NG5)=W(I,NG5)+PINO(L)*POLI(I,L)*SADA(NG5,L)/(1.000+UM(I)*
1RT(NG5))
100   CONTINUE
      FUNC=FUNC+A(I)*W(I,NG5)/2.000
110   CONTINUE
      DO 120 L=1,NG
      OSADA(NG5,L)=0.000
      DO 120 I=1,NG
      OSADA(NG5,L)=OSADA(NG5,L)+A(I)*POLI(I,L)*W(I,NG5)/2.000
      FNOR=OSADA(NG5,1)
      DO 130 L=1,NG
      OSADA(NG5,L)=OSADA(NG5,L)/FNOR
      SADA(NG5,L)=OSADA(NG5,L)
130   CONTINUE
      IS=2
      ORT=(1.00-PINO(1))/OSADA(NG5,2)
      OIFNG=1.00-FUNC
      DIFF=DABS(1.00-FNOR)
      IF (DIFF.LE.1.0-7) GO TO 150
      IF (COUNT.GE.90.0) GO TO 140
      GO TO 70
140   WRITE (6,410) DIFF
      GO TO 400
150   CONTINUE

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DO 160 L=1,NG
SADA(NGS,L)=OSADA(NGS,L)
LL=L+1
KL=MOD(LL,2)
NGS2=NGS+2
IF (KL.EQ.0) SADA(NGS2,L)=SADA(NGS,L)
IF (KL.EQ.1) SADA(NGS2,L)=-SADA(NGS,L)
160 CONTINUE
170 CONTINUE
DO 172 I=1,INT
J=I+INT
SAVE(I)=PT(I)
SAVE(J)=BEXT(I)
172 CONTINUE
BSCA=PT(JAY)*BEXT(JAY)
NBOT=NEWLEV(1)
NTOP=NEWLEV(KTHK+1)
DTAUAB=TAUABS(NBOT,JAY)-TAUABS(NTOP,JAY)
BABS=DTAUAB/THKNS(KTHK)
BEXT(JAY)=BEXT(JAY)+BABS
PT(JAY)=BSCA/BEXT(JAY)
PRINT 175,JAY,PT(JAY),BEXT(JAY)
OPTH(KTHK)=THKNS(KTHK)
GN=FLOAT(NLYR)
OPTHI=OPTH(KTHK)/GN
DO 180 I=1,NLYR
TAU(I)=DBLE(FLOAT(I))*OPTHI*BEXT(JAY)
180 CONTINUE
C
C VALUES OF W OF CHARACTERISTIC EQ. FOR SEARCHING THE EIGENVALUES.
C
DO 200 I=1,NG
DO 200 K=1,NG
W(I,K)=0.000
DO 190 LL=1,NG
W(I,K)=W(I,K)+PINO(LL)*POLI(I,LL)*SADA(K,LL)/(1.000+(UM(I))*RT(K))
190 CONTINUE
200 CONTINUE
C
C MAXIMUM DIMENSIONS OF SOLUTION MATRIX = NGL
NGL=NLYR+NG
DO 210 I=1,NGL
DO 210 K=1,NGL
W1(I,K)=0.000
210 CONTINUE
C
C FOR W(-U)---CLOUD TOP BOUNDARY CONDITION. USE W(9,K) TO W(16,K)
DO 220 I=1,NG2
II=I+NG2
DO 220 K=1,NG
W1(I,K)=W(II,K)
220 CONTINUE
IF (NLYR.EQ.1) GO TO 240
L1=NLYR-1
DO 230 L=1,L1
DO 230 I=1,NG
II=I+NG2+(L-1)*NG
DO 230 K=1,NG
KK=K+NG*(L-1)
NK=K+NG*L
W1(II,KK)=W(I,K)*DEXP(-RT(K)*TAU(L))
W1(II,NK)=-W1(II,KK)
230 CONTINUE
240 CONTINUE

```

```

C
C FOR W(+U)---CLOUD BASE BOUNDARY CONDITION, USE W(1,K) TO W(8,K)
  DO 250 I=1,N62
    II=I+NLYR*N62+(NLYR-1)*N62
    DO 250 K=1,N6
      KK=K+(NLYR-1)*N6
      W1(II,KK)=W(1,K)*DEXP(-RT(K)*TAU(NLYR))
250  CONTINUE
  DO 260 I=1,N6L
    DO 260 K=1,N6L
      W2(I,K)=W1(I,K)
260  CONTINUE
C
C EVALUATE THE 'L' COEFFICIENTS OF THE SOLUTIONS FOR INTEGRAL-
C DIFFERENTIAL TRANSFER EQUATION.
C THE CLOUD TEMPERATURE GRADIENT IS DETERMINED FROM ATMOSPHERIC PROFILE
C INPUT DATA.
C
  HGT=THKNS(KTHK)
  CALL TC (NLYR,OPTH1,HGT,JAY,KTHK)
  DO 270 I=1,NLYR
    U=NU(JAY)
    BINTC(I)=TEMPC(I)
270  CONTINUE
  DO 280 I=1,N62
    LI(I)=DBLE(DW(I,JAY)-BINTC(I))
C
C THE ARRAY 'LE' IS USED FOR EMISSIVITY CALCULATIONS IN WHICH THE
C UPWARD AND DOWNWARD INTENSITIES (UP,DW) OF THE MOLEC ATMOS
C ARE SET TO ZERO
C
  LE(I)=DBLE(-BINTC(I))
280  CONTINUE
  IF (NLYR.EQ.1) GO TO 300
  DO 290 L=1,L1
    DO 290 I=1,N6
      II=I+N62+(L-1)*N6
      LI(II)=DBLE(BINTC(L+1)-BINTC(L))
      LE(II)=LI(II)
290  CONTINUE
300  CONTINUE
  DO 330 I=1,N62
    IF (NLYR.GT.1) GO TO 310
    II=I+N62
    GO TO 320
310  II=I+N62*(2*NLYR-1)
320  LI(II)=DBLE(UP(I,JAY)-BINTC(NLYR))
    LE(II)=DBLE(-BINTC(NLYR))
330  CONTINUE
C
C COMPUTE UPWARD + DOWNWARD INTENSITIES
C THE MATRIX INVERSION WILL RETURN THE L - COEFFICIENTS
C
  NTOT=N6*NLYR
  CALL SIMU (W1,LI,NTOT,KS)
  CALL SIMU (W2,LE,NTOT,KS)
C
C INTENSITY(I,J) IS FOR THE ITH LAYER - JTH ANGLE 'MU'
  DO 360 I=1,NLYR
    DO 370 J=1,N62
      JJ=J+N62
      INTENS(I,J)=0.0000
      INTENS(I,JJ)=0.0000
      ENTENS(I,J)=0.0000
      ENTENS(I,JJ)=0.0000

```

```

DO 360 K=1,N6
NK=K
AX=DEXP(-RT(K)*TAU(I))
IF (I.NE.1) NK=K+N6*(I-1)
IF (I.GT.1) GO TO 340
INTENS(I,J)=INTENS(I,J)+LI(NK)*W(J,K)
ENTENS(I,J)=ENTENS(I,J)+LE(NK)*W(J,K)
GO TO 350
340 INTENS(I,J)=INTENS(I,J)+LI(NK)*W(J,K)*DEXP(-RT(K)*TAU(I-1))
ENTENS(I,J)=ENTENS(I,J)+LI(NK)*W(J,K)*DEXP(-RT(K)*TAU(I-1))
350 INTENS(I,JJ)=INTENS(I,JJ)+LI(NK)*W(JJ,K)*AX
ENTENS(I,JJ)=ENTENS(I,JJ)+LE(NK)*W(JJ,K)*AX
360 CONTINUE
INTENS(I,J)=INTENS(I,J)+DBLE(BINTC(I))
INTENS(I,JJ)=INTENS(I,JJ)+DBLE(BINTC(I))
ENTENS(I,J)=ENTENS(I,J)+DBLE(BINTC(I))
ENTENS(I,JJ)=ENTENS(I,JJ)+DBLE(BINTC(I))
370 CONTINUE
380 CONTINUE
IEMIS=1
ICCHAD=0
NDUM=1
JDUM=3
C
C EXTRAPOLATE 'ENTENS' (FOR EMISSIVITY) VALUES AT THETA = 00, AND 90
IUPDOWN=0
CALL EXTRPO(THETA,ENTENS,TVERT,NDUM,JDUM,IUPDOWN)
CALL EXTRP9(THETA,ENTENS,THORIZ,NDUM,JDUM,IUPDOWN)
IF(CLUDEM.GT.0.1) EMISS=TVERT/BINTC(1)
C
C EXTRAPOLATE 'INTENS' VALUES AT THETA = 0 AND 90 DEGREES FOR
C TRANSMITTED AND REFLECTED RADIANCE USED IN DETERMINING THE CLOUD
C TRANSMISSIVITY AND REFLECTIVITY
C THE EXTRAPOLATED CLEAR COLUMN RADIANCE(UP) FOR THETA=0 IS 'CUVERT'
C
IEMIS=0
CALL EXTRPO(THETA,INTENS,TVERT,NDUM,JDUM,IUPDOWN)
CALL EXTRP9(THETA,INTENS,THORIZ,NDUM,JDUM,IUPDOWN)
IUPDOWN=1
NDUM=3
CALL EXTRPO(THETA,INTENS,RVERT,NDUM,JDUM,IUPDOWN)
CALL EXTRP9(THETA,INTENS,RHORIZ,NDUM,JDUM,IUPDOWN)
IF(CUVERT(JAY).EQ.0) GO TO 384
TRANS=TVERT/CUVERT(JAY)
REFLECT=RVERT/CUVERT(JAY)
384 THRUPT(JAY)=THRUPT(JAY)
IF(ITER.EQ.0) GO TO 385
IF(CUVERT(JAY).NE.0) GO TO 383
THRUPT(JAY)=2.0
CUVERT(JAY)=RVERT
GO TO 385
383 THRUPT(JAY)=RVERT/CUVERT(JAY)
385 WRITE (6,420) JAY,NU(JAY)
WRITE (6,430) OPTH(KTHK),TAU(NLYR)
WRITE (6,440) TVERT,RHORIZ
DO 390 J=1,N62
JJ=N6+1-J
WRITE (6,450) THETA(J),INTENS(I,J),THETA(JJ),INTENS(NLYR,JJ)
390 CONTINUE

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      YYY=INTENS(1,1)
      WRITE (6,460) THOR1Z,RVERT
      WRITE (6,470) EMISS,TRANS,REFLCT
      IF(CLODEM,GT,0.1) EMISSM=ENTENS(1,1)/BINTC(2)
      WRITE (6,480) EMISSM,BINTC(2)
      CALL TEMTOP(TEMP,NBASE,KTHK,JAY,EMISSM,TVERT,YYY,
1NU(JAY),REFLCT)
      IF(JAY.NE.INT) GO TO 400
      IF(IFLAG(2).EQ.0) GO TO 400
      WRITE(20,1000) LVNUM,IFLAG(2),NBASE,THKNS(KTHK)
      WRITE(20,1010) ((TAUABS(I,J),I=1,LVNUM),J=1,INT)
      WRITE(20,1020) (TEMP(I),I=1,LVNUM)
      WRITE(20,1020) (ALT(I),I=1,LVNUM)
      WRITE(20,1010) ((PANI(I,J),I=1,NG),J=1,INT)
      WRITE(20,1010) (BEXT(I),I=1,INT)
      WRITE(20,1010) (PT(I),I=1,INT)
400  CONTINUE
      DO 405 I=1,INT
      J=I+INT
      PT(I)=SAVE(I)
      BEXT(I)=SAVE(J)
405  CONTINUE
      RETURN
C
175  FORMAT(1H0,'REVISED SINGLE SCAT ALBEDO AND BEXT FOR CHANNEL ',I2,
*  '/', ' SINGLE SCAT ALBEDO=',F5.3,2X,'BEXT=',F5.3)
410  FORMAT (10X,'$$$$$ ITERATION FAILED $$$$ FIFF =',E20.14)
420  FORMAT (//,'---SPECTRAL REGION',I2,'--- NU=',F6.4)
430  FORMAT (32X,'CLOUD THICKNESS=',F6.2,'(KM)    OPTICAL THICKNESS=',
1F8.5,'/,36X,'THETA',6X,'UPWARD INTENSITY',5X,'THETA',5X,'DOWNWARD I
2NTENSITY',/)
440  FORMAT(38X,'0.0000',5X,F13.5,7X,'90.0000',5X,F13.5)
450  FORMAT(36X,F8.4,5X,F13.5,5X,F9.4,5X,F13.5)
460  FORMAT(37X,'90.0000',5X,F13.5,6X,'180.0000',5X,F13.5)
470  FORMAT (/,20X,'EMISSIVITY= ',F9.6,5X,'TRANSMISSIVITY= ',F9.6,5X,'R
1EFLECTIVITY= ',F9.6,/)
480  FORMAT (/,10X,'MEAN EMISSIVITY = ',F9.6,
1'  CORRESPONDING TO T = ',F7.2,'DEG K')
1000 FORMAT(2I2,2F10.5)
1010 FORMAT(350E15.7)
1020 FORMAT(50F10.5)
      END

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      SUBROUTINE TC (LA,OPHI,HGT,JAY,KTHK)
C *****
C THIS SUBROUTINE CALCULATES THE CLOUD LAYER TEMPERATURES FROM THE
C INPUT ATMOSPHERIC TEMPERATURE PROFILE, HEIGHT OF CLOUD BASE, AND THE
C NUMBER OF ISOTHERMAL LAYERS APPROXIMATING THE TEMPERATURE GRADIENT
C *****
C
      PARAMETER LEV=40, INT=6, MZ=25, NG=16, NG2=8, NMOST=LEV+10
      REAL NBASE
      COMMON/INPUT/ALT(I,MOST),TEMP(NMOST),PRE(NMOST),H2O(NMOST),
      * THRS(9),NTHICK,I,BASE,TRN(NMOST,INT),NEWLEV(10)
      * ,TAUABS(NMOST,INT)
      COMMON /CLOUD/ TEMPC(10)
      COMMON/LRMULT/UBC,ATMD,ATMR,ATMA,SFC,THRUSV(INT),THRUPT(INT),ITER
      * ,ECR,IPCHAN(INT),CLDEM
      DIMENSION TEMPNT(10)
C
C ASSUME LINEAR TEMP GRADIENT ....MONOTONICALLY DEC W/ HEIGHT
C
      IUXBOT=NEWLEV(1)
      IUXTOP=NEWLEV(KTHK+1)
      TEMPC(1)=TEMP(IUXTOP)*CLDEM
      TEMPC(LA)=TEMP(IUXBOT)*CLDEM
      DT=TEMPC(LA)-TEMPC(1)
      DZ=FLOAT(LA)-1.
      LA1=LA-1
      DO 10 I=2,LA1
      TEMPC(I)=TEMPC(1)+(DT/DZ)*(FLOAT(I-1))*CLDEM
10  CONTINUE
      IF (JAY.GT.1) GO TO 30
      WRITE (6,40)
      DO 20 I=1,LA
      TEMPNT(I)=TEMPC(I)-273.16
      IF (I.EQ.1) WRITE (6,50) I,TEMPNT(I)
      IF (I.EQ.LA) WRITE (6,60) I,TEMPNT(I)
      IF (I.NE.1.AND.I.NE.LA) WRITE (6,70) I,TEMPNT(I)
20  CONTINUE
30  CONTINUE
      RETURN
C
40  FORMAT (//,4HX,'CLOUD LAYER TEMPERATURES')
50  FORMAT (51X,I1,F7.2,' (TOP)')
60  FORMAT (51X,I1,F7.2,' (BASE)')
70  FORMAT (51X,I1,F7.2)
      END

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      SUBROUTINE TEMTOP(TEMP,NHASE,KTHK,JAY,EMISSH,X,Y,U,REFLECT)
C*****
C THIS SUBROUTINE FINDS THE BRIGHTNESS TEMPERATURE AT
C THE SATELLITE'S POINT OF VIEW IN A CLOUDY ATMOSPHERE
C*****
C
      PARAMETER LEV=40,INT=6, NMOST=LEV+10
      REAL NHASE
      COMMON/INPUT/ALT1(NMOST),TEMP1(NMOST),PRE(NMOST),H2O(NMOST),
      * THICK(9),NTHICK,IBASE1,TRN(NMOST,INT),NEWLEV(10)
      * ,TAUABS(NMOST,INT)
      COMMON/NEWLEV/ADDHT(10),NUMNEW,LVNUM
      DIMENSION TAU(NMOST,INT),TEMP(NMOST),TEMTOP(2),TISFC(2),ANGMOD(2)
      EQUIVALENCE(TRN(1,1),TAU(1,1))
      ROLT=1.38046E-16
      C=2.997929E10
      CONVKT=2.*ROLT*C*(U**2)
      ANGMOD(1)=1.0
      ANGMOD(2)=1.01110C
      TISFC(1)=X
      TISFC(2)=Y
      LEV1=LVNUM
      IF(JAY.NE.1) GO TO 60
80 ISFC=NEWLEV(KTHK+1)
      ISFCF1=ISFC+1
      DO 100 L=1,2
      H1=TEMP(ISFC)
      TAU1=TAU(ISFC,JAY)*ANGMOD(L)
      TEMTOP(L)=0.0
      DO 90 K=ISFCF1,LEV1
      H2=TEMP(K)
      TAU2=TAU(K,JAY)*ANGMOD(L)
      TEMTOP(L)=TEMTOP(L)+0.5*(H1+H2)*(TAU2-TAU1)
      TAU1=TAU2
90 H1=H2
      HS=TISFC(L)*TAU(ISFC,JAY)
100 TEMTOP(L)=TEMTOP(L)+HS
      WRITE(6,900) TEMTOP(1),TEMTOP(2)
900 FORMAT (/,10X,' VERT. R. TEMP. =',F12.6,10X,' B. TEMP. AT 8.5 =',
      * F12.6,/)
      RETURN
      END

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```

      SUBROUTINE USADA (Y,X,N)
C*****
C GENERATION OF SADA FUNCTION
C*****
C
      DOUBLE PRECISION Y,X,PINO,FI,YI1
      COMMON /POLY/ PINO(16)
      DIMENSION Y(1)
      Y(1)=1.000
      IF (N) 10,10,20
10 RETURN
20 Y(2)=- (1.00-PINO(1))/X
      IF (N-1) 10,10,30
30 DO 40 I=2,N
      A=I
      FI=USLE(A)
      YI1=(FI-1.000)/FI*Y(I-1)
      Y(I+1)=- (2.000*(FI-1.000)+1.000-PINO(I))/(FI*X)*Y(I)-YI1
40 CONTINUE
      RETURN
      END

```

```

      SUBROUTINE PFILE
C
C WRITE ATMOSPHERE PROFILES TO FILE 20
C
      PARAMETER LEV=40, NUMCHN=6, NMOST=LEV+10
      COMMON/INPUT/H(NMOST),T(NMOST),P(NMOST),H20(NMOST),THK(9),
      * I,THICK,NHASE,TRN(NMOST,NUMCHN),NEWLEV(10)
      COMMON/FREQ/FNU(NUMCHN)
      COMMON/NEWLEV/ADDIT(10),NUMNEW,LVNUM
      COMMON/SFCEMS/EMIS(NUMCHN)
      DIMENSION TR(NMOST,NUMCHN),HT(NMOST),TP(NMOST)

      ISTAT=NERTRN(6,'DASG,AX 20.,F/O/TRK/100 . ')
      PRINT 200,ISTAT
      DO 50 I=LVNUM,1,-1
      K=LVNUM-I+1
      H(K)=H(1)
      T(K)=T(1)
      DO 50 J=1,7
      L=8-J
      IF(J.GT,4) L=4-J-J*INT(5/J)
      TR(K,L)=TRN(I,J)
50 CONTINUE

      PRINT 400,(FNU(J),J=1,NUMCHN)
      DO 100 I=1,LVNUM
      WRITE(20,300),HT(I),TP(I),H20(I),(TR(I,J),J=1,NUMCHN)
      PRINT 500,HT(I),T(I),H20(I),(TR(I,J),J=1,NUMCHN)
100 CONTINUE
200 FORMAT(1H1,'STATIS OF ASSIGN FOR FILE 20. = ',I3)
300 FORMAT(F7.3,F6.1,F7.3,2X,F13.6)
400 FORMAT(1H1,'HEIGHT TEMP H2O ',2X,7(' TAU(',F5.2,') '))
500 FORMAT(1X,F7.3,F6.1,F7.3,2X,F13.6)
      RETURN
      END

```

BLOCK DATA

 C DATA OF ANTENNA GAIN OF RADIOMETER, AND CLIMATOLOGICAL PROFILES
 C OF TEMPERATURE, HUMIDITY, AND HEIGHT.

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PARAMETER NUMLV=40, NMOST=NUMLV+10, NUMCHN=6
COMMON/PRES/ P(NMOST)
COMMON/FREQ/FNU(NUMCHN)
COMMON/TANDW/T(NMOST),W(NMOST)
COMMON/HANDU/H(NMOST),U(NMOST)
COMMON/GAIN/GANT(127,NUMCHN)
DATA (P(I),I=1,NUMLV)/
* .1.,.2.,.5.,1.,1.5.,2.,3.,4.,5.,7.,10.,15.,20.,25.,30.,50.,60.,
* 70.,85.,100.,115.,135.,150.,200.,250.,300.,350.,400.,430.,475.,
* 500.,570.,620.,670.,700.,780.,850.,920.,950.,1000./
DATA FNU/22.235,31.65,52.85,53.85,55.45,37.00/
DATA(GANT(I,1),I=1,127)/ 63*0.,
* 96.5, 93.7, 85.3, 78.9, 63.9, 51.7, 38.5, 28.2, 18.9,
* 11.7, 6.12, 1.53, 1.39, 3.21, 3.60, 3.64, 2.51, 1.33,
* 0.065, 0.889, 1.43, 1.31, 1.05, 0.355, 0.124, 0.707, 0.9,
* 0.846, 0.556, 0.431, 0.695, 0.875, 0.984, 0.873, 0.703, 0.455,
* 0.279, 0.364, 0.584, 0.707, 0.865, 0.733, 0.549, 0.361, 0.153,
* 0.125, 0.321, 0.539, 0.565, 0.513, 0.457, 0.309, 0.279, 0.215,
* 0.347, 0.409, 0.520, 0.544, 0.621, 0.594, 0.645, 0.539, 0.477,
* 0.422 /

DATA(GANT(I,2),I=1,127)/ 63*0.,
* 96.5, 93.7, 85.3, 78.9, 63.9, 51.7, 38.5, 28.2, 18.9,
* 11.7, 6.12, 1.53, 1.39, 3.21, 3.60, 3.64, 2.51, 1.33,
* 0.065, 0.889, 1.43, 1.31, 1.05, 0.355, 0.124, 0.707, 0.9,
* 0.846, 0.556, 0.431, 0.695, 0.875, 0.984, 0.873, 0.703, 0.455,
* 0.279, 0.364, 0.584, 0.707, 0.865, 0.733, 0.549, 0.361, 0.153,
* 0.125, 0.321, 0.539, 0.565, 0.513, 0.457, 0.309, 0.279, 0.215,
* 0.347, 0.409, 0.520, 0.544, 0.621, 0.594, 0.645, 0.539, 0.477,
* 0.422 /

DATA(GANT(I,3),I=1,127)/ 63*0.,
* 96.5, 93.7, 85.3, 78.9, 63.9, 51.7, 38.5, 28.2, 18.9,
* 11.7, 6.12, 1.53, 1.39, 3.21, 3.60, 3.64, 2.51, 1.33,
* 0.065, 0.889, 1.43, 1.31, 1.05, 0.355, 0.124, 0.707, 0.9,
* 0.846, 0.556, 0.431, 0.695, 0.875, 0.984, 0.873, 0.703, 0.455,
* 0.279, 0.364, 0.584, 0.707, 0.865, 0.733, 0.549, 0.361, 0.153,
* 0.125, 0.321, 0.539, 0.565, 0.513, 0.457, 0.309, 0.279, 0.215,
* 0.347, 0.409, 0.520, 0.544, 0.621, 0.594, 0.645, 0.539, 0.477,
* 0.422 /

DATA(GANT(I,4),I=1,127)/ 63*0.,
* 96.5, 93.7, 85.3, 78.9, 63.9, 51.7, 38.5, 28.2, 18.9,
* 11.7, 6.12, 1.53, 1.39, 3.21, 3.60, 3.64, 2.51, 1.33,
* 0.065, 0.889, 1.43, 1.31, 1.05, 0.355, 0.124, 0.707, 0.9,
* 0.846, 0.556, 0.431, 0.695, 0.875, 0.984, 0.873, 0.703, 0.455,
* 0.279, 0.364, 0.584, 0.707, 0.865, 0.733, 0.549, 0.361, 0.153,
* 0.125, 0.321, 0.539, 0.565, 0.513, 0.457, 0.309, 0.279, 0.215,
* 0.347, 0.409, 0.520, 0.544, 0.621, 0.594, 0.645, 0.539, 0.477,
* 0.422 /

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```

DATA (GANT(I,5), I=1,127) / 63*0.,
* 96.5, 93.7, 85.3, 78.9, 63.9, 51.7, 38.5, 28.2, 18.9,
* 11.7, 6.12, 1.53, 1.39, 3.21, 3.60, 3.64, 2.51, 1.33,
* 0.065, 0.889, 1.43, 1.31, 1.05, 0.355, 0.124, 0.707, 0.9,
* 0.846, 0.556, 0.431, 0.695, 0.875, 0.984, 0.873, 0.703, 0.455,
* 0.279, 0.364, 0.584, 0.707, 0.865, 0.733, 0.549, 0.361, 0.153,
* 0.125, 0.321, 0.539, 0.565, 0.513, 0.457, 0.309, 0.279, 0.215,
* 0.347, 0.409, 0.520, 0.544, 0.621, 0.594, 0.645, 0.539, 0.477,
* 0.422 /

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```

DATA (GANT(I,6), I=1,127) / 63*0.,
* 96.5, 93.7, 85.3, 78.9, 63.9, 51.7, 38.5, 28.2, 18.9,
* 11.7, 6.12, 1.53, 1.39, 3.21, 3.60, 3.64, 2.51, 1.33,
* 0.065, 0.889, 1.43, 1.31, 1.05, 0.355, 0.124, 0.707, 0.9,
* 0.846, 0.556, 0.431, 0.695, 0.875, 0.984, 0.873, 0.703, 0.455,
* 0.279, 0.364, 0.584, 0.707, 0.865, 0.733, 0.549, 0.361, 0.153,
* 0.125, 0.321, 0.539, 0.565, 0.513, 0.457, 0.309, 0.279, 0.215,
* 0.347, 0.409, 0.520, 0.544, 0.621, 0.594, 0.645, 0.539, 0.477,
* 0.422 /

```

```

DATA (T(I), I=1, NUMLV) /
* 228.2, 244.2, 256.7, 266.2, 262.2, 258.2, 251.2, 246.2, 242.2,
* 235.2, 231.2, 227.2, 223.2, 221.2, 219.2, 214.2, 215.2, 212.2,
* 211.2, 209.2, 216.3, 216.3, 216.3, 218.4, 225.5, 233.4, 240.4,
* 251.7, 255.4, 259.9, 262.2, 268.4, 272.4, 276.2, 278.3, 283.6,
* 287.5, 290.7, 291.8, 293.6 /

```

```

DATA (W(I), I=1, NUMLV) /
* 20*0., .6669E-3,
* 0.8156E-3, .9703E-3, .4807E-2, .3021E-1, .8973E-1, .1703E-0, .2957E00,
* 0.3668E00, .5653E00, .6897E00, .1204E+1, .1808E+1, .2639E+1, .3138E+1,
* 0.5306E+1, .7582E+1, .1010E+2, .1140E+2, .1348E+2 /

```

```

DATA (H(I), I=1, NUMLV) /
* 66.25, 61.28, 54.30, 48.85, 45.60, 43.35, 40.24, 38.27,
* 37.16, 34.14, 31.79, 28.93, 27.10, 25.53, 24.20, 20.89,
* 19.71, 16.73, 17.48, 16.44, 15.55, 14.54, 13.88, 12.08,
* 10.59, 9.358, 8.268, 7.465, 6.930, 6.186, 5.796, 4.773,
* 4.102, 3.472, 3.116, 2.228, 1.505, .8297, .5533, .1113 /

```

END